

# GeoScience Abstracts

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June 1960

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## GEOSCIENCE ABSTRACTS

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## SERIALS

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 Explosives Engineer. Wilmington, Delaware.  
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 Oklahoma Geological Survey, Bulletin; Circular. Norman.  
 Oklahoma Geology Notes (Oklahoma Geological Survey). Norman.  
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 Pennsylvania State University, Mineral Industries Experiment Station, Bulletin. University Park.  
 Quebec (Province), Dept. of Mines, Geological Report. Quebec.  
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 Virginia, Division of Mineral Resources, Bulletin; Mineral Resources Report. Charlottesville.  
 World Oil. Houston, Texas.

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Those wishing to purchase items abstracted herein should address their orders to the agency, society, or organization indicated in the bibliographic citations preceding the abstracts, or to their local book dealer. The city and state for the serials cited are given above. The American Geological Institute, publisher of GeoScience Abstracts, regrets that it cannot fill purchase orders for abstracted publications other than its own.



# GeoScience Abstracts

## 1. GEOLOGIC MAPS, AREAL AND REGIONAL GEOLOGY

### PART 1. GEOLOGIC MAPS

See also: Structural Geology 2-1396.

2-1309. Mountjoy, Eric Walter. GEOLOGY, NIJETTE, WEST OF FIFTH MERIDIAN, ALBERTA: Canada, Geol. Survey, Map 40-1959, Scale 1:63,360, marginal notes, 1960, 2 refs.

Preliminary series geological map of some 260 sq. mi. in the foothills and front ranges of the central Alberta Rocky Mountains, just E. of Jasper. Several V.-dipping thrust sheets, some of which are folded and faulted, occur, and the whole area has a strongly developed northwesterly trend of parallel structural pattern that is reflected in the exposure of the rock units. Cambrian, Ordovician, Devonian, Mississippian, Permian and possibly Pennsylvanian, Triassic, Jurassic, and Cretaceous marine rocks are exposed. Coal seams occur in the Lower Cretaceous Luscar formation, and structures suitable for the accumulation of gas and oil may be present at depth.--P. Harker.

2-1310. Fyles, John G. SURFICIAL GEOLOGY, CUSTER RIVER, COMOX, NANAIMO AND SAYWARD DISTRICTS, BRITISH COLUMBIA: Canada, Geol. Survey, Map 49-1959, scale 1:63,360, 1960.

Preliminary series map covering about 380 sq. mi. near the midlength of the NE. shore of Vancouver Island. Deposits of the last regional glaciation (Vashon drift) are underlain by the Quadra sediments of fluvial origin that record a major interstadial interval. There are also some other extensive deposits beneath the Vashon ground moraine whose relationship to Quadra sediments is not known. Marine and fluvial deposits overlying the Vashon drift are designated as the Capilano sediments and the Salish sediments. The Capilano sediments accumulated during an early postglacial interval when sea level was higher than at present; the Salish deposits were deposited under modern conditions and are still being formed.--P. Harker.

2-1311. Anderson, Francis D., and W.H. Poole. GEOLOGY, WOODSTOCK-FREDERICTON, YORK, CARLETON, SUNBURY AND NORTHUMBERLAND COUNTIES, NEW BRUNSWICK: Canada, Geol. Survey, Map 37-1959, scale 1:126,720, marginal notes, 1959, 13 refs.

Preliminary series geological map covering about 1,000 sq. mi. of southwestern New Brunswick, N. and W. of Fredericton. The area embraces the southwestern end of a belt of granitic batholiths and deformed early Paleozoic rocks that extends northwesterly from Maine to Bathurst. Intensely deformed sedimentary and volcanic strata of Ordovician, Silurian, and Devonian age have been intruded by granitic batholiths and overlain unconformably by undeformed coarse clastic sediments and few volcanic rocks of Carboniferous age. A central southwesterly trending belt of Ordovician rocks is flanked by Silurian strata on the SE. and Silurian and Devonian strata on the NW. The nature of the Ordovician-Silurian contact is uncertain. Ferruginous and manganeseiferous strata occur in both Ordovician and Silurian systems. Those of Silurian age near Woodstock are being investigated with a view to producing ferro-manganese. Base metal deposits occur sporadically in Ordovician strata. Stibnite has been mined from quartz veins cutting Silurian strata.

Radioactive minerals were recently found in Carboniferous rhyolitic rocks. Coal prospects occur in Pennsylvanian sediments.--W.H. Poole.

2-1312. MacLaren, A.S. AEROMAGNETIC MAP, WHOLDAIA LAKE EAST, DISTRICT OF MACKENZIE, NORTHWEST TERRITORIES: Canada, Geol. Survey, Map 1080A, scale 1:253,440, marginal notes, 1959, 4 refs

Aeromagnetic map covering the area 60°-61°N., 102°-104°W. This is a composite map compiled from 16 aeromagnetic maps previously published by the Geological Survey of Canada on a scale of 1 in. to 1 mi. Flight line traces are eliminated, and base map detail is generalized and reduced to a minimum. Flights were made at an altitude of 1,000 ft. above ground level.--P. Harker.

2-1313. Hughes, Owen L. SURFICIAL GEOLOGY, IROQUOIS FALLS, COCHRANE DISTRICT, ONTARIO: Canada, Geol. Survey, Map 46-1959, scale 1:126,720, marginal notes, 1960.

Preliminary series map covering about 800 sq. mi. The area lies due N. of Toronto and just W. of Timmins. Wisconsin ice covered the area and deposited glacial till containing lenses of gravel. Concurrent with northward retreat of the ice margin, the area was occupied by glacial lake Barlow-Ojibway. The lake extended from the S. end of Lake Timiskaming northward across the continental divide which was then depressed below the level of the outlet through Timiskaming gorge. Subglacial streams, entering into the lake at the ice margin, deposited sand and gravel, forming elongate esker complexes. Silt and clay carried by the subglacial streams were deposited on the lake bottom as annual layers or varves. Northward retreat of the ice margin caused shallowing of the glacial lake, and bedrock knobs appeared as islands and were swept free of drift by wave erosion. Subsequent southward expansion of a lobe of ice deposited the Cochrane till. Thin deposits of lacustrine sediments were deposited in a shallow glacial lake or lakes which persisted along the margin of the lobe as it retreated after a brief stand.--P. Harker.

2-1314. Canada, Geological Survey and Ontario, Dept. of Mines. ANENIMUS RIVER, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 890, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat. 50°30'-50°45'N., long. 91°30'-92°W., 1960.

2-1315. Canada, Geological Survey and Ontario, Dept. of Mines. OTATAKAN LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 891, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat. 50°45'-51°N., long. 91°30'-92°W., 1960.

2-1316. Canada, Geological Survey and Ontario, Dept. of Mines. WESLEYAN LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 892, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat. 51°-51°15'N., long. 91°30'-92°W., 1960.

2-1317. Canada, Geological Survey and Ontario, Dept. of Mines. ZIONZ LAKE, KENORA DISTRICT,



ONTARIO: Canada, Geol. Survey, Geophysics Paper 893, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}15' - 51^{\circ}30'N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1318. Canada, Geological Survey and Ontario, Dept. of Mines. CAT LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 894, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}30' - 51^{\circ}45'N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1319. Canada, Geological Survey and Ontario, Dept. of Mines. WHITESTONE LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 895, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}45' - 52^{\circ}N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1320. Canada, Geological Survey and Ontario, Dept. of Mines. McCAULEY LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 896, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ} - 52^{\circ}15'N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1321. Canada, Geological Survey and Ontario, Dept. of Mines. SHINBONE LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 897, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}15' - 52^{\circ}30'N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1322. Canada, Geological Survey and Ontario, Dept. of Mines. WINDIGO LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 898, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}30' - 52^{\circ}45'N.$ , long.  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1323. Canada, Geological Survey and Ontario, Dept. of Mines. NIKIP LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 899, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}45' - 53^{\circ}N.$ ,  $91^{\circ}30' - 92^{\circ}W.$ , 1960.

2-1324. Canada, Geological Survey and Ontario, Dept. of Mines. ST. RAPHAEL LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 900, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $50^{\circ}30' - 50^{\circ}45'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1325. Canada, Geological Survey and Ontario, Dept. of Mines. LINDBERGH LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 901, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $50^{\circ}45' - 51^{\circ}N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1326. Canada, Geological Survey and Ontario, Dept. of Mines. BLACKSTONE LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 902, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale

1:63,360, lat.  $51^{\circ} - 51^{\circ}15'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1327. Canada, Geological Survey and Ontario, Dept. of Mines. OBASKAKA LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 903, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}15' - 51^{\circ}30'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1328. Canada, Geological Survey and Ontario, Dept. of Mines. GITCHE RIVER, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 904, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}30' - 51^{\circ}45'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1329. Canada, Geological Survey and Ontario, Dept. of Mines. UPTURNEDROOT LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 905, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}45' - 52^{\circ}N.$ ,  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1330. Canada, Geological Survey and Ontario, Dept. of Mines. HINTON LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 906, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ} - 52^{\circ}15'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1331. Canada, Geological Survey and Ontario, Dept. of Mines. STIRLAND LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 907, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}15' - 52^{\circ}30'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1332. Canada, Geological Survey and Ontario, Dept. of Mines. YOYOH LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 908, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}30' - 52^{\circ}45'N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1333. Canada, Geological Survey and Ontario, Dept. of Mines. WEAGAMOW LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 909, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}45' - 53^{\circ}N.$ , long.  $91^{\circ} - 91^{\circ}30'W.$ , 1960.

2-1334. Canada, Geological Survey and Ontario, Dept. of Mines. DE LESSEPS LAKE, THUNDER BAY AND KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 910, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $50^{\circ}30' - 50^{\circ}45'N.$ , long.  $90^{\circ}30' - 91^{\circ}W.$ , 1960.

2-1335. Canada, Geological Survey and Ontario, Dept. of Mines. MINISS LAKE, THUNDER BAY AND KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 911, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $50^{\circ}45' - 51^{\circ}N.$ , long.  $90^{\circ}30' - 91^{\circ}W.$ , 1960.



## GEOLOGIC MAPS, AREAL AND REGIONAL GEOLOGY

2-1336. Canada, Geological Survey and Ontario, Dept. of Mines. LAKE ST. JOSEPH WEST, KENORA AND THUNDER BAY DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 912, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}0'-51^{\circ}15'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1337. Canada, Geological Survey and Ontario, Dept. of Mines. KAWINOGANS LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 913, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}15'-51^{\circ}30'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1338. Canada, Geological Survey and Ontario, Dept. of Mines. DOBIE RIVER, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 914, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}30'-51^{\circ}45'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1339. Canada, Geological Survey and Ontario, Dept. of Mines. OBABIKA LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 915, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $51^{\circ}45'-52^{\circ}N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1340. Canada, Geological Survey and Ontario, Dept. of Mines. KECHEOKAGAN LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 916, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}-52^{\circ}15'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1341. Canada, Geological Survey and Ontario, Dept. of Mines. MAWLEY LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 917, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}15'-52^{\circ}30'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1342. Canada, Geological Survey and Ontario, Dept. of Mines. DONNELLY RIVER, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 918, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}30'-52^{\circ}45'N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1343. Canada, Geological Survey and Ontario, Dept. of Mines. NORTH CARIBOU LAKE, KENORA DISTRICT, ONTARIO: Canada, Geol. Survey, Geophysics Paper 919, aeromagnetic map, contour intervals 10, 20, 100, and 1,000 gammas, scale 1:63,360, lat.  $52^{\circ}45'-53^{\circ}N.$ , long.  $90^{\circ}30'-91^{\circ}W.$ , 1960.

2-1344. Karrow, P.F. SURFICIAL GEOLOGY, FRONZINES, CHAMPLAIN, PORTNEUF, LOTBIN-ERE, AND NICOLET COUNTIES, QUEBEC: Canada Geol. Survey, Map 41-1959, scale 1:63,360, marginal notes, 1959.

Preliminary series map covering about 380 sq. mi. on the N. shore of the St. Lawrence River. Earliest Pleistocene sediments are varved clays in a pro-

glacial lake. Ice of possible Illinoian age overrode the varved clays, depositing Becancour till. Following a nonglacial interval, Wisconsin ice disrupted drainage and formed a second proglacial lake, lake Deschaillons. After deposition of the Gentilly till, the ice retreated, allowing marine waters of the Champlain sea to flood the depressed lowland.--P. Harker.

2-1345. Jewett, John M., GRAPHIC COLUMN AND CLASSIFICATION OF ROCKS IN KANSAS: Kansas, State Geol. Survey, chart (columnar sec.), scale (vert.) approx. 1 in. to 75 ft., 1959, pub. 1960.

A revision of a chart published in 1952. Chart includes complete generalized section of surface and subsurface rocks of Kansas. Notes in italics call attention to significant deviations from average lithology or thickness. Section is arranged in 5 columns on 25 x 38 in. sheet.--R. H. King.

2-1346. Deasy, George F., and Phyllis R. Griess. ATLAS OF PENNSYLVANIA COAL AND COAL MINING: PART I, BITUMINOUS COAL: Pennsylvania State Univ., Mineral Industries Expt. Sta., Bull. no. 73, 68 p., maps, Dec. 1959.

This paper-backed atlas brings together selected data on Pennsylvania's bituminous coal resources and coal-mining industry and presents them in graphic form. Included are maps dealing with geological, historical, engineering, economic, production, employment, safety, and transportation facets of the subject.

The approximately 100 maps, most of which are drawn to scales of 1 in. to 32 or 50 mi. are arranged in 6 groups. Introductory maps represent Pennsylvania's relative position among states in the mineral and coal-mining industries of the United States. A second group of maps deals with Pennsylvania's bituminous coal fields as well as selected coal seams within those fields, and a third group analyzes the composition and heat value of the state's bituminous coal. The fourth group of maps is devoted to employment and labor factors in the bituminous industry, and the fifth is concerned with mines, mining methods, and accidents. A final group represents the production and shipment of coal from the state's coal fields.

The authors plan to prepare a companion volume to the above atlas that will deal with anthracite in Pennsylvania.--J. D. Ridge.

### PART 2. AREAL AND REGIONAL GEOLOGY

See also: Structural Geology 2-1393; Stratigraphy 2-1420; Sedimentary Petrology 2-1564; Geohydrology 2-1577, 2-1578; Mineral Deposits 2-1585.

2-1347. Ross, Stewart H. CÉLORON-CARQUEVILLE AREA, ABITIBI-EAST AND ABITIBI-WEST ELECTORAL DISTRICTS: Quebec, Dept. Mines., Geol. Rept. 89, 18 p., 2 illus., on pl., geol. map (in pocket), scale 1:63,360, 1959, 8 refs.

This area ( $49^{\circ}-49^{\circ}15'N.$   $78^{\circ}20'-78^{\circ}45'W.$ ), mapped in 1951, is situated within the clay belt of the Abitibi region, and bedrock exposures are rare. All consolidated rock is Precambrian. About 60% of the area is underlain by Keewatin volcanic, intrusive, and sedimentary rocks and the remainder by later intrusive rocks. The southern and western quarters of the area are underlain mainly by Keewatin rocks, and the northwestern third by later granite. The Keewatin rocks are mainly andesitic to basaltic



lavas, with associated gabbroic to dioritic intrusive rocks. Agglomerates and tuffaceous sedimentary rocks are present locally. The granite in the northwestern part of the area is part of a large mass lying mainly to the N. of the area, between Harricana and Turgeon rivers. A syenite intrusion towards the southwestern corner of the area is part of a mass that extends some 4 mi. to the S. Other minor intrusions consist of diorite and granodiorite; the latest, referred to the Keweenawan, are of diabase and gabbro. Hybrid rocks, consisting of injection gneiss, contact breccia, and diorite, occur along the contacts between the northwestern granite mass and the Keewatin volcanics.

Some prospecting has been done in various parts of the area since 1940, and some copper showings have been found.--From auth., p. 1, 4, 13.

2-1348. Tiphane, Marcel. CHASTE-MAZARIN AREA, ABITIBI-EAST ELECTORAL DISTRICT: Quebec, Dept. Mines, Geol. Rept. 88, 15 p., 8 illus. on 4 pls., geol. map (in pocket), scale 1:63,360, 1959, 12 refs.

The map-area ( $49^{\circ}0'-49^{\circ}15'N$ ,  $77^{\circ}45'-78^{\circ}20'W$ .), covering 550 sq. mi., was examined in 1948 and 1949. It may be divided into 2 parts along a N.-S. line near the Dalet-Maizerets and Mazarin-Glendale township boundaries. E. of this line outcrops are scarce and scattered, and most contacts can only be approximated. They are more abundant along Harricana River, Coigny River, and one of its tributaries, and on low hillocks in the central and north-eastern parts of the area. Outcrops are more common in the second or western part, a belt 5 or 6 mi. wide towards the western boundary of the area.

Beaches composed of rounded pebbles of different rock types were observed along the slope of a few hills. Their elevations, according to barometric measurements, is 150 ft. above the level of Octave River.

All the consolidated rocks are Precambrian. Keewatin-type volcanic rocks and slates, the latter probably of sedimentary origin, occupy about 85% of the area and are intruded by younger granitic masses and by diabase dikes. Feldspar porphyry, aplite, and pegmatite dikes are also found in places. Peridotite occurs in the N.-central part of the area and is assumed to be postvolcanic in age.

Granitic rocks form a stocklike mass near the northwestern corner, cropping out over an area of about  $4\frac{1}{2}$  sq. mi., and a batholithic mass near and along the southern boundary in a zone up to 3 mi. wide. The batholithic mass seems to be the extension of the "Bernetz gneiss." A small area underlain by granitic rocks diagonally cuts the NE. corner and appears to join with the granitic masses of Harricana and Bell rivers.--From auth., p. 4-5.

During 1957 and 1958, a group of mining companies and syndicates covered most of the area by airborne magnetic and electromagnetic surveys supplemented, in some cases, by ground surveys. A limited amount of diamond drilling was also done, the results of which appear to have been inconclusive.--J. E. Gilbert.

2-1349. Churkin, Michael, Jr., and Ralph L. Ingenheim, Jr. SILURIAN STRATA OF THE KLAMATH MOUNTAINS, CALIFORNIA: Am. Jour. Sci., v. 258, no. 4, p. 258-273, 2 illus., map, 4 secs., Apr. 1960, 11 refs.

A eugeosynclinal suite comprising shale, volcanic graywacke, quartz arenite, bedded chert, limestone,

and tuff crops out along Willow Creek, W. of Gatzelle, California. The limestone and associated interbeds of shale contain Silurian trilobites, brachiopods, and corals. The bulk of the sequence, however, consists of unfossiliferous sedimentary rocks beneath the limestone. Trilobites in the limestone belong to genera previously reported from eastern North America, but the assemblage appears most closely related to Middle and Upper Silurian faunas of Alaska and Asia.

Silurian sedimentary rocks are succeeded in thrust fault contact by phyllite and semischist in the western part of the area. These metamorphic rocks have been derived from sedimentary rocks similar to those of the Silurian sequence chiefly by regional cataclastic deformation. Small bodies of diorite and quartz diorite and a large pyroxenite-peridotite pluton intrude both the sedimentary and metamorphic rocks. In addition Tertiary and/or Quaternary basalt flows and alluvium locally truncate all rocks.

Relatively open folds and high-angle faults generally trending N. characterize structural patterns within the sedimentary rocks in marked contrast to a complex pattern of crumpling and shearing in the strongly lineated metamorphic rocks. In addition the thrust plane is deformed by broad open folds plunging gently NW. Thus it is possible to infer at least 3 periods of diastrophism in this area.--Auth.

2-1350. Woodford, A.O. BEDROCK PATTERNS AND STRIKE-SLIP FAULTING IN SOUTHWESTERN CALIFORNIA: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 400-417, map, 1960, 59 refs.

In southwestern California the pre-Turonian bedrock may be divided into 8 units, which may then be used to distinguish 3 rock belts. One belt comprises the glaucophane-bearing rocks near the shore, a second is characterized by quartz-rich graywackes, and the third, which is farthest NE., contains limestones. The second and third belts include Mesozoic volcanics and extensive Cretaceous quartz plutonites. The bedrock mass has been disrupted by faults, some of which have been the loci of strike slips measurable in tens of miles. Much greater displacements have also been suggested. The simplicity of these suggestions seem inconsistent with the distribution of the extrusive and intrusive Mesozoic igneous rocks. Perhaps on some faults 2 slip solutions are possible. In such cases tentative choices may be made by the use of the rule *Disjunctiones minimae, disjunctiones optimae*.--Auth.

2-1351. Vanderpool, Robert E. GEOLOGY OF THE FEATHERSTON AREA, PITTSBURG COUNTY, OKLAHOMA: Oklahoma Geol. Survey, Circ. 53, 36 p., 9 illus., 2 maps (col. geol. map in pocket), 1960, 33 refs.

Rocks of the upper part of the McAlester formation, the Savanna formation, and the lower part of the Boggy formation [all Pennsylvanian] crop out in the area. Three anticlines and 2 synclines are the major structural features. Coal has been produced and is now of little importance. Small gas fields are yielding natural gas from the Hartshorne sandstone [Pennsylvanian].--Auth.

2-1352. Gould, Wilburn J. GEOLOGY OF THE NORTHERN NEEDLE RANGE, MILLARD COUNTY, UTAH: Brigham Young Univ., Dept. Geology, Brig-



## GEOMORPHOLOGY

am Young Univ. Research Studies, Geology Ser.,  
6, no. 5, 47 p., 8 pls. incl. illus., fold. geol.  
map, sec., Aug. 1959, 32 refs.

The Needle Range lies in the extreme southwest-  
ern part of Millard County, approximately 25 mi. S.  
of Garrison. State Highway 21 bounds the area on  
the N., Antelope Valley on the E., the Millard-  
Weaver county line on the S., and the Utah-Nevada  
state line on the W.

The exposed stratigraphic section, aggregating  
slightly more than 8,000 ft., includes rocks from  
Middle Devonian to lower Permian age. The upper  
part of the Guilmette limestone and Pilot shale, a  
combined thickness of 1,800 ft., represents the  
Devonian system. Mississippian strata aggregate  
more than 3,000 ft. and include the upper part of  
the Pilot shale, Joana limestone, Chairman shale,  
and the Illipah formation. The Pennsylvanian system  
includes the lower 2,000 ft. of Ely limestone. Only  
the basal Wolfcampian part of the Permian system  
is present, represented by 690 ft. of the upper Ely  
limestone and 1,000 ft. of the Arcturus limestone.  
An asymmetrical syncline forms the major  
structural feature of the range. Thrust plates moving  
in a NE. direction have overridden the W. limb.  
In the northwestern part of the range has been involved  
in a major fan fold. The folds and accompanying  
thrust faults probably formed during the various  
phases of the Laramide orogeny. Numerous normal  
faults occur throughout but have only minor dis-  
placement. Flow rock in the SE. indicates a slight  
westward tilting of the region.

Structural complexities possibly make the im-  
mediate area unfavorable for accumulation of gas  
and oil.--Auth.

2-1353. Bushnell, Vivian, ed. SCIENTIFIC  
STUDIES AT FLETCHER'S ICE ISLAND, T-3, 1952-  
1955. VOLUME I: U.S. Air Force, Cambridge Re-  
search Center, Geophysics Research Directorate,  
Geophys. Research Papers, no. 63, 219 p., illus.,  
map, profiles, diags., graphs, tables, Sept. 1959,  
32 refs.

Scientific studies based on field investigations at  
Fletcher's Ice Island T-3 during the period 1952 to  
1955 are reported, with appendixes containing data  
tables. Included are general bathymetry, sub-bottom  
seismic studies, gravity studies, oceanographic  
observations, ice movement studies, micropaleonto-  
logical and lithological analysis of sediment cores,  
general Arctic Ocean biology, detailed studies of  
certain marine animals, tritium measurements of  
ice, and O isotope measurements on sediment cores.  
In addition to the works of Air Force Cambridge  
Research Center scientists and contractors, the series  
includes studies prepared by, and for, several U.S.  
Government agencies.

The following papers are included in this report.

- Bushnell, Vivian. Introduction, p. 1-6.
- Crary, A.P., and Norman Goldstein. Geophys-  
ical Studies in the Arctic Ocean, p. 7-30.
- Worthington, L.V. Oceanographic Observations,  
p. 31-58.
- Browne, Irene M., and A.P. Crary. The Move-  
ment of Ice in the Arctic Ocean, p. 37-49.
- Ericson, David B., and Goata Wollin. Micro-  
paleontology and Lithology of Arctic Sediment Cores,  
p. 51-58.
- Green, Keith E. Ecology of some Arctic For-  
aminifera, p. 59-81.
- Mohr, John L. Marine Biological Work, p. 83-103.
- Paynter, Raymond A., Jr. Birds in the Upper  
Arctic, p. 104.
- Knox, George A. Pelagic and Benthic Polychaetes  
of the Central Arctic Basin, p. 105-114.
- Barnard, J. Laurens. Epipelagic and Under-Ice  
Amphipoda of the Central Arctic Basin, p. 115-152.
- Giletti, B.J., and J. Laurence Kulp. Tritium  
Observations, p. 153-158.
- Ault, Wayne U. Oxygen Isotope Measurements  
on Arctic Cores, p. 159-168.
- General Bibliography, p. 169
- Appendixes
  1. Ocean Depths, p. 172-174.
  2. Geomagnetic Observations, p. 174-180.
  3. Gravity, p. 181-219.

## 2. GEOMORPHOLOGY

See also: Geologic Maps 2-1310, 2-1313, 2-1344; Stra-  
tigraphy 2-1418, 2-1419; Engineering Geology 2-1618,  
2-1619.

2-1354. Tanner, William F. NUMERICAL COM-  
PARISON OF GEOMORPHIC SAMPLES: Science,  
131, no. 3412, p. 1525-1526, May 20, 1960, ref.

The distribution of elevations representing a re-  
gion can be shown as a cumulative frequency curve  
plotted on probability paper. Many elevation distri-  
butions are "zig-zag" curves which can be repre-  
sented conveniently by measures of skewness and  
kurtosis. A plot of skewness versus kurtosis per-  
mits the recognition of 6 major, non-Gaussian forms,  
with countless gradations.--Auth.

2-1355. Rocky Mountain Nature Association.  
GLACIERS IN ROCKY MOUNTAIN NATIONAL PARK:  
p. 9 illus., Estes Park, Colorado, 1959.

A brochure intended for visitors to Rocky Moun-  
tain National Park, Colorado. It describes in non-  
technical terms the action of glaciers, their forma-  
tion and movement. Also given are brief descriptions  
and directions for reaching the following glaciers:

Andrew's, Tyndall, Sprague, Rowe, and Taylor.--  
M. Russell.

2-1356. Johnson, Arthur. U.S. GEOLOGICAL  
SURVEY AND U.S. NATIONAL PARK SERVICE  
GLACIER OBSERVATIONS, GLACIER NATIONAL  
PARK, MONTANA, 1959: U.S. Geol. Survey, Repts.,  
Open-File Ser., no. 507, 23 p., 2 illus., fold. map-  
profiles, 3 tables, March 1960.

Data recorded during the 1959 season on Grinnell  
and Sperry glaciers in Glacier National Park are  
presented. Measurements were made of surface  
changes, movement, recession, precipitation, and  
runoff.--M. Russell.

2-1357. Rutten, M.G. ICE-PUSHED RIDGES,  
PERMAFROST AND DRAINAGE: Am. Jour. Sci.,  
v. 258, no. 4, p. 293-297, 2 maps, Apr. 1960, 8  
refs.

Ice-pushed ridges in regional development are the  
main glacial characteristics of the northwestern  
European plain. They were formed during the Riss  
glaciation and consist of preglacial material pushed



up into long, straight or curved ridges with an imbricate structure. They developed as a consequence of the formation of a thick permafrost layer in front of the advancing ice sheet.

Ice-pushed ridges of such a regional development are absent from the northern plains of the United States. It is thought that a relative difference in drainage pattern - away from the ice front in the U. S. A. and toward the ice front in Europe - can be held responsible. The presence of ice-pushed ridges in northwestern Canada, where again the drainage is toward the ice front, and their absence in the European Würm glaciation, where outward drainage prevailed, corroborates this theory.--Auth.

2-1358. Marcus, Melvin G. PERIODIC DRAINAGE OF GLACIER-DAMMED TULSEQUAH LAKE, BRITISH COLUMBIA: *Geog. Rev.*, v. 50, no. 1, p. 89-106, 6 illus., 2 maps, 2 secs., 4 diags., 3 graphs, 2 tables, Jan. 1960, approx. 15 refs.

Ice-dammed Tulsequah Lake has been subject to periodic (usually annual) summer outbursts beneath the 4 1/2 mi. surface of its glacier barrier for over 50 years. The formation and morphology of this lake type is associated with specific phases of climate and glacier retreat common to the North Pacific coast of North America during the last 100 years. At its maximum dimensions (1910-1920), the lake was 195 m. deep and discharged about 239 billion gallons of water. In July 1958, it was 73 m. deep, discharging about 60 billion gallons in a 4-day period. Bathymetric mapping and geobotanical techniques provided fair accuracy.

Drainage frequency is a direct response to climate and ice barrier height. Drainage mechanics are more complex. Physical conditions prevented close examination of water inlets at the foot of the ice dam. Changes in glacial structure and measurement of escape flow, and observation of outlets at the glacier terminus do, however, indicate an hypothesis. It is believed that ice flotation allows access of the water to remnant subsurface pipes. Although the ice barrier resettles, continuous flow is established, and escaping water enlarges the tunnels by melting. At the completion of drainage the pipes begin to close gradually, and barrier outlets are plugged during the autumn freeze. When the lake refills, the process is repeated.--Auth.

2-1359. Terasmae, Jaan, and Owen L. Hughes. GLACIAL RETREAT IN THE NORTH BAY AREA, ONTARIO: *Science*, v. 131, no. 3411, p. 1444-1446, map, May 13, 1960, 10 refs.

Geological and palynological studies in Ontario and Quebec, supported by radiocarbon dates, suggest that the opening of the North Bay outlet and the initiation of the Stanley-Chippewa stages in the Huron and Michigan basins took place 10,000 to 11,000 years ago.--Auth.

2-1360. Tipton, Merlin J. A NEW GLACIAL DRIFT SHEET IN SOUTH DAKOTA?: *South Dakota Acad. Sci., Proc.*, v. 38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 45-48, map, table, 1959, pub. 1960, 11 refs.

The glacial drift in the Big Sioux River valley of South Dakota between Sioux Falls and Watertown has been previously mapped as deposits of 1 glacial age or subage. Chamberlin, Rothrock and Newcombe, Leverett, and Rothrock and Otton mapped the drift

as Kansan, but Flint mapped it as Iowan. On the basis of topographic evidence discovered in 1958 in this area, it is believed that 2 drifts rather than 1 are present. The northern part of the area is believed to be Iowan(?) in age and the southern part Illinoian(?).--F. V. Steece.

2-1361. Cook, Frank A. SELECTED BIBLIOGRAPHY ON PERIGLACIAL PHENOMENA IN CANADA: Canada, Dept. Mines & Tech. Surveys, Geog. Branch, Bibliog. Ser. 24, 23 p., 1959.

This bibliography presents 93 selective references to periglacial phenomena in Canada in the form of annotations and abstracts.--Auth.

2-1362. Lofgren, Ben E. NEAR-SURFACE LAND SUBSIDENCE IN WESTERN SAN JOAQUIN VALLEY, CALIFORNIA: *Jour. Geophys. Research*, v. 65, no. 3, 1053-1062, 9 figs. incl. 2 illus., 2 maps, diags., graphs, March 1960, 4 refs.

Numerous evidences of land subsidence have been observed on unusual soils in the San Joaquin Valley immediately after the application of water. This subsidence results in extensive settling and cracking of soil along ditches and in irregular, undulating topography in irrigated areas. Experiments on quarter-acre test plots are being made to determine the nature and magnitude of near-surface subsidence on different types of alluvial deposits. After 27 months of observation of test plot B, to Dec. 29, 1958, surface bench marks had settled an average of 10.50 ft.; a bench mark anchored at a depth of 25 ft. had settled 9.44 ft.; a bench mark at 50 ft., 8.13 ft.; a bench mark at 75 ft., 6.13 ft.; a bench mark at 100 ft., 3.93 ft.; and a bench mark at 150 ft., 0.39 ft.

Infiltration at a continuing rate of 0.20 ft. of water per day for a 15-month period of inundation was more than sufficient to fill the pore space beneath the test plot to the wetted front, and considerable lateral spreading of the water occurred. Core samples collected before and during the spreading operation show the increase in density and moisture content of the subsurface deposits. The low field density and extreme dryness of the alluvial deposits apparently are 2 important factors in land subsidence due to application of water.--Auth.

2-1363. Tanner, William F. HELICOIDAL FLOW, A POSSIBLE CAUSE OF MEANDERING: *Jour. Geophys. Research*, v. 65, no. 3, p. 993-995, 2 diags., March 1960, 4 refs.

Sediment-free model streams, suspended from the lower sides of nearly horizontal glass plates, meander in the same way as do alluvial rivers. Meandering cannot, therefore, be ascribed to granular bed load. It is suggested that an alternative causative mechanism is helicoidal flow, which, in turn, is the product of turbulence within the stream, as Einstein and Li have pointed out.--Auth.

2-1364. Dury, G.H. MISFIT STREAMS: PROBLEMS IN INTERPRETATION, DISCHARGE, AND DISTRIBUTION: *Geog. Rev.*, v. 50, no. 2, p. 219-242, map, 2 diags., Apr. 1960, 38 refs.

Misfit streams are shown to be regionally distributed, and river capture is rejected as a general cause of their condition. Large meandering channels beneath the alluvium of misfit streams are



interpreted as former stream beds, largely infilled after a change in regimen caused by a change of climate. Statistical analysis of dimensions of channels, of valley meanders, and of stream meanders is used to indicate, within fairly wide limits, the order of reduction in near-bankfull discharge which is expressed in the misfit condition. Data from very small catchments show that present streams, even in conditions highly favorable to rapid surface runoff, fail to develop meanders comparable in size, area or area, with valley meanders, while data from very large catchments demonstrate the unitary nature of the problem discussed. Tentative or preliminary comments and suggestions are made on dating and on distribution.--Auth.

1-1365. Hack, John T. INTERPRETATION OF EROSIONAL TOPOGRAPHY IN HUMID TEMPERATE REGIONS: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 80-97, 2 maps, profile, sec., 1960, 48 refs.

Since the period 1890 to 1900, the theory of the geographic cycle of erosion has dominated the science of geomorphology and strongly influenced the theoretical skeleton of geology as a whole. Some of the principal assumptions in the theory are unrealistic. The concepts of the graded stream and of lateral planation, although based on reality, are misapplied in an evolutionary development, and it is unlikely that a landscape could evolve as indicated by the theory of the geographic cycle.

The concept of dynamic equilibrium provides a more reasonable basis for the interpretation of topographic forms in an erosionally graded landscape. According to this concept every slope and every channel in an erosional system is adjusted to every other. When the topography is in equilibrium and erosional energy remains the same, all elements of the topography are downwasting at the same rate. Differences in relief and form may be explained in terms of spatial relations rather than in terms of an evolutionary development through time. It is recognized however that erosional energy changes in space as well as time, and that topographic forms evolve as energy changes.

Large areas of erosionally graded topography in humid regions have been considered to be "maturely dissected peneplains." According to the equilibrium theory, this topography is what we should expect as the result of long continued erosion. Its explanation does not necessarily involve changes in base level. Sediments in humid regions and some terraces are so equilibrium forms and commonly occur on a lowland area at the border of an adjacent highland.--Auth.

1-1366. Radforth, Norman W., and Lynda S. Muir. DEFINITIVE MICROFOSSILS PERTINENT TO PHYSIOGRAPHIC DIFFERENCE IN MUSKEG: *Royal Soc. Canada, Trans.*, Sec. 5, v. 53, p. 35-40, illus., table, 1959, 8 refs.

It is commonly difficult to assign muskeg areas to specific categories because of admixture of cover types. Microfossil analysis was made of muskeg areas at Copetown, Ontario, and Fort Churchill,

Manitoba, in an attempt to reveal the primary organization. Spore and pollen analyses were made of samples taken at 2.5-cm. increments. Ponding is a major factor in mixed cover areas.--M. Russell.

2-1367. Plapp, John E., and James P. Mitchell. A HYDRODYNAMIC THEORY OF TURBIDITY CURRENTS: *Jour. Geophys. Research*, v. 65, no. 3, p. 983-992, 2 figs., 2 tables, March 1960, 7 refs.

A theory of steady-state turbidity currents along sloping and horizontal ocean or lake bottoms is developed on the basis of hydrodynamic boundary-layer assumptions. The theory is applied to the Grand Banks turbidity current of 1929; reasonable values of current thickness and sediment content are obtained, as well as satisfactory agreement between observed and calculated velocities.--Auth.

2-1368. Hunkins, Kenneth L., Maurice Ewing, Bruce C. Heezen, and Robert J. Menzies. BIOLOGICAL AND GEOLOGICAL OBSERVATIONS ON THE FIRST PHOTOGRAPHS OF THE ARCTIC OCEAN DEEP-SEA FLOOR: *Limnology & Oceanography*, v. 5, no. 2, p. 154-161, 15 illus., Apr. 1960, 4 refs.

The first series of Arctic bottom photographs give the following indications regarding conditions on the Arctic sea floor: Bottom life appears to be less abundant than in the Atlantic Ocean at similar depths. Loose rocks are more numerous. Ripples seem to be less abundant in the Arctic area photographed than they do at comparable depths in the Atlantic. A bottom current of 0.3 to 0.6 cm./sec. is indicated from movements of sediment clouds. The area is abundantly supplied with curious short "tracks" of uncertain origin. Small irregular holes in the pictures probably represent holes produced by ice-rafted stones and pebbles.--Auth.

2-1369. Hare, F. Kenneth. A PHOTO-RECONNAISSANCE SURVEY OF LABRADOR-UNGAVA: Canada, Dept. Mines & Tech. Surveys, Geog. Branch, Mem. 6, 83 p., 14 pls. 11 figs. incl. 2 maps, 5 tables, 1959, 38 refs.

A terrain analysis of Labrador-Ungava based on aerial photography. In developing methods of interpretation many thousands of square miles of Labrador-Ungava were covered on the ground and many thousands of air miles flown across it. The following physiographic divisions are recognized:

A. Bedrock-controlled plateau elements: 1) western plateau belt, 2) eastern plateau belt, 3) southeastern plateau belt.

B. Upland areas - massifs: 1) center of Kaniapiskau massif, 2) NE. Labrador massif, 3) Mealy Mountain massif, 4) Red Wine-Romaine-Fleur-de-Mai axis, 5) Laurentide massif.

C. Upland areas - Appalachian type: 1) Cape Smith ranges, 2) Labrador trough, 3) Naskaupi ranges, 4) Otish ranges, 5) Belcher-Nastopoka belt.

D. Drift-covered belts: 1) the Lake plateau, 2) southeastern drift belts - Moraine-Mecatina-Hamilton-Naskaupi, Double Mer belts, 3) Whale River depression and associated drift belts W. of the Labrador trough, 4) the Eastmain belt.--F. A. Cook.



### 3. STRUCTURAL GEOLOGY

See also: Areal and Regional Geology 2-1350; Stratigraphy 2-1404, 2-1415; Geophysics 2-1483, 2-1506, 2-1507; Mineral Deposits 2-1587.

2-1370. Kravchenko, G.G. AN EXAMPLE OF PLASTIC DEFORMATION OF LIMESTONES IN TECTONIC FRACTURE ZONES: *Akad. Nauk SSSR, Bull., Geol. Ser.*, in translation, 1958, no. 3, p. 75-85, 8 figs. incl. 4 illus., map, pub. 1959, 10 refs.

In the area of the Kansk polymetallic deposit (southwestern part of Kirghiz SSSR), widely distributed lower Carboniferous limestone (Visean stage) contains numerous fragments of serpentinite. The limestone containing these fragments, that is, the limestone-serpentinite dikes, occurs as linear zones in serpentine rocks. Mapping and study of their structural details led to the conclusion that these formations originated during the process of plastic flow of the limestone in the tectonic fracture zones. --Auth.

2-1371. Griggs, David T., and John Handin, eds. ROCK DEFORMATION (A SYMPOSIUM): *Geol. Soc. America, Mem. 79*, 382 p., illus., diags., graphs, tables, 1960, approx. 250 refs.

This volume contains most of the papers presented at the Symposium on Rock Deformation held in conjunction with the Annual Technical Conference of the Institute of Geophysics, University of California, Los Angeles, Nov. 1956. In most instances these papers have been expanded, extended, and revised as a result of new ideas and further experimental work. This collection of papers gives some recent results of most of the active workers in the United States in experimental rock deformation and also presents several new theoretical contributions. Each paper reports research which was separately conceived and executed and so is necessarily specialized and detailed. Throughout there is the common goal of a better understanding of the mechanics of the deformation of rocks in the earth. --From pref.

Each of the 13 papers included in the symposium is abstracted separately below in the order in which it appears in the book.

2-1372. MacDonald, Gordon J.F. ORIENTATION OF ANISOTROPIC MINERALS IN A STRESS FIELD (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: *Geol. Soc. America, Mem. 79*, p. 1-8, 1960)

The equilibrium orientation of an anisotropic mineral in a stress field is one in which the potential energy of the external forces plus the potential energy of deformation is minimized. For materials obeying a linear elastic stress-strain law, the minimum total potential energy is equivalent to a maximum Helmholtz free energy in an isothermal system or a maximum internal energy in an adiabatic system. The minimum potential-energy theorem is applicable to the orientation of minerals in rocks undergoing plastic deformation provided temperature and elastic strain are sufficient parameters to describe the thermodynamic state of the mineral undergoing isothermal plastic flow. --Auth.

2-1373. Brace, William F. ORIENTATION OF ANISOTROPIC MINERALS IN A STRESS FIELD: DISCUSSION (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: *Geol. Soc. America, Mem. 79*, p. 9-20, 5 diags., 3 tables, 1960)

Based on MacDonald's thermodynamic prediction that minerals should so orient that strain energy is maximized, the preferred orientations of calcite, high and low quartz, and ice are calculated for a stress field consisting of a single component. The following planes should be approximately normal to the stress: ice and calcite, the planes {1011}; high quartz, the planes {1012}; and low quartz, the planes {0221}. A number of grains at equilibrium will have c-axes which occupy girdles about the single-stress component.

The equations for strain energy of hexagonal minerals in a stress field of 3 different principal stresses are derived and applied to the orientation of ice ( $-2^{\circ}\text{C}.$ ). The preferred orientation depends upon ratios of principal stresses and upon stress difference and confining pressure independently. The c axes of grains tend to lie on small-circle girdles about the unique principal stress in a uniaxial stress field. --Auth.

2-1374. Griggs, David T., and M.S. Paterson, Hugh C. Heard, and Francis J. Turner. ANNEALING RECRYSTALLIZATION IN CALCITE CRYSTALS AND AGGREGATES (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: *Geol. Soc. America, Mem. 79*, p. 21-37, 12 illus., 6 diags., table, 1960)

Annealing recrystallization has been produced in cold-strained single crystals of calcite, fine-grained undeformed aggregates, highly sheared aggregates, and in Yule marble deformed cold under controlled conditions.

The critical temperature for recrystallization of Yule marble is about  $500^{\circ}\text{C}.$  In our limited experience annealing time does not seem to influence recrystallization.

In contrast to recrystallization during shear strain (syntectonic recrystallization), which develops high preferred orientation, annealing recrystallization tends to produce a random orientation, but traces of a preferred orientation inherited from the original material remain. Textures of calcite fabrics completely recrystallized by annealing closely resemble granoblastic or porphyroblastic textures of metamorphic rocks.

Recrystallization of calcite is similar in every respect that we have yet explored to recrystallization of metals. Water and  $\text{CO}_2$  have not affected the process in our experiments. --Auth.

2-1375. Griggs, David T., Francis J. Turner, and Hugh C. Heard. DEFORMATION OF ROCKS AT  $500^{\circ}$  TO  $800^{\circ}\text{C}.$  (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: *Geol. Soc. America, Mem. 79*, p. 39-104, 22 illus., (1 col.), 48 diags., 6 tables, 1960)

Extensive experiments on the deformation of rocks at temperatures from  $25^{\circ}$  to  $800^{\circ}\text{C}.$  have been performed in triaxial-test apparatus at 5 kb. confining pressure and in shearing apparatus at pressures to 20 kb. A variety of rocks, crystals, and mixtures have been tested, including peridotite, pyroxenite, basalt, granite, dolomite, marble, and quartz crystals and aggregates.

Strength is reported as a function of temperature. High quartz is found to be very strong. Peridotite, pyroxenite, and granite have nearly the same strength over the entire temperature range - 20 kb. at  $25^{\circ}\text{C}.$  and 7 kb. at  $800^{\circ}\text{C}.$  Basalt has a similar strength to  $600^{\circ}\text{C}.$ , but its strength decreases rap-



dly at higher temperatures. Dolomite is weaker than these rocks at low temperature but at 500°C. has nearly the same strength as peridotite, pyroxenite, and granite. Marble is much weaker at low temperature and decreases in strength more rapidly than the other rocks as temperatures increase.

The principal mechanism of deformation of diopside was found to be {100} translation. Grains of enstatite highly deformed at 500°C., 5 kb. locally invert to clinoenstatite. The mechanism of deformation of olivine was not identified.

Stress-strain curves of calcite single crystals change character at 500°-600°C. At 600°C., translation becomes important, and at 800°C. basal glide may occur. At 800°C., twinning is little easier than translation, which is still the easiest translation. This large change in the relative ease of translation and twinning at high vs. low temperature has important consequences in the deformation of marble at 800°C.

Syntectonic recrystallization reaches a maximum at 600°C. in marble deformed at 3% per minute. At a lower strain rate the maximum syntectonic recrystallization seems to occur at a lower temperature. The recrystallized grains show a pronounced orientation with the c-axes parallel to the maximum principal compressive stress. Water and CO<sub>2</sub> do not affect this recrystallization in our experience to date. The recrystallization observed in these experiments is thought to be of the same type as that which occurs in metamorphic marmorization.

Only very limited plastic deformation of quartz has been observed at 500° and at 800°C., 5 kb. The mechanism could not be determined. High preferred orientations have been developed in quartz in shear-experiments, particularly when the quartz recrystallized from an amorphous phase subject to large shear stress and strain.--Auth.

1376. Maxwell, John C. EXPERIMENTS ON COMPACTION AND CEMENTATION OF SAND (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: Geol. Soc. America, Mem. 9, p. 105-132, 12 illus., 11 diags., 7 tables, 1960)

Compaction and cementation of quartz sands and natural sandstones have been studied experimentally under conditions simulating deep burial. Approximately 230 experiments, lasting a few hours to 100 days, were successfully completed.

Compaction, resulting in loss of porosity, comes about largely by mechanical fracturing, rotation, and interpenetration of grains. High temperatures and the presence of distilled water and alkaline saline waters facilitate compaction. Flowing solutions are more effective in this respect than are static solutions.

Cementation occurs more or less independently of compaction. Again, flowing solutions are the most efficacious. Very little cementation resulted from any of the experiments with distilled water. With sea water and saline-formation waters the degree of cementation is a function of temperature and develops with increasing rapidity above 270°C. It occurs at all, the cementation is apparent within the first 24 hours.

The fabric of many natural sandstones suggests that compaction has been brought about primarily by solution of quartz grains at points of contact. We were unable to demonstrate this process experimentally, however, although it may have occurred to some degree in experiments at high temperatures

with flowing solutions. One experiment involving an alkaline solution of colloidal silica duplicates the fabric of a natural sandstone quite closely, in that the fractures resulting from compaction apparently have been healed by deposition of secondary quartz. It is suggested that mechanical straining, fracturing, and crushing of quartz grains at points of contact, followed by solution of the strained and fractured quartz and healing of fractured grains, may be important in consolidation of natural sandstones.

An anomalous increase in apparent compressive strength of quartz with rising temperature was observed in the interval between 235° and 270°C. This effect is possibly related to incipient plastic deformation of the quartz grains and may be a factor in the preservation of porosity of deeply buried sandstones.--Auth.

2-1377. Borg, Iris, Melvin Friedman, John Handin, and Donald V. Higgs. EXPERIMENTAL DEFORMATION OF ST. PETER SAND: A STUDY OF CATACLASTIC FLOW (In: Griggs, David T., and John Handin, eds. *Rock Deformation (A Symposium)*: Geol. Soc. America, Mem. 79, p. 133-191, 10 illus., 58 diags., 4 graphs, 14 tables, 1960)

In order to make a quantitative study of a purely cataclastic deformation, jacketed cylindrical specimens of dry St. Peter quartz sand of different grain sizes have been deformed under uniform confining pressure and differential load. Results presented illustrate the influence of the deformation on porosity, pore size, grain size, grain fracturing and orientation, and strength. The present experiments, conducted under deliberately restricted conditions, are intended as a first step toward an adequate understanding of the complicated problem of sandstone deformation.

Undeformed St. Peter sand grains are unfractured to slightly fractured; contain planes of liquid and gas inclusions which exhibit a decided preference for the prism zone; exhibit a predominance of grains with no undulatory extinction as well as grains which are slightly, moderately, and highly undulatory; contain very few grains showing "deformation" lamellae; have an intercept ratio of 1:1.3:1.7, and a high intercept sphericity (0.77 to 0.79); and appear to have grain shapes controlled by  $r\{10\bar{1}1\}$ ,  $z\{01\bar{1}1\}$ , and the prisms. The undeformed sand aggregates exhibit an apparent elongation so that the long axes of the grains tend to lie parallel to the circular section. In addition, the aggregates show a tendency for the optic axes of the grains to form a girdle parallel to the circular section.

Fracturing is the most conspicuous feature of the deformed sands. The fracture pattern of specimens deformed under uniform pressure is random. Fracture-orientation patterns reflect the symmetry of the deformation under differential loading conditions. In compressed specimens the fractures tend to lie at small angles to the direction of loading and are probably shear fractures. In extended specimens the late-formed fractures tend to lie normal to the axis of extension and are most probably tensile fractures.

Apparent elongation and optic-axis orientations for compressed samples are much like those of undeformed sands. In extension, however, the apparent elongations of the grains undergo a progressive rotation through 90° from their initial positions. This accompanied by a similar shift in optic-axis orientations and therefore indicates that bodily rotation of grains does occur.

There is no evidence that the experimental de-

formation has produced any "deformation" lamellae or that the over-all undulatory extinction index has been changed. The percentage of moderately and highly undulatory grains decreases, however, indicating that these grains are particularly susceptible to fracturing.

For sands of uniform grain size the coarsest sand (250-300 microns) exhibits the greatest compressibility (uniform confining pressure), which correlates with the facts that this sand also exhibits the greatest reduction in porosity, median pore size, median grain size, and the lowest percentage of unbroken grains. The coarsest sand is the strongest in triaxial tests (differential loading). On the other hand, the finest sand (105-125 microns) is the least compressible; it exhibits the least reduction in porosity, median pore size, and median grain size, and the largest percentage of unbroken grains, and it is the weakest in triaxial tests.

It is significant that uniform loading (approximately simulating an overburden pressure alone) does not reorient the fabric of the sands, whereas differential loading results in preferred orientations of fractures, optic axes, and apparent grain elongations, all of which reflect the orientations of the principal stresses across the boundaries of the specimens. This suggests that petrographic studies of natural sands may be used to distinguish between the effects of overburden and of tectonics.--Auth.

2-1378. Heard, Hugh C. TRANSITION FROM BRITTLE FRACTURE TO DUCTILE FLOW IN SOLENHOFEN LIMESTONE AS A FUNCTION OF TEMPERATURE, CONFINING PRESSURE, AND INTERSTITIAL FLUID PRESSURE (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 193-226, 3 illus., 22 diags., 5 tables, 1960)

More than 115 triaxial compression and extension experiments have been performed on the mechanically isotropic, homogeneous Solenhofen limestone to determine the transition from brittle fracture to ductile flow as a function of temperature, confining pressure, and interstitial fluid pressure. Temperature and confining pressures ranged from 25° to 700°C. and from 1 to 7,500 atm.; interstitial fluids included water and carbon dioxide. Strain rates remained constant at 10<sup>-4</sup>/second. Transitional behavior was arbitrarily defined as that point at which 3-5% strain may be induced without notable loss in cohesion. In dry extension tests, the confining pressure required to induce transitional behavior decreased from 7,300 atm. at 25°C. to about 700 atm. at 700°C. In dry compression tests, the pressure required for the transition was 1,000 atm. at 25°C., decreasing to 1 atm. at 480°C. For transitional behavior in compression tests with interstitial fluids, with increasing confining pressure, the difference between confining pressure and interstitial pressure decreased almost exponentially from 1,000 atm. at 25°C. and 850 atm. at 150°C. to nearly zero at 5,000 atm.

Data for all stress-strain curves are summarized; most stress-strain curves are plotted and compared for a wide variety of test conditions. From these, plots of transition stress or ultimate stress at transition conditions, maximum and minimum principal stresses, and mean stress are derived and reported as a function of temperature.

None of the various theories of strength examined were able to correlate results in compression and extension. The Mohr criterion predicted shear

fracture angles within 4° at the brittle-ductile transition for dry compression tests and within 7° for extension tests at temperatures ranging from 25° to 400°C. By applying the Mohr theory, similar angles were predicted within 3° for transitional compression tests with interstitial water at both 25° and 150°C.

The geological implications of these experiments are briefly discussed. Even though the experimental strain rate is vastly greater than tectonic strain rates, the qualitative difference observed between compression and extension would probably apply to naturally deformed limestones. The analogy of compression experiments to reverse faulting and of extension experiments to normal faulting is discussed. In the event that strain rate does not affect the brittle-ductile transition, these experiments predict normal faulting of dry limestone to a depth of 15 km. and reverse faulting to a depth of 3.5 km. Interstitial fluid pressure would increase the depth to which faulting could occur.--Auth.

2-1379. Robertson, Eugene C. CREEP OF SOLENHOFEN LIMESTONE UNDER MODERATE HYDROSTATIC PRESSURE (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, 227-244, 9 diags., 2 tables, 1960)

Stress, strain, and time observations were made of 10 solid cylinders of Solenhofen limestone compressed into the plastic range or beyond rupture; 32 observations were made of transient creep under nearly constant hydrostatic pressure and stress difference. Eighteen creep tests were made on Danby marble, Rutland White marble, and calcite, and 10 creep tests on the punching of Solenhofen limestone discs.

One to six consecutive creep tests on each specimen were made under hydrostatic pressures up to 4,000 bars under stress differences from 3,000 to 8,000 bars at room temperature. Although the data are somewhat inexact and refer only to transient creep (measured for periods less than 3 hours), certain generalizations can be made: 1) The effect of hydrostatic pressure in decreasing the creep rate is demonstrated by the 100-fold decrease in the creep rate per unit stress difference as hydrostatic pressure is increased from 1,000 bars to 2,000 bars. 2) The similarity of creep behavior of obviously highly fractured specimens to those apparently deformed wholly by plastic flow suggests that fracturing is one of the principal mechanisms of creep in limestone. Healing of fractures is indicated by recovery of deformation upon unloading. 3) Two empirical laws were found to satisfy the creep data: a logarithmic relation for strain-time data,  $\epsilon = kt^{-1}$ , and a linear power law for the strain rate-stress difference data,  $\dot{\epsilon} = k_1 \sigma_D^{-k_2}$ .

A dimensionless strain ration,  $t\dot{\epsilon}/\epsilon$  and a dimensionless stress ratio,  $(\sigma_z - \sigma_m)/\sigma_m$ , were found to be parameters respectively of fracturing and of ductility for limestone.--Auth.

2-1380. Handin, John, Donald V. Higgs, and Joseph K. O'Brien. TORSION OF YULE MARBLE UNDER CONFINING PRESSURE (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 245-274, 14 illus., 11 diags., 7 tables, 1960)

Jacketed cylindrical specimens of Yule marble and calcite single crystals have been twisted at a constant rate of 0.1 radian per minute under confin-



pressures of 1,000 to 2,750 kg./cm.<sup>2</sup>, at temperatures from 24°C. to 300°C., and under axial loads 0 to 1,000 kg./cm.<sup>2</sup> Permanent twists of as much as 1.6 rad/cm. have been obtained. The yield stress and ductility of the marble increase with confining pressure and are further enhanced by an axial compressive load. The yield stress decreases, and the ductility increases as the temperature is raised. Friction under confining pressure alone leads to tensile fracture across a helical surface which is everywhere normal to the least principal stress, a tension of about 200 kg./cm.<sup>2</sup> In specimens under axial compressive load, failure is by shear fracturing approximately parallel to the plane of maximum shearing stress.

Petrographic studies reveal that macroscopically the shear occurs parallel to circular sections. The sliding elements in the grains of the aggregate and the single crystals have not yet been delineated, but the evidence suggests translation gliding on certain planes not previously observed in marble formed under uniaxial compression or extension. Mechanical twinning on {0112} is common but cannot be the predominant flow mechanism.--Auth.

1381. Higgs, Donald V., Melvin Friedman, and John E. Gebhart. PETROFABRIC ANALYSIS BY MEANS OF THE X-RAY DIFFRACTOMETER (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 275-292, illus., chart, 28 diag., plate, 1960)

The sample holder designed for adaptation of a Perkin-Elmer X-ray diffractometer to petrofabric analysis permits relatively rapid definition of preferred crystallographic orientations. A complete diagram is obtained in about 3 hours by examination of one sample. Moreover, diagrams are obtained by plotting directly the intensities recorded on a strip chart. Absorption and geometric corrections are unnecessary because, by studying hemispherical surfaces, absorption is held constant, and, by construction of the sample holder, misalignment problems are eliminated.

Three rocks - Poughquag quartzite, Yule marble, and Solenhofen limestone - were studied. Orientations in the quartzite and marble determined by X-rays are in excellent agreement with fabrics obtained from microscopic measurements. The quartzite exhibits a preferred  $a$  axis as well as a  $c_v$  orientation. X-ray diagrams record changes in crystallographic orientations for experimentally deformed Yule marble. The limestone, which is too fine-grained to be studied optically, exhibits a slight tendency for preferential  $c_v$  orientation.--Auth.

1382. Odé, Helmer. FAULTING AS A VELOCITY DISCONTINUITY IN PLASTIC DEFORMATION (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 293-321, 25 diag., 1960)

Triaxial loading of certain rocks sometimes results in faulting without fracture in the ordinary sense, or sudden displacements. A brief discussion of the theory of brittle fracture of Griffith is given to indicate why it cannot apply to this kind of ductile faulting. Such ductile faulting can be explained by the theory of plasticity for plane strain. This theory is presented for more general yield conditions than that of von Mises, and it is assumed that the stress-strain rate relations can be derived from the yield

condition by a process of differentiation. This formalism is by no means original, but it is little known among geologists. Across certain planes in the plastic mass, discontinuities in the velocity are possible, whereas the stresses remain continuous across these planes. These "characteristic planes" of the velocity equations are identified with the planes of ductile faulting. The results obtained in the theory of plane strain for the von Mises, Coulomb, and Torre yield conditions do not hold for the more general 3-dimensional theory, unless very stringent conditions on the strain rates are satisfied. Smooth characteristic surfaces in the plastic domain are possible only if these conditions are satisfied. Possibly, however, there are yield conditions which lead to smooth characteristic surfaces under weaker restrictions on the strain rates.--Auth.

2-1383. Orowan, E. MECHANISM OF SEISMIC FAULTING (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 323-345, 10 diag., 1960)

According to the classical theory, earthquakes would be caused by fracture, or the release of static friction, followed by sliding between the fault walls and a consequent drop of stress around the fault. At a depth of 600-700 km., where the pressure is about 200,000 bars, the lowest possible estimate of the coefficient of friction (about 1) would demand principal stress differences of the order of 400,000 bars for overcoming the friction of the fault wall and producing a drop in seismic stress. Such high stresses cannot be assumed to exist, if only because the yield stress of the asthenospheric rocks can hardly have an order of magnitude exceeding a few tens or at most hundreds of bars. This estimate is obtained by extrapolation of the observed creep behavior of hard materials with the assumption that the effect of a hydrostatic pressure is similar to that of a correspondingly increased molecular cohesive pressure; its order of magnitude agrees with that of the stress drop estimated from the energies released in earthquakes.

A detailed discussion of this difficulty shows that earthquakes, except at focal depths less than about 5-10 km., cannot arise in the manner implied in the classical theory. The only plausible alternative available at present is that they are due to an instability of plastic deformation (creep) such as gives rise to slip bands. Lüders bands, the Hanson-Wheeler creep-deformation bands, and many other similar phenomena. If creep produces structural changes that accelerate further creep, the deformation concentrates gradually into thin layers in which high flow rates can develop, and finally even shear melting may occur by the heat development due to plastic deformation.

Such a mechanism would explain the sequence structure of earthquakes (fore- and aftershocks). If fracture and frictional sliding are impossible and faulting can occur only by the gradual concentration of creep deformation into thin zones, the stress concentration around the edges of a fault cannot propagate this fault immediately over the entire stressed volume, as would be the case with faulting due to fracture. The fault can extend only after the occurrence of a certain amount of creep leading to progressive strain concentration in the region put under increased stress by the preceding faulting. The first faults in a sequence are likely to be small; after several preparatory faults have thrown a high stress upon a larger volume, however, extensive faulting with larger shocks may develop. Toward



the end of the sequence the intensity of the shocks is likely to decrease again by depletion of stress.

Since creep instability seems at present the only feasible mechanism of deep and intermediate-depth earthquakes, the fact that it is a characteristic of crystalline materials indicates that the seismogenic parts of the Earth's mantle are substantially crystalline.--Auth.

2-1384. Griggs, David T., and John Handin. OBSERVATIONS ON FRACTURE AND A HYPOTHESIS OF EARTHQUAKES (In: Griggs, David T., and John Handin, eds. Rock Deformation (A Symposium): Geol. Soc. America, Mem. 79, p. 347-364, 15 illus., chart, diag., 1960)

Experiments have shown that the deformation of rocks under high confining pressures cannot be described by terms based on ordinary experience. The writers believe there are only 3 fundamental macroscopic processes at work: extension fracturing, faulting, and uniform flow. The first involves separation across a plane of no shear normal to the direction of least principal stress and includes tensile fracture as a special case. Faulting involves a shearing displacement and occurs with or without loss of cohesion. It includes shear fracturing as a special case and may occur along a plane inclined at from a few degrees to  $45^{\circ}$  to the direction of maximum principal stress. From the standpoint of the stress-strain curve, faulting without loss of cohesion is indistinguishable from flow. (The 3 flow mechanisms - cataclasis, gliding, and recrystallization - cannot always be distinguished either.) Examples of experimental boudins are presented to illustrate these categories of deformation.

The field evidence is that earthquakes are accompanied by shearing displacements and are therefore due to faulting in the general sense. Since stored elastic strain energy is released, there must be at least a momentary and local loss of cohesion. A crack then propagates at close to the speed of sound. For deep-focus earthquakes (down to 700 km.) certainly, and most probably even for shallow disturbances (a few tens of kilometers), ordinary Coulomb fracture is impossible. The internal friction of dry rocks under tens or hundreds of thousands of bars pressure would demand impossibly high shearing stresses of many kilobars. The most reasonable mechanism of energy release at great depth is a phase change, and the most probable phase change is melting.

Calculations suggest that, once a crack or a flaw exists, there is ample elastic energy to propagate the crack by shear melting even if the stress difference is only a few tens of bars. It is, of course, inconceivable that an open crack could exist at depth, so that the most baffling problem is the nature of a flaw of the Griffith type under these conditions. Although it is little more than speculation, we suggest that the flaw may be a pocket of already molten rock or of its more volatile constituents. This idea receives some support from the intimate association of earthquake epicenters and zones of volcanic activity.--Auth.

2-1385. Ball, T.K. A PETROFABRIC ANALYSIS OF A FOLD: *Am. Jour. Sci.*, v. 258, no. 4, p. 274-281, 7 figs., Apr. 1960, 9 refs.

A petrofabric method whereby polygenetic folds can be analyzed is described. The method has been applied to the study of the tectonic history of a fold

from the Grampian Highlands of Scotland. An earlier deformation was produced by shearing, followed by accordion folding, utilizing the same S-planes.--Auth.

2-1386. Clark, Lorin D. FOOTHILLS FAULT SYSTEM, WESTERN SIERRA NEVADA, CALIFORNIA: *Geol. Soc. America, Bull.*, v. 71, no. 4, p. 483-496, 4 illus., 2 maps (1 fold.), 2 fold. secs., Apr. 1960, 37 refs.

A large fault system, here named the Foothills fault system, is the dominant structural feature of the western Sierra Nevada. The steeply dipping to vertical component faults trend northwestward through an area about 200 mi. long and 30 mi. wide N. of  $37^{\circ}30'N$ . The faulted Paleozoic and Mesozoic rocks are overlapped by unfaulted younger rocks, and the total extent of the fault system is not known. It is probably not limited to the western Sierra Nevada. Faults are marked by belts as much as 4 mi. wide, of cataclastically deformed and recrystallized rocks and by truncated folds. Along 1 fault, Upper Jurassic rocks are juxtaposed against Paleozoic rocks for at least 100 mi. The direction of fault movement has not been determined. Net displacement on some of the component faults exceeds 3,000 ft. and may be measurable in miles. Major faults cut beds of Late Jurassic age and are in turn cut by plutonic rocks of probable Late Jurassic and Middle Cretaceous age. Faults that controlled deposition of quartz veins and Au ore bodies of the Mother Lode belt are apparently younger and structurally less important features superimposed on one of the fault zones of the large system.--Auth.

2-1387. Allen, Clarence R., Leon T. Silver, and Francis G. Stehli. AGUA BLANCA FAULT - A MAJOR TRANSVERSE STRUCTURE OF NORTHERN BAJA CALIFORNIA, MEXICO: *Geol. Soc. America, Bull.*, v. 71, no. 4, p. 457-482, 7 illus., 10 maps, table, Apr. 1960, 44 refs.

Agua Blanca fault is a major right-handed strike-slip fault at least 80 mi. in length that cuts transversely across the peninsula of Baja California about 70 mi. S. of the international border. Its trend is anomalous in being more nearly parallel to the Transverse Ranges of southern California than to the San Andreas fault system that elsewhere dominates the tectonic grain of the peninsula. Geographic features delineating the fault trace are, from E. to W: Paso San Matías, Valle de La Trinidad, Cañon de Dolores, Valle de Agua Blanca, Valle de Santo Tomás, Bahía Soledad (S. branch of fault), Punta Banda (N. branch). Farther W., both branches of the fault control submarine topography, and possibly the fault system is continuous with the NW-trending San Clemente fault off the southern California coast.

Physiographic expression of the Agua Blanca fault is remarkably similar to that of the San Andreas. Typical features are Recent scarps, offset streams, shutterridges, fault sags and saddles, side-hill ridges, and fault-controlled valleys. Most of these features are particularly well exhibited in Valle de Agua Blanca, which is designated as the type locality.

Rocks cut by the fault are mainly Cretaceous plutonic rocks of the southern California batholith and Lower Cretaceous (Albian) metavolcanic rocks. Along the Pacific Coast, the fault cuts Upper Cretaceous (Maestrichtian) postbatholithic sedimentary rocks that are otherwise surprisingly little deformed as compared to rocks of similar age in most of



California.

Agua Blanca fault shows a history of right-lateral displacement throughout its length. Recent stream offsets occur from Valle de La Trinidad nearly to the Pacific Ocean; distinctive Quaternary (?) fan travels in Valle de Agua Blanca are offset laterally 1 mi. from their most likely source area across the fault; in the same area, a fault slice of distinctive antiperthitic granodiorite is best explained by 7 mi. of lateral displacement, and a nearby slice of quartz diorite may indicate displacement as great as 14 mi. In general, evidence of both Recent activity and amount of total displacement appear to increase westward. The fault coincides in gross aspect with a broad E.-W. zone of seismic activity in a region elsewhere characterized by relative quiescence, but no large historical earthquakes can be positively correlated with this fault.

Despite its orientation athwart the regional tectonic grain, Agua Blanca fault does not appear to represent a deep-seated structural feature analogous to those of the Transverse Ranges. Instead, it is probably one of several paths by which the San Andreas fault tends to break around the "knot" caused by the great bend of the San Andreas in southern California. -- Auth.

-1388. King, Philip B. THE ANATOMY AND HABITAT OF LOW-ANGLE THRUST FAULTS: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 115-125, secs., 1960, 39 refs.

Low-angle thrust faults, commonly called "overthrusts," are faults with reverse movement, which dip at low angles. Along them, great sheets of rock have moved over the rocks beneath for distances as much as many miles. How they were formed is poorly understood, but much has been learned in recent decades regarding details of their geometry and space relations. Two relatively simple examples, the Pine Mountain fault of the Appalachian region and the Johnnie-Wheeler Pass fault of the Cordilleran region, illustrate this geometry. Initial fractures of these faults, rather than having been regularly dipping shear planes, consisted of a succession of flats and pitches, formed, respectively, on incompetent and competent units of the stratified sequence. This and other characteristics of these examples may also apply to more complex low-angle thrust faults, which have had a more eventful history and have been more obscured by superposed structures. Low-angle thrust faults characterize mainly the tectogeosynclinal realm, and in many places seem to have formed near the margins of sedimentary basins or along the boundaries between one sedimentary basin and another. -- Auth.

-1389. Gilluly, James. A FOLDED THRUST IN NEVADA - INFERENCES AS TO TIME RELATIONS BETWEEN FOLDING AND FAULTING: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 68-79, 9 maps, secs., 2 diags., 1960, 15 refs.

A segment of the Roberts thrust, a fault of regional extent in central Nevada, is strongly overturned at the border of a window in the northern Shoshone Range. The several branch faults that spring from the main thrust at and near the axis of overturning are, in order of decreasing age, successively less warped. This sequence suggests that folding of the thrust was concurrent with its development. The map pattern resembles that of many other windows in large thrusts; it is suggested that "eyelid" windows

(which are framed by 2 or more thrust surfaces) are generally formed by concurrent folding and thrusting. -- Auth.

2-1390. Billings, Marland P. DIASTROPHISM AND MOUNTAIN BUILDING: *Geol. Soc. America, Bull.*, v. 71, no. 4, p. 363-397, 34 figs., incl. maps, secs., diags., 3 tables, Apr. 1960, 154 refs.

For many years the terms orogeny and mountain building have been used so loosely that they no longer convey definite meanings. Because of this lack of precision, hypotheses of diastrophism are often based on false premises.

Folding and thrusting, often referred to as tectogenesis or orogeny, is best displayed by stratified rocks. Some geologists believe that folds and thrusts are due to vertical movements accompanied by lateral spreading, that the strata are stretched an amount commensurate with the intensity of the folding, and that the opposite sides of the sedimentary packet do not move toward each other. Most geologists, however, agree that during folding and thrusting the strata are shortened and that the 2 opposite sides of the sedimentary packet move toward each other by an amount commensurate with the intensity of the folding. But it is not clear to what extent the entire crust in the deformed belt is shortened. In one extreme view, the 2 opposite ends of the deformed belt do not move toward each other, and the folds and thrusts result from gravity sliding. At the other extreme, the entire crust is believed to be shortened to the same extent as the sedimentary strata.

An excellent example of broad vertical movements (epeirogenesis) that has never been sufficiently emphasized is in eastern North America, involving the Appalachians, Coastal Plain, and Continental Shelf. During the Mesozoic and Cenozoic the Fall-Line surface was depressed as much as 20,000 ft. beneath the Continental Shelf, whereas it was uplifted at least 8,000 ft. over the Appalachians. The north-eastern two-thirds of the Appalachian-Ouachita belt has been going up since the Jurassic, whereas the southwestern third has been going down. Thus the present Appalachians are unrelated to the late Paleozoic folding. Certain metamorphic minerals - such as kyanite - or certain metamorphic assemblages - such as jadeite and quartz - form at confining pressures found only at depths of tens of miles, implying tremendous erosion wherever such minerals or assemblages are exposed.

Broad vertical movements, accompanied by high-angle faulting, are illustrated by the fault-block mountains of the Basin and Range province and by such features as the Rhine graben and the rift valleys of Africa.

Although the nature of the displacement along the San Andreas fault [California] has been known for more than 50 years, only in recent decades have other large strike-slip faults been recognized. But it is now the fad to assign all kinds of faults and even nonexistent faults to the strike-slip category. In recent years seismologists have emphasized the importance of strike-slip faults. The author suggests that in the Fairview Peak-Dixie Valley, Nevada, earthquakes there is evidence that strike-slip movements are invading an area previously characterized by Basin-Range structure. The geologic record indicates that large strike-slip faults have been distinctly subordinate to folding, broad vertical movements, and broad vertical movements accompanied by high-angle faulting.

Possible displacement of the crust or the entire

earth relative to the axis of rotation has been emphasized in recent years. Continental drift, if it occurs, is one such type of movement. Slipping of the entire crust over the mantle is another. Much research has been accomplished in paleomagnetism in recent years, but the data are too scanty and conflicting to permit any definite conclusions.

Among the possible causes of diastrophism are 1) contraction of the earth, 2) convection currents, 3) formation of large pockets of magma, 4) sialic material leaking out of the mantle, 5) conversion of sial to mantle by change of low-pressure minerals to high-pressure minerals, or the reverse, and 6) serpentinization or deserpentinization of the upper part of the mantle.

Mountain building is merely one manifestation of vertical movements of the crust. In the past, mountain building has been erroneously considered by many to be primarily the result of folding and thrusting. Certainly many modern ranges are the result of vertical uplift unrelated to folding. Many such uplifts are accompanied by high-angle faulting to produce fault-block mountains. Blocks caught between strike-slip faults may be squeezed upward if the blocks move toward each other. Mountainous uplifts have also resulted from folding and thrusting. Conversely, in some folded belts erosion appears to have kept pace with the rise of the folds so that no mountains developed.--Auth.

2-1391. Knopf, Adolph. ANALYSIS OF SOME RECENT GEOSYNCLINAL THEORY: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 126-136, table, 1960, 40 refs.

"Geosyncline" was invented by Dana in 1873. It was defined as "a down-bending of the crust"; the context supplied the connotation that a mountain chain would eventually arise from sediments accumulated in such a down-bending. However, "geosyncline" has not yet been adequately defined.

Stille's 1940 classification of geosynclines was introduced into America by Kay in 1942. Stille's 2 principal classes are based upon capacity to be folded. Highly foldable geosynclines are orthogeosynclines, weakly foldable, parageosynclines. Orthogeosynclines include eugeosynclines and miogeosynclines, distinguished by position, time of folding, and magmatism. Eugeosynclines supposedly determine the sites of autochthonous palaeogenic batholithic magma. However, batholiths occur also in so-called miogeosynclines, which necessitates the auxiliary hypothesis that magmas can travel laterally from their eugeosynclinal place of origin to a miogeosynclinal environment. Recent absolute age determinations on batholiths render this hypothesis untenable.

Lawson maintained that geosynclines do not collapse until they have accumulated a critical thickness - approximately 40,000 ft. Nevertheless, geosynclines have been shown to fail under loads of 10,000 ft. or less. The maximum thickness of accumulation appears to be 40,000 ft.; consequently this limit suggests that stratigraphic thicknesses estimated to exceed 40,000 ft. require reduction.

In 1944 Stille introduced the concept of "regeneration," meaning thereby restoration of the orthogeosynclinal state in ancient fold-belts. This idea has stimulated, among other things, active interest in the fate of ore deposits in regenerated orthogeosynclines. Critical evidence of the effects of regeneration on the dispersal of preexistent ores is given by the isotopic composition of leads in galenas formed during successive orogenies.--Auth.

2-1392. Longwell, Chester R. POSSIBLE EXPLANATION OF DIVERSE STRUCTURAL PATTERNS IN SOUTHERN NEVADA: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 192-203, 3 maps, 1960, 18 refs.

In southern Nevada a belt of folds and thrusts striking generally N. is transected by a NW.-trending fracture zone on which right-lateral movement has occurred. This movement was in progress while some of the thrusts and folds were active. A suggested explanation assumes that the 2 structural patterns reflect independent controls at different levels in the crust.--Auth.

2-1393. Webb, Philip K. GEOLOGY OF THE CAVANAL SYNCLINE, LE FLORE COUNTY, OKLAHOMA: *Oklahoma Geol. Survey, Circ. 51*, 65 p., 2 maps (col. geol. map in pocket), 1960, 61 refs.

The Cavanal syncline is expressed at the surface by the Cavanal Mountain and the Potato Mountains. Exposed rocks are of the Savanna and Boggy formations, Krebs group, Des Moines series [Pennsylvanian]. The syncline is a broad downwarp cut by N.-S. faults.--Auth.

2-1394. Osmond, John C. TECTONIC HISTORY OF THE BASIN AND RANGE PROVINCE IN UTAH AND NEVADA: *Mining Engineering*, v. 12, no. 3, p. 251-265, 8 maps, 9 secs., profile, March 1960, 50 refs.

N.-trending discontinuous ranges consisting of tilted fault blocks with complicated internal structure are characteristic of the area from the Wasatch Range to the Sierra Nevada. Generalized maps and cross sections illustrate the structural evolution of the region.

Beginning in Cambrian time, the Cordilleran geosyncline occupied the western margin of North America. The geosyncline was divided lengthwise into 2 parallel parts by a tectonic zone which trended NNE. across Nevada from near Tonapah.

This zone separated shallow water carbonates on the E. from deeper water clastics and volcanics on the W. during the lower Paleozoic. The tectonic rose as an elongate welt in Mississippian time and expanded eastward and westward until, by lower Tertiary time, it included the former area of the geosyncline. Zones of subsidence flanking the welt migrated to E. and W. as it expanded.

Deformation occurred in various parts of the welt in each period since Devonian with greater intensity in Mississippian, Jurassic-Cretaceous (Nevadan), and Cretaceous-Tertiary (Laramide). Widespread intrusion of granitic stocks occurred in middle Cretaceous and middle Tertiary.

Toward the end of lower Tertiary the relief of the region had been reduced, and most of it was covered by a large volume of essentially horizontal volcanic extrusives. Following the extensive intrusion and extrusion, the crust of the welt adjusted to regional rise by movement along preexisting fractures, and the present ranges and valleys began to develop.

Movement along faults has continued since Miocene, and many of the valleys have been filled with over 5,000 ft. of tuffaceous clays and sands.--Auth.

2-1395. Lomashov, I. P. ON THE RELIEF OF LIMESTONE FOUNDATION IN THE SUB-MOSCOW BASIN: *Akad. Nauk SSSR, Bull., Geol. Ser.*, in



## STRATIGRAPHY AND HISTORICAL GEOLOGY

translation, 1958, no. 3, p. 86-96, 9 maps, pub. 1959, 6 refs.

The foundation of the Moscow basin consists of predominantly carbonate lower Carboniferous and Upper Devonian rocks. The surface of the foundation has a general dip toward the center of the Moscow syncline of 1.5 to 2 m. per km. and is characterized by locally significant irregularities of tectonic origin.--M. Russell.

-1396. Rezanov, I. A. TECTONIC MAP OF THE

TURKMEN-KHORASSAN MOUNTAINS: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 51-67, map, pub. 1959, 41 refs.

This paper describes the map which has been constructed by the separation of structural units which differ from each other in the time of formation of flexures and the time of upheavals which altered these flexures. The nonuniform character of folding movements is shown by the age of folding and morphology of the folded structures.--Auth.

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See also: Geologic Maps 2-1345; Areal and Regional Geology 2-1349; Geomorphology 2-1359; Structural Geology 2-1395; Paleontology 2-1433, 2-1469, 2-1471; Geophysics 2-1503.

-1397. FORMATION CORRELATOR OF THE UNITED STATES: Oil & Gas Jour. v. 58, no. 17, Supplement (fold. chart), Apr. 25, 1960.

A 22 in. x 24 in. table with small insert map of the United States illustrating the major basins. The table gives the formations for the West Coast, Gulf Coast, Rocky Mountains, Midcontinent, and Eastern states areas. The tabulation is by states within the areas. Producing formations are indicated.--N. Street.

-1398. Zhemchuzhnikov, Yu. A. SIMILARITIES AND DIFFERENCES BETWEEN FACIES, FACIES-CYCLIC, AND FACIES-TECTONIC METHODS OF THE STUDY OF COAL MEASURES: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 1-7, pub. 1959, 18 refs.

Periodicity in sedimentation is a characteristic of coal beds. Both the facies-tectonic and the facies-cyclic methods of analysing this feature agree in recognizing facies and rhythms as factors and regard the latter as the result of oscillating motion against the background of general subsidence. The pure facies analysis ignores cyclic regularity as a factor. Whether tectonic activity is a factor in such alternation of beds is not proven one way or the other, but clearly cycles or rhythms in sedimentation have been established.--M. Russell.

-1399. James, Harold L. PROBLEMS OF STRATIGRAPHY AND CORRELATION OF PRECAMBRIAN ROCKS WITH PARTICULAR REFERENCE TO THE LAKE SUPERIOR REGION: Am. Jour. Sci., 258-A (Bradley Volume), p. 104-114, chart, sec., graph, 1960, 23 refs.

Two great advances in geologic knowledge - the gradually accruing evidence of the immense duration of Precambrian time, which is now believed to represent about 8 times that of the Paleozoic, Mesozoic, and Cenozoic combined, and development of the facies concept, with concomitant destruction of "layer cake" stratigraphy - have forced thorough reappraisal of stratigraphic correlations of Precambrian rocks. This reappraisal, which is being made by many students of the Precambrian, has thrown doubt on heretofore accepted correlations and has led to greater caution in regional syntheses. Most of the familiar time and time-stratigraphic terms have been virtually abandoned except in type areas, but despite this apparent retrogression, Precambrian stratigraphy

probably is on a sounder basis than ever before.

Within individual areas, even areas of highly deformed and metamorphosed rocks, problems of stratigraphy and of correlation are being solved by detailed mapping and application of the physical criteria of correlation. In the Lake Superior region, for example, about 100 formations, plus named stratigraphic equivalents, can be placed in sequential position; these formations represent a probable time span of more than 2 billion years. District-to-district correlations are made on the basis of structural position, highly distinctive sequences and "carry-over" units, and bracketing by isotope age determinations on older and younger crystalline rocks. The resultant stratigraphic classifications and correlations cannot attain the refinement possible in fossiliferous strata, but they provide useful approximations for reconstruction of regional geologic history.--Auth.

2-1400. Sudovikov, N. G. NEW DATA ON THE GEOLOGY OF THE PRECAMBRIAN OF THE ALDAN: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 73-79, pub. 1959, 9 refs.

The main problems of Precambrian geology are examined in the light of data obtained from new investigations by the Laboratory of Precambrian Geology [Academy of Sciences, U.S.S.R.]. Some statements of A. A. Kadensky are critically analyzed as well as the scheme of succession of geological events which he proposes for the Aldan Precambrian.--Auth.

2-1401. Kindle, Cecil H., and Harry B. Whittington. SOME STRATIGRAPHIC PROBLEMS OF THE COW HEAD AREA IN WESTERN NEWFOUNDLAND: New York Acad. Sci., Trans., v. 22, no. 1, p. 7-18, illus., 2 maps, 4 tables, Nov. 1959, ref.

As a result of field work undertaken in 1955, the authors published a paper in 1958 on "Stratigraphy of the Cow Head Region, Western Newfoundland" (Geological Abstracts, v. 6, no. 2, p. 66, June 1958). This paper details the results of some of the work done since that time.

A Cambrian and Ordovician sequence of rocks in the Cow Head area contains within a normal sequence of limestone and shale a series of conglomerates having fossiliferous boulders. From a faunal study of the trilobite fauna, the stratigraphic sequence is currently being deciphered. The Cambrian faunas in the boulders range from the Bathyriscus-Elrathina zone of the Middle Cambrian to the top of the Cambrian. A fault at Cormorant Head on the Port-au-Port Peninsula probably is not related in origin to the conglomerates in associated rocks.--M. Russell.

2-1402. Keller, B. M. SILURIAN DEPOSITS ON THE AK-KERME PENINSULA: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 1-8, 3 illus., map, pub. 1959, 3 refs.

English translation of GeoScience Abstracts 1-379.

2-1403. Nosow, Edmund. SOME DEVONIAN-SILURIAN STRATIGRAPHIC RELATIONSHIPS ON THE WEST FLANK OF THE CINCINNATI ARCH (In: McGrain, Preston, and Thomas J. Crawford, eds. Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959: Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 67-81, 4 maps, 3 charts, 3 secs., log, 1959).

Recent [oil] activity in S.-central Kentucky has created a need for a simplified explanation of Silurian-Devonian relationships on the W. flank of the Cincinnati arch. A discussion of the formations involved, and suggested areas for additional shallow potential oil and gas reservoirs are presented.--Auth.

2-1404. Varentsov, I. M. ON THE DISTRIBUTION OF CLASTICS IN THE GOTLANDIAN AND DEVONIAN OF THE TUVA DOWNWARP: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 107-110, pub. 1959, 3 refs.

The Tuva depression formed during the Gotlandian (Silurian). The 3 stages of the period (Llandoveryan, Wenlock and Ludlow) correspond to a transgressive cycle. The Llandoveryan sand-conglomerate, from 200 to 2,000 m. thick, was formed from vulcanogenic siliceous rocks in a basin with an intensively depressed bottom. The sediments of the Wenlock stage were formed in a shallow sea, and adjacent land was of a flat, slightly peneplain character. During the Ludlow stage the Gotlandian sea was in regression, and the sub-aqual continental-lagoonal multicolored sandy-aleuritic sediments were formed. The Lower Devonian was characterized by increased volcanic activity, and the sediments mixing with pyroclastic materials formed graywackes. During the Eifelian stage, the deposited clastic rocks varied in different parts of the depression. In the central part, the clastogene rocks are similar to the Lower Devonian rocks. They are mostly represented by the tuffite quartz-graywacke and fieldschist-graywacke varieties. The volcanic activity at this stage diminished, and the lava covers are found only in the lower parts of this stage. The accumulation of large strata of red sandstone was accomplished in shallow-water lagoons. The terrigenous deposits of the Givetian stage were mainly composed of tuffite-arkosic varieties. Different types of schist formations are also found here. Thus a Gotlandian and Devonian terrigne complex fills the middle-upper Paleozoic synclinal Tuva depression, and, by studying the changes in the composition and structures of the rocks, we can build a picture of the development of the depression.--L.C.

2-1405. Semikhatova, S. V. ON THE HISTORY OF THE RUSSIAN PLATFORM DURING THE TOURNAISIAN AND EARLY VISEAN: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 9-21, 3 tables, pub. 1959, 35 refs.

The facies and stratigraphic divisions of a lower Carboniferous formation which forms a narrow zone

of thick sediments in the Tatar republic, the Kuybyshev area and Udmurtia are described. The author distinguishes a zone containing *Productus sublaevis* Kon. and sandstone formations containing spores which are older than those of the Stalinogorsk horizon. Paleontologic and geologic arguments are given for placing the boundary between the Tournaisian and Visean stages in the Volga-Ural region at the bottom of the zone containing *Productus sublaevis*.--Auth.

2-1406. Koperina, V. V. FACIES AND TYPES OF COAL ACCUMULATION IN THE DONETS COAL MEASURES: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 1, p. 8-17, pub. 1959, 20 refs.

A review of facies content of rocks in the Donets coal measures shows that facies changes through the section may be explained by the change in relationship between the rate of subsidence and the rate of sediment accumulation and not by oscillatory movements.

Two types of coal accumulation occurred in the Donets: 1) along a flat littoral zone, protected by sand bars; 2) in a delta area with a migratory river mouth.--Auth.

2-1407. Jordan, Louise. PERMIAN SALT BEDS IN LAVERNE GAS AREA, HARPER COUNTY, OKLAHOMA: Oklahoma Geology Notes, v. 20, no. 2, p. 23-28, 3 logs, Feb. 1960, 5 refs.

Gamma, neutron, caliper, and resistivity logs of the Cities Service Oil Company No. 1 Dunnaway "B" well reveal the presence of 3 intervals of Permian salt beds interbedded with shale at depths ranging from 475 to 2,490 ft. below the derrick floor. The shallowest section begins 60 ft. below the Blaine formation at 475 ft. and extends to 625 ft. The section below the Cimarron anhydrite, 1,236 to 1,476 ft., is approximately 50% salt. Ten salt strata, ranging in thickness from 5 to 20 ft., are interbedded with shale and anhydrite from 2,040 to 2,265 ft. in the upper part of the Wellington formation; 49 ft. of salt are present in the 90-ft. interval from 2,400 to 2,490 ft. in the lower part of the Wellington.--Auth.

2-1408. Nesterenko, L. P. ON THE METHOD OF CORRELATION OF THE LOWER PERMIAN SECTIONS OF THE DONBAS, ITS NORTHWESTERN MARGINS AND THE EASTERN PART OF THE DNEIPER-DONETS DEPRESSION: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 111-113, pub. 1959, 6 refs.

English translation of GeoScience Abstracts 1-382.

2-1409. Azarov, A. A. ON THE AGE OF COAL-BEARING DEPOSITIONS IN THE TRANS-BAIKAL: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 1, p. 56-57, pub. 1959, 3 refs.

The Transbaikalian coal measures consist of both Jurassic and Cretaceous rocks. Flora-faunal lists are given.--M. Russell.

2-1410. Pettyjohn, Wayne A. THE DAKOTA CONTROVERSY: South Dakota Acad. Sci., Proc., v. 38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 34-38, map, 1959, pub. 1960, 14 refs.



The name Dakota was applied at the type locality near southeastern South Dakota to a sequence of sandy strata of Early and Late(?) Cretaceous age. However, in the Black Hills region the name "Dakota" has been applied to all or parts of an Early Cretaceous threefold sequence (Lakota sandstone-Fuson shale-Fall River sandstone) that is lithologically similar to the sequence at the type locality. The Fall River sandstone is overlain by the Skull Creek shale, above which a second sand-shale-sand sequence of Early Cretaceous age (Newcastle-Mowry-Frontier") is present. In central South Dakota the name "Dakota" has been applied to at least 2 of the lower 3 sandstones.--Auth.

1411. Dane, Carle H. THE BOUNDARY BETWEEN ROCKS OF CARLILE AND NIOBRARA AGE IN SAN JUAN BASIN, NEW MEXICO AND COLORADO: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 46-56, map, 2 secs., 1960, 23 refs.

In the northern part of the San Juan basin, New Mexico and Colorado, an unconformity at the base of beds of Niobrara [Upper Cretaceous] age cuts out 100 to 400 ft. of beds of late Carlile age [Upper Cretaceous] that in the southern part of the basin are represented by beds from the top of the Juana Lopez sandstone member of Rankin at least up to the top of the Gallego sandstone member of the Gallup sandstone. The Juana Lopez sandstone member of the Carlile shale and the Tres Hermanos sandstone member in the lower part of the Mancos shale are shallow-water marine sandstones, the tops of which are early time equivalent over very extensive areas. The Gallego sandstone member of the Gallup sandstone of the Mesaverde group may be similar to them. The stratigraphic relations suggest that in a belt 30 to 75 mi. wide between the northern and southern parts of the basin downflexing toward the S. occurred during the time of deposition.--Auth.

1412. Maslakova, N.I., and D.P. Naydin. ON CENOMANIAN DEPOSITIONS OF THE CRIMEAN MOUNTAINS: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 97-99, secs., pub. 1959.

Clayey and sandy marls of Cenomanian (Cretaceous) age in the Crimean mountains are divided into distinct stages, the lower characterized by mass accumulations of *Rotalipora apenninica*, and the upper containing *R. reicheli* and *R. montsalvensis*.--M. Russell.

1413. Maslakova, N.I., and D.P. Naydin. ON THE SANTONIAN DEPOSITIONS IN SOUTH-WEST CRIMEA: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 1, p. 57-58, pub. 1959, 5 refs.

Santonian (Upper Cretaceous) deposits in the Akhchisaray area of SW. Crimea are 65 m. thick. Micropaleontological data and the occurrence of *Marsupites testudinarius* indicates their age to be Santonian, Horizon 5.--M. Russell.

1414. Finko, V.I., and E.D. Zaklinskaya. ON THE STRATIGRAPHY OF THE FRIABLE FORMATIONS OF THE ZEYSKO-BUREINSKAYA DEPRESSION: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 22-42, 7 figs. incl. maps, charts, secs., 3 tables, pub. 1959, 22 refs.

Previously existing stratigraphic breakdowns of

the sequences of friable formations in the Zeysko-Bureinskaya depression are subjected to a critical examination. On the basis of geologic, lithologic, and paleobotanical data, a new stratigraphic plan for the friable Tsagayanskaya formations of the Zeysko-Bureinskaya depression is proposed. They are to be classified as the Tsagayanskaya series. The series, in turn, is divided into suites: the Vodorazdel'naya (Oligocene), the Kivdinskaya (Paleocene-Eocene), and the Tsagayanskaya (Upper Cretaceous).--Auth.

2-1415. Love, J. David. CENOZOIC SEDIMENTATION AND CRUSTAL MOVEMENT IN WYOMING: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 204-214, illus., map, table, 1960, 20 refs.

One of the thickest and most complete sequences of Cenozoic nonmarine sedimentary rocks in North America is in Wyoming. During the Laramide revolution, which, in this region, began in latest Cretaceous time and was essentially over by the close of Eocene time, some basins were downwarped 30,000 ft. or more and some mountains upwarped at least 14,000 ft. Most basins were filled before the beginning of Oligocene time, and most mountain ranges were then partially buried by Oligocene, Miocene, and Pliocene strata. Some thrust faulting occurred along the SW. flank of the Gros Ventre Mountains during early Pliocene time. The highest plain of aggradation is thought to have been developed in central and northwestern Wyoming during middle Pliocene time and in eastern Wyoming in late Pliocene or early Pleistocene time. The later history of Wyoming was one of regional uplift, normal faulting, warping, and regional degradation, with minor local deposition of Pleistocene sediments.--Auth.

2-1416. Hooper, Kenneth. THE MARINE TERTIARY ROCKS OF BINNINGI AT THE NORTH END OF LAKE COWAN, WESTERN AUSTRALIA: Carleton Univ., Dept. Geology, Geol. Paper 59-3, 13 p., fold. map, 1959, 12 refs.

Scattered outcrops of Tertiary (presumed Eocene) marine fossiliferous rocks of calcareous and siliceous facies rest unconformably upon the Precambrian at the N. end of Lake Cowan, Western Australia. The fauna of lamellibranchs, gastropods, echinoids, bryozoans, sponge spicules, and foraminifers is poorly preserved. The rocks (of shallow-water type) are tentatively correlated with the Plantagenet beds of the S. coast of Western Australia and probably indicate the most northern extension of the Eocene marine transgression in this region.

The petrology and succession of the Precambrian rocks - sediment, metasediment, greenstone, granite, pegmatite, gabbro, dolerite, and acid porphyry - are briefly described.--Auth.

2-1417. Malhotra, Chamen L., and Edward R. Tegland. A NEW TERTIARY FORMATION IN HARDING COUNTY, SOUTH DAKOTA: South Dakota Acad. Sci., Proc., v. 38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 263-274, illus., map, 1959, pub. 1960, 2 refs.

A fossil assemblage recently collected in the southern portion of the Slim Buttes of Harding County indicates that the stratigraphic unit formerly mapped as Chadron is actually Duchasnean (basal Oligocene) in age. The authors propose the name Slim Buttes formation for this unit.

The configuration of surface of deposition, type of sorting and bedding, lack of basal conglomerate, position of clay beds, and presence of conglomerate and channel fills at the top of the Slim Buttes formation indicate that the sediments of this unit were probably deposited in a lake. At least, during the latter part of the period of deposition, the general direction of stream flow was from the NW.

Study of heavy mineral assemblages indicate that the sediments were largely derived from igneous rocks, probably either from the northern Black Hills or the Big Horn mountains. Direction of stream flow within the Slim Buttes formation, plus differences between mineral assemblages in the Black Hills and in the Slim Buttes samples, indicates that the northern Black Hills were probably not the source of sediments.--Auth.

2-1418. Golubeva, L.V. A STRATIGRAPHIC SCHEME FOR THE QUATERNARY DEPOSITS OF THE NORTHWESTERN PART OF THE WEST SIBERIAN LOWLANDS AND ITS PALEOFLORESTIC BASIS: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 43-52, 2 diags., pub. 1959, 19 refs.

A paleophytologic basis for stratigraphic subdivision of the Quaternary deposits is presented to aid in determining the number of glaciations, and also for correlation of the different Quaternary stratigraphic units of the northern west Siberian lowlands.

Spore-pollen analyses in conjunction with determinations of the macroscopic vegetation remains (seeds and diatoms) show a regular change in the constitution of the vegetation and, consequently, may be used as the basis for a stratigraphic classification. The changes in vegetation point to a rhythmic oscillation in the climate. Starting from the time of maximum (Samarovian) glaciation, 4 cool periods alternated with warmer intervals.--Auth.

2-1419. Steece, Fred. V. PLEISTOCENE VOLCANIC ASH IN SOUTHEASTERN SOUTH DAKOTA: South Dakota Acad. Sci., Proc., v. 38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 41-44, 2 illus., map, 1959, pub. 1960, 2 refs.

Volcanic ash or ashy silts have been collected from several localities in southeastern South Dakota, and are correlated with the Pearlette ash bed of Kansas. Criteria for identification are shard shape, color, presence of inclusions, and refractive index.

The Pearlette ash bed in Kansas lies in the Meade formation, and is late Kansan in age. In Nebraska the Pearlette ash is a member of the Sappa formation, also of late Kansan age. Because its stratigraphic position is known, the ash is an important late Kansan time marker to which other Pleistocene deposits may be related.--Auth.

2-1420. Mezhevik, A.A. GEOLOGIC DEVELOPMENT HISTORY OF NORTH KHARA-ULAKH: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 68-74, map, pub. 1959, 14 refs.

This paper attempts to review the geologic history of N. Khara-Ulakh in the light of new data obtained as the result of a geologic survey on the scale 1:200,000. These data lead to the following general conclusions:

1) The N. Khara Ulakh territory and the edge of the Siberian platform experienced 13 major marine transgressions during the Paleozoic, Mesozoic, and Cenozoic eras.

2) The character of the tectonic movements changed often. In early and middle Paleozoic time, predominantly undulatory movements of platform type occurred, which were followed by the accumulation of carbonate deposits. During late Paleozoic time, differential movements of large amplitude occurred. At that time, the upper Paleozoic flexure, filled with terrigenous deposits, was formed. During the Triassic, Jurassic, and Lower Cretaceous periods, N. Khara-Ulakh experienced mainly undulatory movements, which also affected the edge of the Siberian plateau.

3) After early Carboniferous time, the platform was replaced by the late Paleozoic geosyncline, the formation of which was accompanied by deep fractures.

4) During late Cretaceous time, due to plicative dislocations, the folding spread towards the W. at the expense of the platform.

5) The deposition of rocks in the early transgressive stages in the early Permian, early Triassic, early Jurassic and early Cretaceous periods is characterized not by coarse-grained beds, as is generally assumed, but by fine-grained, clay-bearing formations, which during the regressive phase were succeeded by coarse-grained formations.--Auth. concl.

2-1421. Fairbairn, H.W., Patrick M. Hurley, William H. Pinson, Jr., and R. F. Cormier. AGE OF THE GRANITIC ROCKS OF NOVA SCOTIA: Geol. Soc. America, Bull., v. 71, no. 4, p. 399-413, fold. map, 2 charts, diags., 3 graphs, 7 tables, Apr. 1960, 22 refs.

The ages of biotite, muscovite, and K feldspar in granitic rocks from 27 localities in Nova Scotia, as determined by A/K and Sr/Rb methods, fall in the range 300 to 400 million years. At 2 localities, where nearby intrusive contacts with upper Lower Devonian fossiliferous sedimentary rocks are established, the best estimate of the age of the intrusive rocks is  $365 \pm 20$  m.y. Biotite and K feldspar in granite from Boisdale Hills (Cape Breton Island), associated with nearby Middle Cambrian sedimentary rocks, have an estimated age of  $500 \pm 20$  m.y. A single age analysis of biotite in the metamorphosed Meguma series gives 340 m.y. Radiation-damage ratios for zircon at 9 localities where the age has been determined show 1 concordant result and 8 which are lower than the A/K and Sr/Rb ages by 16 to 33%.

The evidence of the mineral ages in the granitic rocks investigated thus far in Nova Scotia strongly suggests a protracted period of intrusion in the Devonian, with the upper Lower Devonian older than 365 m.y. Middle Cambrian is tentatively believed to be older than 500 m.y.

Holmes' B time scale, with Cambrian limits of 430 and 510 m.y. and Devonian limits of 255 and 313 m.y., would require an extension of 10-25% to meet the Nova Scotia data.--Auth.

2-1422. Kupsch, W.O. RADIOCARBON-DATED ORGANIC SEDIMENT NEAR HERBERT, SASKATCHEWAN: Am. Jour. Sci., v. 258, no. 4, p. 282-292, map, chart, 2 diags., Apr. 1960, 9 refs.

Organic sediments from a depth of 11 ft. near Herbert, Saskatchewan, contain willow wood which has been radiocarbon dated at  $10,050 \pm 300$  years B.P. The sediments contain in addition to the wood a faunule of predominantly snails and ostracodes.



plant fragments are common and include a few spruce cones. Marl with abundant remains of algae (Chara) is locally present. Spores and pollen recovered from the sediment provide an insight into the type of vegetation at the Herbert site 10,000 years ago.

Fauna and flora indicate sedimentation in quiet water of a pond or small lake in a spruce-pine park-

land with a few northern, broadleaved trees such as birch, willows, and probably aspen. Such an environment exists today in the forest belt of Saskatchewan, 200 mi. to the N. of Herbert. The change from forest to grassland, which now prevails near Herbert, is indicated by pollen analyses of the sediments overlying the radiocarbon dated wood.--Auth.

## 5. PALEONTOLOGY

See also: Areal and Regional Geology 2-1353; Geomorphology 2-1366; Stratigraphy 2-1401, 2-1418; Geochemistry 2-1526; Fuels 2-1597.

-1423. Beerbower, James R. **SEARCH FOR THE PAST, AN INTRODUCTION TO PALEONTOLOGY:** 62 p., 227 figs., 29 tables, Englewood Cliffs, New Jersey, Prentice-Hall, Inc., 1960, approx. 300 refs.

This book is an introduction to the principles, problems, and conclusions of paleontology for the geologist, biologist, and prospective paleontologist. The initial chapter describes the nature of paleontology and the methods of approach to paleontologic problems. The following chapter introduces the principles of descriptive, developmental, and functional morphology as they apply to fossils. The next topic is the relation of form to classification, and this chapter includes a survey of the major animal groups. The species is then approached from the population viewpoint, and the genetics, morphology, and ecology of species populations, both fossil and recent, are discussed. The distribution of species is considered in terms of habitat, historical factors, and interspecies relationships; the concepts developed are applied to paleoecologic problems. Both mechanisms and patterns of evolution are described and discussed with reference to paleontologic interpretation. The phylogenetic basis of classification is considered with examples of phylogenetic interpretation. Correlation is related to evolutionary sequence, to paleoecologic factors in distribution, and to practical problems of guide fossils, fossil zones, and time-stratigraphic classification. Succeding chapters deal with the major fossil groups and their living relatives, including the vertebrates. The phylogeny (and its morphologic basis), adaptations, and paleoecology of each group is discussed. Morphology is illustrated and described in accompanying glossaries; a simple classification is provided in a tabular form. The book concludes with a survey of human evolution and its place in the Cenozoic mammalian radiations.--Auth.

-1424. Simpson, George Gaylord. **NOTES ON THE MEASUREMENT OF FAUNAL RESEMBLANCE:** Am. Jour. Sci., v. 258-A (Bradley Volume), p. 300-11, 2 tables, 1960, 6 refs.

Of various measures of taxonomic resemblance, the percentage in the smaller of 2 faunal samples of the number of taxa common to both is most useful. It tends to eliminate the effects of discrepancy in size between the 2 faunas or samples, and when that discrepancy is a factor in the problem being studied the percentage of common taxa may be used. When faunas closely similar taxonomically are compared, it may be desirable to take into account differences in the relative abundances of taxa in common. For that purpose, measures based on rank correlation are suggested, but no such measure seems fully satisfactory. Several other measures used in current literature are discussed.--Auth.

2-1425. Hutchinson, G.E. **ON EVOLUTIONARY EURYHALINITY:** Am. Jour. Sci., v. 258-A (Bradley Volume), p. 98-103, 1960, 15 refs.

The Monera and many of the lower groups of Protista exhibit extreme evolutionary euryhalinity, and it is impossible to state whether the lowest groups of organisms are characteristically inhabitants of fresh or salt water. In the Metazoa, the capacity to invade fresh water from the ocean is very irregular and probably represents a statistically superdispersed system. Such a system could be explained formally by supposing that the various physiological preadaptations required to invade fresh waters are distributed with characteristic but different probabilities. Once the osmotic problems have been surmounted by an invader, the problem of salt uptake from very dilute solutions remains. It is probable that many animals of marine origin which are found associated with fresh-water species only in coastal localities have simply failed to develop really efficient salt uptake mechanisms, even though they are osmotically well adapted to fresh water. There is probably some correlation between small size and evolutionary capacity to change from one environment to the other. The evidence from the Caspian basin suggests that chemical adaptation is only one aspect of invasion into fresh water; it is suggested that the preexisting fresh-water fauna now constitutes the major barrier against invasions from the sea.--Auth.

2-1426. Nursall, J.R. **THE ORIGIN OF THE METAZOA:** Royal Soc. Canada, Trans., Sec. 5, Biol. Sci., v. 53, p. 1-5, 1959, 24 refs.

It may be assumed that at the time of the origin of life on earth 2 conditions existed: 1) a supply of organic matter was present as nutrient, formed probably by physico-chemical means from inorganic compounds, and 2) eobionts could utilize the nutrient by adsorption or absorption. Under such conditions it is probable that life need not have originated from one single primeval progenitor but could have evolved separately at many different times and in many different forms. This implies that several phyla had independent origins, and it does not follow that all phyla must ultimately be related.--M. Russell.

2-1427. Woodring, W.P. **PALEOECOLOGIC DISSONANCE; ASTARTE AND NIPA IN THE EARLY EOCENE LONDON CLAY:** Am. Jour. Sci., v. 258-A (Bradley Volume), p. 418-419, 1960, 6 refs.

Astarte (a chiefly circumarctic and boreal marine pelecypod) and Nipa (an Indo-Malayan palm) lived in the same region during early Eocene time when the London clay was deposited, whereas now they are separated by a gap of several thousand kilometers. Such examples of paleoecologic dissonance are more common than indicated by paleoecologic

generalities concerning Tertiary marine faunas.--Auth.

2-1428. Emerson, William K. RESULTS OF THE PURITAN-AMERICAN MUSEUM OF NATURAL HISTORY EXPEDITION TO WESTERN MEXICO. 11. PLEISTOCENE INVERTEBRATES FROM CERALVO ISLAND: Am. Mus. Nat. History, Am. Mus. Novitates, no. 1995, 6 p., fig., March 1960.

Pleistocene metazoan invertebrates are reported for the first time from Ceralvo Island, Baja California, Mexico. An assemblage totaling 34 species of megafossils, 32 mollusks, and 2 stony corals, was collected from a low emergent terrace on the southern end of the island. All the species are extant forms that live in this region of the Gulf of California at the present time.--Auth.

2-1429. Sutherland, Patrick K., and J.M. Cocke. A SOLITARY RUGOSE CORAL OF EXCEPTIONAL SIZE FROM THE MIDDLE PENNSYLVANIAN OF OKLAHOMA: Oklahoma Geology Notes, v. 20, no. 4, p. 78-82, 3 illus., Apr. 1960, 4 refs.

Specimens of the largest caninid coral known from the Pennsylvanian of North America are *Pseudozaphrentoides* sp., collected from the upper Des Moinesan Oologah formation, in Rogers County, Oklahoma. They are up to 10 in. long and 3 in. in diameter.--M. Russell.

2-1430. Cronoble, William R. AN OCCURRENCE OF ULOCINUS BUTTSI MILLER AND GURLEY IN OKLAHOMA: Oklahoma Geology Notes, v. 20, no. 4, p. 96-99, 3 illus., table, Apr. 1960, 5 refs.

An excellently preserved dorsal cup of *Ulocinus buttsi* Miller and Gurley was recently discovered in the Hogshooter formation (Missourian [Late Pennsylvanian]) of Nowata County, Oklahoma. Few well-preserved specimens of this species have been found; none of the previous descriptions of this species have been complete. The new Oklahoma specimen is described in detail; 16 of the more important measurements are tabulated. A comparison of the new specimen with the holotype is included. One photographic plate illustrating the posterior, anterior, and dorsal views of the specimen is also included.--Auth.

2-1431. Chenoweth, Philip A. STARFISH IMPRESSIONS FROM THE HILLTOP SHALE: Oklahoma Geology Notes, v. 20, no. 2, p. 35-36, 2 illus., Feb. 1960, ref.

Heretofore only a single specimen of a starfish has been reported in Oklahoma, this being an unidentified form found near Barnsdall. Several small starfish impressions are described and figured from the Hilltop shale (Missourian [Late Pennsylvanian]). The fossils occur in thin silty sandstones near the upper part of the formation and are found in a quarry near Wewoka, Seminole County. The fossils occur both as shallow depressions and as slightly raised molds. All are small, averaging 14 mm. from tip to tip. None can be identified, but they appear to belong to the subclass Asteroidea. Conditions for starfish preservation, as outlined by Schuchert, appear to have been well satisfied during deposition of the Hilltop shale, and it is thought likely that further search will yield more readily identified specimens.--Auth.

2-1432. Gill, Edmund D., and Kenneth E. Caster. CARPOID ECHINODERMS FROM THE SILURIAN AND DEVONIAN OF AUSTRALIA: Bulls. Am. Paleontology, v. 41, no. 185, 71 p., 8 illus., 10 pls., map, 2 charts (1 fold.), diag., March 1960, 70 refs.

The Silurian and Devonian carpod echinoderms described in this paper are the first representatives of the class to be recorded from Australia. A new genus and species of the Carpoidea Mitrata, *Victoriacystis wilkinsi* (Upper Silurian, Lower Silurian?), and Lower Devonian?) are diagnosed. The Carpoidea Soluta genus *Rutroclypeus* (type species *R. junori* Withers), formerly assigned to the Xiphosura, is reanalyzed and 2 new species, *R. victoriae* and *R. withersi*, (Lower Devonian all) described. The Rhenish Lower Devonian *Dendrocystites globulus* Dehm is provisionally referred to the genus *Rutroclypeus*, and definitely to the new family based thereon. A new classification of the Carpoidea is proposed: 2 new subclasses, Homostealea and Homoiostealea are defined, and 2 new superorders of the Homoiostealea, Stylophora and Astylophora, proposed. In the order Soluta, 2 new families, lowacystidae and Rutroclypeidae are diagnosed and the new genus of the Dendrocystitidae, *Heckericystis* (type species *Dendrocystites kuckersianus* Hecker), from the Russian Ordovician, proposed. Certain new names required by the Rules of Zoological Nomenclature are introduced. The distribution of the Carpoidea is considered, and the significance of the Australian forms for paleogeography is assessed.--Auth.

2-1433. Huffman, George G., and John M. Starke, Jr. SPIRIFER GRIMESI FROM THE ST. JOE LIMESTONE NEAR TAHLEQUAH, OKLAHOMA: Oklahoma Geology Notes, v. 20, no. 3, p. 50-52, 2 illus., March 1960, 15 refs.

The pedicle valve of a specimen of *Spirifer grimesi* Hall, 120 mm. wide and 80 mm. long, from the Lower Mississippian St. Joe limestone is described. The stratigraphic relations and faunal content of the formation are reviewed.--M. Russell.

2-1434. Valentine, James W. HABITATS AND SOURCES OF PLEISTOCENE MOLLUSKS AT TORREY PINES PARK, CALIFORNIA: Ecology, v. 41, no. 1, p. 161-165, 2 figs., Jan. 1960, 11 refs.

A molluscan fauna of 74 forms is recorded from an upper Pleistocene marine terrace carved on rocks on which sandy beaches are developed today. A large rocky-shore molluscan element is present, but the northern forms so characteristic of that element were not found. A recent thanatocoenose from a nearby sandy beach also contained rocky-shore forms, which seem to have come from the S. Perhaps the rocky-shore Pleistocene element was washed from the S. where warm surface waters may have piled up at a rocky headland (Soledad Mountain).--Auth.

2-1435. Ladd, Harry S. ORIGIN OF THE PACIFIC ISLAND MOLLUSCAN FAUNA: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 137-150, illus., 2 maps, 1960, 50 refs.

It has long been recognized that the marine mollusks of the Pacific islands are related to the fauna of Indonesia, there being few ties between the mollusks of the islands and those of western America. Indonesia has been regarded as the center of dis-



ersal of the "Indo-Pacific fauna," yet prevailing winds, currents, and even the major storm tracks tend from the islands toward Indonesia and the W.

Until recently all of the islands of the Pacific basin were thought to be geologically very young, but data obtained from submarine mapping, dredging, and drilling now indicate that: 1) a shallow-water fauna, including mollusks, was present in the area as early as middle Cretaceous; 2) rich Tertiary molluscan faunas (possibly richer than those of today) were widespread both inside and outside the Pacific basin proper; 3) many, if not all, of the hundreds of atolls that lie between Hawaii and Indonesia stood above the sea (some as high islands) at intervals during the Tertiary; 4) some 50 guyots, now far beneath the sea, projected into shallow water as additional stepping stones for the distribution of marine life, many of them in the broad gap SW. of Hawaii. These discoveries suggest that the Pacific islands once formed a giant archipelago and that the islands could have been the home of many elements of the Indo-Pacific fauna. Faunal migration, favored by winds and currents, was toward Indonesia rather than from

With some islands or parts of islands projecting above the sea at all times since the Cretaceous there could be a reasonable explanation for the ancient stocks of land shells and plants found in Hawaii and other volcanic islands. The newly discovered stepping stones appear to be satisfactory replacements for the land bridges once called for by many biologists. --Auth.

1436. Taylor, Dwight W. DISTRIBUTION OF THE FRESHWATER CLAM PISIDIUM ULTRAMONTANUM; A ZOOGEOGRAPHIC INQUIRY: *Am. Jour. Zool.*, v. 258-A (Bradley Volume), p. 325-334, illus., maps, 1960, 20 refs.

Pisidium ultramontanum Prime, a freshwater clam of the Sphaeriidae, is a relict species found living only in northeastern California and S.-central Oregon. Tertiary and Pleistocene fossils, however, document its former occurrence in the Snake River drainage as far E. as southeastern Idaho. Analysis of the distribution of other mollusks and fishes associated with P. ultramontanum reveals a common pattern, which can be traced by faunal links from one drainage basin to another. These links form a chain which extends from Walker Lake in western Nevada across Big Lake and the upper Pit River, California, to Klamath Lake, Oregon; thence across Fossil Lake and the Malheur basin, Oregon, to the Snake River; and through Gentile Valley and Bear Lake, Idaho, to Lake Powell, Utah. All of the relict mollusks and fishes of lakes in the northern Great Basin, Snake River drainage, and Klamath River drainage are related to this pattern of distribution. These lakes are therefore thought to have been connected, although not necessarily all at once. --Auth.

1437. Yochelson, Ellis L., and J. Thomas Dutro, LATE PALEOZOIC GASTROPODA FROM NORTHERN ALASKA: *U.S. Geol. Survey, Prof. Paper* 334-D, p. 111-147, 9 maps, 3 pls., 2 charts, g., table, 1960, 50 refs.

Late Paleozoic gastropods from northern Alaska occur in rocks of both Mississippian and Permian age; most of the fossils studied came from the Mississippian. Although the gastropods are of use for dating the rocks only in the broadest terms, locally they are useful in establishing informal faunal zones.

On the basis of predominant occurrence of certain gastropods, it is possible to distinguish rocks of Early Mississippian, Late Mississippian, and Permian age. Two divisions, based on the distribution of gastropods, are distinguished in the Upper Mississippian. One collection, consisting entirely of specimens of Glabrocingulum and Trepospira, may indicate the presence of rocks of Pennsylvanian age. Occurrence and distribution data are summarized in tabular form.

The gastropod faunule is composed primarily of euomphalaceans, platycerataceans, pleurotomariaceans, neritaceans, and bellerophonaceans, in that order of abundance. There is no indication that a distinct boreal fauna is represented. Many of the specimens are poorly preserved, though some well-preserved shells occur, particularly among the pleurotomariaceans.

Thirty-four species are recognized in the systematic treatment; 9 of these are formally named as new species and one is referred to a previously described species. One new pleurotomariacean genus, Nodospira, is described. At least 2 other new genera may occur in the faunule, but specimens are too incomplete for adequate taxonomic description. The occurrence of 2 scaphopod specimens is reported.

Most of the fossils were collected from the Brooks Range during fieldwork in connection with geologic investigations of Naval Petroleum Reserve no. 4 and adjacent areas. Some earlier collections from northern Alaska were restudied. --Auth.

2-1438. Pierce, W. Dwight. FOSSIL ARTHROPODS OF CALIFORNIA. NO. 22. A PROGRESS REPORT ON THE NODULE STUDIES: *Southern California Acad. Sci., Bull.*, v. 58, pt. 2, p. 72-78, 1959.

A preliminary report on the fossil findings from 17,203 nodules mostly collected in the Calico Mts., San Bernardino County, California; some from Mt. Pinos, Frazier Mts., Ventura County; 2 from Tehachapi Mts., Kern County; and from a hill at Lenwood, San Bernardino County. The insects are silicified, 3-dimensional, with evidence of instantaneous kill and preservation. The nodules are numbered, weighed, measured, and sketched before treatment, usually with 20% formic acid. When the digestion of the nodule has proceeded a week, the sludge is filtered, washed, transferred by brushes to alcohol, and examined under microscope. Specimens are lifted by brush into xylol and mounted on slides in balsam. The groups so far found and recorded by Palmer, Snyder, and the author, as well as the groups yet to be reported on, include Mollusca, Lymnaeidae; Ostracoda; Anostraca; Araneida; Acarina; Chelonethida; Scorpionida; Myriapoda; Collembola; Odonata, dragonflies and damselfly; Orthoptera; Plecoptera; Isoptera; Hemiptera; Lygaeidae, Miridae, Notonectidae, and Saldidae; Homoptera; Aphidoptera; Aleuroptera; Thysanoptera; Lepidoptera; Coleoptera, Dytiscidae, Staphylinidae, Cybocephalidae, Hydrophilidae; Hymenoptera, ants and wasp; Diptera, Heleidae, Tendipedidae; fish and fish eggs; bird feathers; mammal hairs; algae, mosses, and stems and leaf of higher plants. An average of 1.9% of all nodules have evidence of insects externally. The silicified insects occur throughout the nodules. --Auth.

2-1439. Frederickson, Edward A. NEW EVIDENCE CONCERNING DALMANITES OKLAHOMAE:

Oklahoma Geology Notes, v. 20, no. 3, p. 53-54, 2 illus., March 1960, ref.

The species *Dalmanites oklahomae* Richardson, based on a specimen found in the collections of the Chicago Natural History Museum, and assumed to be from the Silurian of Oklahoma, is revised on the basis of conspecific specimens from the Devonian Haragan formation of Oklahoma and assigned to the genus *Neoprobolium*. The corrected name of the species is *Neoprobolium oklahomae* (Richardson), age: lower Helderbergian.--Auth.

2-1440. Pierce, W. Dwight. SILICIFIED EGGS OF VERTEBRATES FROM CALICO MTS. NODULES: Southern California Acad. Sci., Bull., v. 58, pt. 2, p. 79-83, 3 illus., 1959.

In digesting Miocene nodules from the Calico Mts., California, for fossil insects, using 20% formic acid, there have been recovered from 3 quarter sections along Mule Cañon Drive, 26 silicified eggs, lacking amnion and showing the embryo in various stages of development. From the same area, 2 fish skeletons have also been found in nodules. The eggs measured 0.67 to 1.11 mm. (average 0.908 mm.) in length, and 0.34 to 0.63 mm. (average 0.50 mm.) in width. The fish skeleton measures 32.7 mm. in length (not complete), with thickness 9.9 mm.--Auth.

2-1441. Vallentyne, J.R. ON FISH REMAINS IN LACUSTRINE SEDIMENTS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 344-349, 3 tables, 1960, 5 refs.

The uppermost 15 cm. of the deep-water sediments of Little Round Lake, Ontario, contain an abundance of fish bones and scales mostly derived from *Perca flavescens*. The average numbers found per square meter of sediment surface were 16,000, 650, 400, and 205, for scales, vertebrae, arches (neural and hemal), and ribs, respectively. The deep-water sediments of 2 other lakes were devoid of fish remains. The number of scales found in the surface sediments of a fish hatchery pond was 6% of the number expected from fish that had died in the pond.--Auth.

2-1442. Dunkle, David H., and Stanley J. Olsen. DESCRIPTION OF A BERYCIFORM FISH FROM THE OLIGOCENE OF FLORIDA: Florida Geol. Survey, Spec. Pub. no. 2 (Contr. Florida Vertebrate Paleontology, Paper no. 3), 20 p., 4 pls., 1959.

Beryciform fishes (snappers and squirrel-fishes) were well established and of wide distribution around the world by upper Cretaceous times. Details of their origin and anatomy remain obscure due to the paucity of their remains in sediments in the western hemisphere. A specimen of *Holocentrites ovalis* from the Florida Marianna limestone, of Oligocene age, is described, figured, and compared with related forms.--S. J. Olsen.

2-1443. Brattstrom, Bayard H., and Ann Sturn. A NEW SPECIES OF FOSSIL TURTLE FROM THE PLIOCENE OF OREGON, WITH NOTES ON OTHER FOSSIL CLEMMYS FROM WESTERN NORTH AMERICA: Southern California Acad. Sci., Bull., v. 58, pt. 2, p. 65-71, illus., 4 maps, 1959, 7 refs.

New or noteworthy fossil pond turtles, genus *Clemmys*, from western North America are listed or described. This includes *C. owyheensis*, n. sp.,

from the Pliocene of Oregon and Pliocene (Los Angeles, Riverside, and Kern counties, California) and Pleistocene (Los Angeles, Kern, Santa Barbara, Shasta and Fresno counties, California, and White Bluffs, Washington) records of *C. marmorata*. *C. hesperia* Hay is shown to fall within the variation of *C. marmorata*. The phylogenetic and distributional relationship of the western U.S. species of *Clemmys* (*marmorata*, *saxea*, *morrisiae*, *bockmani*, and *owyheensis*) are cited.--B. H. Brattstrom.

2-1444. Etheridge, Richard. THE SLENDER GLASS LIZARD, *OPHISAURUS ATTENUATUS*, FROM THE PLEISTOCENE (ILLINOIAN GLACIAL) OF OKLAHOMA: Copeia, 1960, no. 1, p. 46-47, 6 refs.

A brief note on the presence of a single dentary of an unidentified species of *Eumeces* from the Cudahy fauna (Pleistocene, Kansan glacial) of southwestern Kansas, and a single vertebra of *Ophisaurus* from the Doby Springs local fauna (Pleistocene, Illinoian glacial) of northwestern Oklahoma. These 2 fossils represent only lizards obtained from glacial age faunas in this area.

The vertebra is discussed in detail and compared with *Ophisaurus attenuatus*, *O. compressus*, and *O. ventralis*. It is assigned to *Ophisaurus attenuatus*.

The 3 modern species of the glass lizard (*Ophisaurus*) are indicated as probably being differentiated at least by Illinoian glacial time, and comments are made concerning their fossil and Recent distribution.--J. J. Stephens.

2-1445. Harrell, Byron E. NOTES ON FOSSIL BIRDS FROM THE PLEISTOCENE OF KANSAS AND OKLAHOMA: South Dakota Acad. Sci., Proc., v. 38, (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 103-106, 1959, pub. 1960, 7 refs.

Fragmentary fossil bird bones from several Pleistocene sites in Kansas and Oklahoma were collected in 1954 by Dr. Claude W. Hibbard. A number of species can be identified. *Spatula clypeata* and *Fulica americana* are recorded from the Dixon local fauna, which is probably late Nebraskan in age. *Euphagus cyanocephalus* is identified from the Sanders local fauna, probably late Aftonian. *Mareca americana* is reported from the probably Sangamon Nye Sink locality in Oklahoma.--Auth.

2-1446. Marcus, Leslie F. A CENSUS OF THE ABUNDANT LARGE PLEISTOCENE MAMMALS FROM RANCHO LA BREA: Los Angeles County Mus., Contr. Sci. no. 38, p. 1-11, 2 figs., May 1960, 9 refs.

Ten species represented by 18 or more individuals are listed according to their abundance in the 7 major Los Angeles County Museum pits. Abundance was determined by counts of the best represented skeletal element (left or right side) of each species in each pit. The proportional relationship of the separate species in each pit is statistically analyzed, using the  $\chi^2$  test for deviation from similar proportional faunal composition among the several pits. The results show that although the pits had nearly the same fauna, they differed significantly in relative numbers of some of the species. The carnivores contribute the major part to the deviation. To explain the census differences, the hypothesis is advanced that the pits were not all continuously active for the same period of time.--Auth.



-1447. Stewart, T.D. FORM OF THE PUBIC ONE IN NEANDERTHAL MAN: Science, v. 131, p. 3411, p. 1437-1438, 3 illus., May 13, 1960, 7 refs.

Shanidar I and III from Iraq have the same peculiar form of pubis as Tabūn I from Palestine. These are the only such pubes known to exist. These facts suggest that Neanderthal man (Shanidar-Tabūn) and an early variety of modern man (Skhūl) coexisted during Mousterian times in this part of the world.--Ruth.

-1448. Jepsen, Glenn L. A NEW JERSEY MAS-ODON: 2d ed., New Jersey State Mus., Bull. 6, 10 p., 7 illus., 2 maps, diag., Feb. 1960, 14 refs.

The details of the 1954 discovery and subsequent restoration of "Matilda," a female *Mammut americanus*, from Pleistocene peat-bog deposits in Sussex County, New Jersey, are given against the background of a general review of what is known of mastodons. The find is unusual in that, although all the parts were recovered by means of dragline equipment in water, most of the hyoid bones were recovered. The skeleton was assembled at the American Museum of Natural History and put on permanent display at the State Museum in Trenton.--M. Russell.

-1449. Matthew, W.D., and J.R. Macdonald. TWO NEW SPECIES OF OXYDACTYLUS FROM THE MIDDLE MIOCENE ROSEBUD FORMATION IN WESTERN SOUTH DAKOTA: Am. Mus. Nat. History, v. 11, p. 203, 7 p., 5 figs., March 1960.

*Oxydactylus exilis* Matthew and *O. iacota* Matthew are described. These species were proposed by Matthew in an unpublished manuscript and are presented in this paper so that they might be properly entered in the literature.--J. R. Macdonald.

-1450. Wood, Albert E. A NEW SCIURAVID RODENT OF THE GENUS *PAUROMYS* FROM THE EOCENE OF WYOMING: Am. Mus. Nat. History, v. 11, p. 203, 7 p., 5 figs., Dec. 1959.

A new sciuravid rodent, *Pauromys schaubi*, from the upper Bridger is described and compared with *P. perditus* from the lower Bridger. It is considered to be somewhat more primitive than the earlier form. *Pauromys* is retained in the Sciuravidae as a possible cricetid ancestor, instead of including it (as does *Schaub*) in the Cricetidae. The only known genus that could be ancestral is the lower Eocene paramyid *Microparamys*.--Auth.

-1451. Hoffmeister, William S. SODIUM HYPOCHLORITE, A NEW OXIDIZING AGENT FOR THE PREPARATION OF MICROFOSSILS: Oklahoma Geology Notes, v. 20, no. 2, p. 34-35, Feb. 1960.

The Heard method of breaking down anthracite coals for separating out acid-insoluble microfossils involves the use of a 5.25% solution of sodium hypochlorite (i. e. Clorox) as the oxidizing agent. It is superior to Schulze's solution and the dry method in that it is cheap, fast, safe, gentle, breaks down anthracite coals as well as bituminous coals and carbonaceous shales, and it is a good bleaching agent.--M. Russell.

2-1452. Wilson, L.R., and R.W. Hedlund. TWO TECHNIQUES FOR STAINING HYSTRICHOSPHAERIDS: Oklahoma Geology Notes, v. 20, no. 4, p. 101-102, Apr. 1960.

Fossil hystrichosphaerids can be stained red with aqueous safranin in a solution of pH8 or higher when heated in a double boiler for several minutes. Green is obtained by placing the microfossils in glycerine jelly and then staining with malachite green powder and baking for several days at 40°C.--A. Nicholson.

2-1453. Todd, Ruth. RECENT LITERATURE ON THE FORAMINIFERA: Cushman Found. Foraminiferal Research, Contr., v. 10, pt. 4, p. 137-140, Oct. 1959.

Eighty-two briefly annotated references are given, arranged alphabetically by author. References are world-wide and are mostly from the period 1958-1959.

2-1454. Todd, Ruth. RECENT LITERATURE ON THE FORAMINIFERA: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 1, p. 42-45, Jan. 1960.

Sixty-one briefly annotated references are given, arranged alphabetically by author. References are world-wide and are mostly from the period 1958-1959.

2-1455. Todd, Ruth. RECENT LITERATURE ON THE FORAMINIFERA: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 2, p. 69-72, Apr. 1960.

Seventy-four briefly annotated references are given, arranged alphabetically by author. References are world-wide and primarily for 1959.

2-1456. Thalmann, Hans E. FORAMINIFERAL HOMONYMS: Cushman Found. Foraminiferal Research, Contr., v. 10, pt. 4, no. 201, p. 127-129, Oct. 1959.

Forty-six homonyms are listed, with references.

2-1457. Thalmann, Hans E. NEW NAMES FOR FORAMINIFERAL HOMONYMS IV: Cushman Found. Foraminiferal Research, Contr., v. 10, pt. 4, no. 202, p. 130-131, Oct. 1959.

Twelve new names of foraminiferal homonyms are proposed.

2-1458. Banner, F.T., and Walter H. Blow. SOME PRIMARY TYPES OF SPECIES BELONGING TO THE SUPERFAMILY GLOBIGERINACEAE: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 1, no. 204, p. 1-41, 8 pls., 2 charts, Jan. 1960, 88 refs.

An attempt is made to clarify the taxonomy of those species of the superfamily Globigerinaceae which possess primary types deposited in London and Paris. The lectotypes of the following species are described: *Globigerina bilobata* d'Orbigny, *G. bulloides* d'Orb., *G. bulloides borealis* Brady, *G. conglobata* d'Orb., *G. cretacea* d'Orb., *G. cristata* Heron-Allen and Earland, *G. dutertrei* d'Orb., *G. eggeri* Rumbler, *G. elongata* d'Orb., *G. helicina* d'Orb., *G. megastoma* Earland, *G. punctulata*

Deshayes, *G. punctulata* d'Orb., *G. quadrilobata* d'Orb., *G. rotundata* d'Orb., *G. rubra* d'Orb., *G. sacculifera* Brady, *Pullenia obliquiloculata* Parker and Jones, *Pulvinulina menardii fimbriata* Brady, *P. menardii tumida* Brady, *P. scitula* Brady, *Rotalia limbata* d'Orb., *R. menardii* Parker, Jones and Brady, *R. nitida* d'Orb., *Sphaeroidina bulloides* var. *dehiscens* Parker and Jones, and *Truncatulina humilis* Brady. Neotypes are proposed and described for *Globigerina conglomerata* Schwager, *G. seminulina* Schwager and *Rotalina cultrata* d'Orbigny. Descriptions of the morphology of the above species are emended, and their taxonomy, stratigraphy, and phylogeny are discussed. The remaining syntypes of *Globigerina depressa* d'Orbigny, *Rosalina linneiana* d'Orb. and *Rotalina canariensis* d'Orb. are discussed. The taxon "*Rotalia menardii* d'Orbigny" is shown to be unavailable. *Globotruncana mariei* is proposed as a new name for the homonym *Globotruncana cretacea* Cushman non *Globotruncana cretacea* (d'Orbigny). -- Auth.

2-1459. Banner, F. T., and Walter H. Blow. SOME PRIMARY TYPES OF SPECIES BELONGING TO THE SUPERFAMILY GLOBIGERINACEAE - A FURTHER TAXONOMIC NOTE: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 2, no. 206, p. 53, Apr. 1960.

Corrections are given for the authorship of 5 taxa discussed in the paper entitled "Some Primary Types of Species Belonging to the Superfamily Globigerinaceae" (see above). The conclusions of the paper are not affected by these corrections.

2-1460. Cole, W. Storrs. REVISION OF *HELICOSTEGINA*, *HELICOLEPIDINA* AND *LEPIDOCYCLINA* (POLYLEPIDINA): Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 2, no. 208, p. 57-63, 4 pls., Apr. 1960, 27 refs.

The genera *Helicostegina* and *Helicolepidina* and the subgenus *Polylepidina* of the genus *Lepidocyclus* are reviewed. New illustrations are given of many of the species discussed. The species *Helicostegina gyralis* Barker and Grimsdale is a synonym of *H. dimorpha* Barker and Grimsdale. The species previously known as *Helicolepidina polygyralis* Barker, *Helicolepidina paucispira* Barker and Grimsdale, and *Helicostegina soldadensis* Grimsdale are combined and assigned to the genus *Helicostegina* under the name *H. polygyralis* (Barker). Two species are recognized in the genus *Helicolepidina*, *H. spiralis* Tobler from the upper middle Eocene and *H. nortoni* Vaughan from the upper Eocene. The genus *Eulinderina* Barker and Grimsdale is a synonym of *Lepidocyclus* (Polylepidina). A brief statement on the probable evolution of the lepidocyclines and these other genera based on this revision is presented. -- Auth.

2-1461. Hanzawa, Shoshiro. THE FORAMINIFERAL SPECIES *FABIANIA CASSIS* (OPPENHEIM), IN JAPAN: Cushman Found. Foraminiferal Research, Contr., v. 10, pt. 4, no. 199, p. 119-122, pl., table, Oct. 1959, 22 refs.

Specimens referred to the European species, *Fabiania cassis* (Oppenheim), have been found in a basal conglomerate of Eocene age on Shikoku island, southwestern Japan. The various species of *Fabiania* are analyzed. All of the species are seemingly synonyms of *Fabiania cassis* which has a world-wide

distribution in the tropical belt. -- Auth.

2-1462. Hedley, R. H. NEW OBSERVATIONS ON *PELOSPHAERA CORNUTA*: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 2, no. 207, p. 54-56, 5 illus., Apr. 1960, 15 refs.

Apart from the test of *Pelosphaera cornuta* the animal itself is characterized by having a relatively thick organic wall which is modified at one part in the form of an oral region with a definite mouth. In many respects this region is remarkably similar to that known in *Gromia oviformis*.

When characters of the test alone are considered, the inclusion of *P. cornuta* in the family *Astrorhizidae* may appear to be justified, whereas, if characters of the protoplasmic body are taken into account a closer affinity to the *Saccamminidae* is indicated. -- Auth.

2-1463. Hofker, Jan. THE TAXONOMIC POSITIONS OF THE GENERA *BOLDIA* VAN BELLEN, 1946, AND *ANOMALINELLA* CUSHMAN, 1927: Cushman Found. Foraminiferal Research, Contr., v. 11, pt. 2, no. 205, p. 47-52, 11 illus., Apr. 1960, 12 refs.

The type species of *Boldia* van Bellen, 1946, was originally designated as *Rotalina lobata* Terquem. The specimens from the Montian of Bunde, Holland, upon which van Bellen based his genus were wrongly identified by him. They are referable to *Boldia madrugensis* Cushman and Bermúdez which should, therefore, be considered the type species of *Boldia*. A second species from the same Montian material, questionably called *Rotalia obesa* Terquem by van Bellen, is referable to *Boldia cubensis* Cushman and Bermúdez.

*Anomalinella rostrata* (Brady) differs from the *Planulina* group of Almaena, *Kelyphistoma*, *Planulinella* only in having embracing chambers on both sides. The so-called supplementary apertures near the margin, even in the last-formed chamber, are closed by porous plates, as is the case in all the other groups mentioned. Comparison with the type species of *Planulina*, *P. ariminensis* d'Orbigny, leads to the conclusion that the whole group is a very homogenous one and must be closely related to the type species of *Planulina*. A simple emendation of the genus *Planulina* enables us to place all these forms in that genus, thus greatly simplifying our system. -- Auth.

2-1464. Hofker, Jan. THE GENERA *EPONIDES*, *LACOSTEINA*, *NUTTALLIDES*, *PLANORBULINA*, AND *HALKYARDIA*: Cushman Found. Foraminiferal Research, Contr., v. 10, pt. 4, no. 198, p. 111-118, 27 illus., Oct. 1959, 19 refs.

Restudy of the type species of *Eponides* and *Alabamina* shows that the 2 genera are synonymous. *Lacosteina* belongs to the dentate Foraminifera. The coiled portion shows the features of *Conorboides*, being usually trochoid with the axis of coiling forming an angle with the triserial portion. Since the toothplates are also typical of *Conorboides*, *Lacosteina* probably represents a wild-growing form of that genus; the relationship between the 2 forms is similar to that between *Dyocibicides* and *Cibicides*. The type species of *Nuttallides* is nearly identical with Eocene *Asterigerina campanella* of Europe which differs only specifically from the type species of *Asterigerina*. *Nuttallides* is, therefore, a synonym of *Asterigerina*. The type species of *Planorbulina*



shows the characteristics of *Cymbaloporetta*, and the latter must be placed in the synonymy of the former. The family Cymbaloporetidae must be incorporated into the Planorbulinidae. The whole group is allied to Discopulvinulina. The type species of Halkyardia shows an umbilical structure which resembles that of Ferayina, but not that of Cymbaloporetta. It may be that Cymbalopora, Halkyardia, and Ferayina belong to one group of the Foraminifera, and it is certain that Cymbaloporetta with its allies belong to quite different one.--Auth.

1465. Herrick, Stephen M. SOME SMALL FORAMINIFERA FROM SHELL BLUFF, GEORGIA: *Bull. Am. Paleontology*, v. 41, no. 187, p. 113-120, 3 pls., map, May 1960, 2 refs.

Microfossiliferous material was collected from the uppermost part of the large oyster bed (Cooke's Unit 6, 1943 description) exposed at Shell Bluff, Burke County, Georgia, in Feb. 1955. From the material a total of 20 species was identified, 19 of which are illustrated. No new species were found. The age of these fossils is considered to be no older than uppermost Eocene and possibly younger. This famous locality, together with the type locality of the McBean formation [Eocene], is described and shown on the map.--Auth.

1466. Keach, John M. PRELIMINARY FORAMINIFERAL POPULATION COUNT IN UPPER NIOBRARA CHALK: *South Dakota Acad. Sci., Proc.*, v. 6 (South Dakota, State Univ., Bull., ser. 60, no. 6), p. 54-56, 1959, pub. 1960, 4 refs.

Three test samples of clean picked Foraminifera from 1 cc. of Niobrara [Upper Cretaceous] chalk, each averaging 290 forms and weighing 0.7 mg., were used to determine the presence of an estimated 14,797 foraminifers in a cubic centimeter.

Out of 2,417 specimens identified, the following genera are arranged in their order of decreasing percentage: *Globigerinella*, *Globigerina*, *Gumbelina*, *Loxostoma*, *Globotruncana*, *Bulimina*.

The fauna suggests an open sea, moderately deep sublittoral environment with a well oxygenated bottom. The dominant forms are planktonic and were accumulated by ocean currents in an environment similar to the foraminiferal oozes of the present oceans.--F. V. Steece.

1467. Barnard, Tom. SOME ARENACEOUS FORAMINIFERA FROM THE LIAS OF ENGLAND: *Cushman Found. Foraminiferal Research, Contr.*, v. 10, pt. 4, no. 203, p. 132-136, 10 illus. on pl., Oct. 1959, 23 refs.

A group of arenaceous Foraminifera is described from the Lias of England. Some species are recorded for the first time from England. The genera *Ammodiscus*, *Jaculella*, and *Protonina* are discussed.--Auth.

1468. Waller, Harry O., and William Polski. PLANKTONIC FORAMINIFERA OF THE ASIATIC HELF: *Cushman Found. Foraminiferal Research, Contr.*, v. 10, pt. 4, no. 200, p. 123-126, pl., map, graphs, Oct. 1959, 4 refs.

About 300 bottom samples from the Gulf of Tonkin to the Yellow Sea were examined for their content of planktonic Foraminifera. Faunal groups were established and related to thermal ranges and latitude. Average surface-water temperatures higher than

20°C.-27°C. yield greater than 75% normal sized, unbroken tests, whereas average surface-water temperatures less than 20°C.-27°C. produce mostly small, easily broken tests.

The distributions of some species suggest a zonation with depth. Coiling direction of *Globorotalia menardii* and *Globorotalia trigonula* is left whereas *Globigerina subcretacea* and *Pulleniatina obliquilocolata* coil right.

Four temperature restricted faunas are established: warm water restricted, 19°C.-30°C.; warm water preferred, 6°C.-30°C.; transitional, 4°C.-30°C.; cold water, 4°C.-26°C.--Auth.

2-1469. Cole, W. Storrs, Ruth Todd, and Charles G. Johnson. CONFLICTING AGE DETERMINATIONS SUGGESTED BY FORAMINIFERA ON YAP, CAROLINE ISLANDS: *Bull. Am. Paleontology*, v. 41, no. 186, p. 73-112, 3 pls., map, 3 tables, March 1960, 39 refs.

The Map formation of Yap contains a good fauna of smaller planktonic and larger Foraminifera. This association of 2 kinds of Foraminifera, found in 2 samples, is the first such association observed by which comparison is possible between known and widely distributed zones of both larger and smaller Foraminifera in the Pacific.

The larger Foraminifera indicate correlation with formations on Guam and Fiji that are dated as Tertiary f (Miocene), and that overlie rocks of Tertiary e (early Miocene) age. The planktonic Foraminifera indicate correlation with a formation on Saipan that was provisionally dated as Tertiary d (late Oligocene) and that unconformably underlies rocks of Tertiary e age.

The Map formation is a variable, irregularly bedded, coarse, and unsorted breccia and conglomerate with distorted lenses of finer material and channels filled with stratified sandstone. The conglomeratic nature of the Map formation makes reworking of the globigerinids seem a reasonable explanation of the conflicting age indicated by the 2 parts of the assemblage. However, that explanation is not wholly satisfactory, as no source for the older Foraminifera has yet been found on Yap. Local expansion of ranges of index species or changes of sequence of zones that seem to apply consistently elsewhere in the Pacific are not acceptable explanations.--Auth.

2-1470. Bé, Allan W.H. SOME OBSERVATIONS ON ARCTIC PLANKTONIC FORAMINIFERA: *Cushman Found. Foraminiferal Research, Contr.*, v. 11, pt. 2, no. 209, p. 64-68, 3 illus., table, Apr. 1960, 10 refs.

*Globigerina pachyderma* (Ehrenberg) is believed to be the only species of planktonic Foraminifera living in the present central Arctic Ocean. It is thin-shelled and possesses a large aperture in its early stages (identified as small specimens of *Globigerina eggeri* or *G. bulloides* by previous workers), which contrast with features attained in its mature stages. The typical form of *G. pachyderma*, living below approximately 200 m. depth, is attained by crystalline thickening of the test and the addition of a reduced final chamber.--Auth.

2-1471. Johnson, Curtis L. MICROFOSSILS OF THE GREGORY SHALE MEMBER OF THE PIERRE FORMATION: *South Dakota Acad. Sci., Proc.*, v.

38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 49-52, 1959, pub. 1960, 6 refs.

Preliminary investigation of the microfauna of the Gregory shale [Upper Cretaceous] has disclosed the presence of species representing 10 different families of Foraminifera and 1 family of Ostracoda. In addition, some radiolarians are present in the upper Gregory.

Most of the foraminiferal tests are calcareous, but a few arenaceous forms are present. The fossil fauna represents benthonic and pelagic forms.

Ecologically, the genera suggest relatively shallow, warm seas of normal marine to brackish salinity, perhaps slightly on the alkaline side as indicated by pH evidence.

The Gregory shale shows possible correlation with the Taylor marl of Texas (Campanian) on the basis of foraminiferal evidence.--Auth.

2-1472. Lohman, Kenneth E. THE UBIQUITOUS DIATOM - A BRIEF SURVEY OF THE PRESENT STATE OF KNOWLEDGE: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 180-191, table, 1960, 20 refs.

Diatoms have distinctive and highly diversified tests or shells, similar to but not identical with opal in composition, accounting for their preservation as fossils in many kinds of sedimentary rocks and their importance to the geologist. The problem of unequivocally assigning diatoms and related minute unicellular or noncellular organisms to either the plant or the animal kingdom is increasingly being met by referring them to the Protista, which may be considered either as a separate kingdom of organisms or a confession of ignorance.

The earliest known large assemblage of marine diatoms occurs in rocks of Late Cretaceous age in California; the earliest known assemblage of non-marine ones occurs in middle Oligocene rocks in Colorado. They are widespread over the globe, living today in virtually all nontoxic waters, occupying a wide variety of habitats. This wide distribution is explained by the ease with which these minute organisms may be transported over great distances in viable condition by means of water, wind, and attachment to larger organisms. The preservation of diatoms in sedimentary rocks is governed largely by the pH of the solutions penetrating the sediment, as the solubility of the diatom silica increases rapidly with increasing pH above 5. Even under such conditions diatoms may be preserved in calcareous concretions formed before the advent of the alkaline solutions.

The percentage of extinct species in any assemblage of diatoms increases with age of the enclosing sediment at rates comparable with those of mollusks, making them useful indicators of geologic time. The species in any assemblage still represented in living assemblages form the basis for paleoecological interpretations. The fact that large and representative assemblages of diatoms can be obtained from very small samples is of equal importance in both lines of investigation. The small size of the diatoms and their resulting ease of transportation introduces problems of reworking and contamination. The former is often difficult to evaluate; the latter can be virtually eliminated by systematic care and cleanliness in the collection and preparation of the samples.--Auth.

2-1473. Beck, Charles B. CONNECTION BETWEEN ARCHAEOPTERIS AND CALLIXYLON: Sci-

ence, v. 131, no. 3412, p. 1524-1525, May 20, 1960, 5 refs.

Characters of 2 Upper Devonian genera, *Archaeopteris*, often considered to be a fern, and *Callixylon*, classified with the gymnosperms, have been recognized in a single specimen.--Auth.

2-1474. Huffman, George G., and John M. Starke, Jr. A NEW FOSSIL PLANT LOCALITY IN THE SYLAMORE MEMBER, CHATTANOOGA FORMATION, NORTHEASTERN CHEROKEE COUNTY, OKLAHOMA: Oklahoma Geology Notes, v. 20, no. 4, p. 89-91, 2 illus., map, Apr. 1960.

A new fossil plant locality in the Mississippian Chattanooga formation contains the shallow molds of tree trunks, tentatively identified as *Cordaites*.--M. Russell.

2-1475. Scott, Richard A., Elso S. Barghoorn, and Estella B. Leopold. HOW OLD ARE THE ANGIOSPERMS?: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 284-299, 1960, 77 refs.

The currently known paleontological evidence bearing on the time and appearance of the angiosperms in the fossil record is examined. Special attention is directed to the very considerable body of data derived from palynological studies made during the past decade in various parts of the world. It is concluded that no bona fide angiosperm remains, either megafossil or microfossil, have yet been described from rocks older than Early Cretaceous sediments. Their record, however, is extraordinarily and increasingly well-documented in post-Early Cretaceous sediments from the major geographic areas of the earth. Considering the efficiency of the atmospheric and fluvial media of transport of plant microfossils, and the well known instances of long distance transport of microfossils, it seems highly implausible that a major, evolving unit of the earth's terrestrial flora could have greatly antedated the Cretaceous without detection. The morphological problems in identification of "early" angiosperms and difficulties in determining their first appearance are briefly considered.--Auth.

2-1476. Wolfe, Jack A., and Elso S. Barghoorn. GENERIC CHANGE IN TERTIARY FLORAS IN RELATION TO AGE: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 388-399, 2 diag., table, 1960, 41 refs.

The question of an orderly, systematic change in generic composition with age of Cenozoic floras is reevaluated on the basis of 48 well-documented floras from the western United States. These data reinforce the conclusion that there is an orderly, progressive increase in modernity of fossil angiospermous floras with decreasing age. The changing ratio of native to exotic genera in fossil floras, when compared to the modern distribution of these same genera, is of such consistency, as a function of geologic time, that generic analysis of temperate zone floras preserved either as megafossils or microfossils provides in itself a reliable adjunct in the determination of age. The age-generic change is viewed as a reflection of broad regional and essentially unidirectional climatic change from late Eocene to Pliocene. Data on the physiognomic characteristics of fossil leaves is correlated with the age of fossil floras and change in generic composition.--Auth.



2-1477. Kremp, G.O.W., H.T. Ames, and A.J. Kovar. CATALOG OF FOSSIL SPORES AND POLLEN. VOLUME 9. TRIASSIC AND JURASSIC SPORES AND POLLEN: 163 p., illus., University Park, Pennsylvania, Pennsylvania State University, College of Mineral Industries, Dept. of Geology, Palynological Laboratories, 1960, refs.

One hundred and fifty-six specific or varietal descriptions of Triassic and Jurassic types from Europe and North America are presented in Vol. 9.

Contents:

- Daugherty, L.H., 1941. The Upper Triassic Flora of Arizona, p. 1-4, 4 new descriptions.  
 Krausel, R., and G. Leschik, 1955. Die Keuperflora von Neuwelt bei Basel, II. Die Iso- und Mikrosporen, p. 5-150, 146 new descriptions.  
 Erdtman, G., 1948. Did Dicotyledonous Plants Exist in Early Jurassic Time?, p. 151-152, 2 new descriptions.  
 Lantz, J., 1958. Étude des Spores et Pollens d'un Échantillon Purbeckien de L'Ile d'Oléron, p. 153-156, 4 new descriptions.

2-1478. Wilson, L.R. FLORINITES PELUCIDUS

AND ENDOSPORITES ORNATUS, WITH OBSERVATIONS ON THEIR MORPHOLOGY: Oklahoma Geology Notes, v. 20, no. 2, p. 29-33, pl., Feb. 1960, 10 refs.

Thin sections of the Pennsylvanian fossil Cordaitanthus shuleri from an Iowa coal ball contain spores which in the dissociated state are known as Florinites, a fossil spore genus. Outside the pollen sacs of the Cordaitanthus sections are spores of Endosporites ornatus. The sections illustrate morphological differences in the 2 genera. The germinal surface of Florinites is in the distal polar hemisphere and that of Endosporites is in the proximal polar hemisphere. The bladder of Endosporites completely encloses the central body, while in Florinites the bladder is attached slightly above the equator in the distal polar hemisphere. The germinal area of the Endosporites is a double wall consisting of the spore body wall and the bladder wall, which are fused in most of the proximal hemisphere. The germinal area of Florinites is a single layer as the bladder does not extend much beyond the equator of the spore body in the distal polar hemisphere. Endosporites has been found to be related to the lycopods and Florinites to the Cordaitales.--Auth.

## 6. GEOPHYSICS

See also: Geologic Maps 2-1312, 2-1314 through 2-1343; Geologic and Regional Geology 2-1353; Structural Geology 2-1383, 2-1384; Stratigraphy 2-1407; Geochemistry 2-1510, 2-1516; Mineral Deposits 2-1580.

2-1479. Burkard, Richard K. GEODESY FOR THE LAYMAN: 77 p., maps, diags., St. Louis, Missouri, U.S. Air Force, Aeronautical Chart and Information Center, Oct. 1959.

The use of intermediate and long range rockets and missiles requires that the distance and direction from launch site to target be accurately known. The science of geodesy can help to provide this information.

The basic principles of geodesy and geodetic surveying are presented in an elementary form. The formation of geodetic datums is introduced, and the necessity of connecting or joining datums is discussed. Methods used to connect independent geodetic systems to a single world reference system are discussed. Particular emphasis is placed on the U.S. Air Force role in developing a World Geodetic system.--Auth.

2-1480. Lum, Daniel. GRAVITY METERS (GRAVIMETERS) USED IN GEOPHYSICAL EXPLORATION: South Dakota Acad. Sci., Proc., v. 38 (South Dakota, State Univ., Bull., ser. 60, no. 2), p. 29-31, 1959, pub. 1960, ref.

The variation of the Earth's gravitational attraction at different locations was first observed by Jean Richer in 1672 with a pendulum-operated clock, and it was with pendulums that the first relative values of gravity were measured. Gravity meters were developed later and have been used intensively during the past 30 years for oil exploration. Because gravimeters can measure changes in gravity of less than 1 part in 1 million of the Earth's total gravity, their sensitivity creates many difficult instrumental problems, which are briefly discussed. However, gravimeters used in the more accurate geodetic work today have an even greater reading

accuracy of 1 part in 1 billion of the Earth's total gravity, as a result of instrument research and development to meet even more exacting requirements in precision.--Auth.

2-1481. Whitworth, Virgil L., Edward F. Haye, and Thomas M. Lindholm. GRAVITY-PHOTOGEOLOGY METHOD BOOSTS ACCURACY, CUTS COSTS: World Oil, v. 150, no. 5, p. 99-100, 104, 3 ports., illus., 2 maps, Apr. 1960.

Gravity surveys can be greatly improved by combining with photogeology and surface geologic data, particularly where there is alluvial material present. Densities of masses closest to the meter affect the readings most critically. As an example, an anomaly supposedly favorable according to the meter was followed as customary with an expensive seismic survey, after which it was discovered that this anomaly was due to alluvium at less than 80 ft. instead of the reported 8,000 ft. Other significant surface features determinable by photo or surface geology will lead to much improvement in gravity mapping and interpretation.--K.M. Willson.

2-1482. Bancroft, A.M. GRAVITY ANOMALIES OVER A BURIED STEP: Jour. Geophys. Research, v. 65, no. 5, p. 1630-1631, May 1960.

The expression  $\underline{d} = \frac{t}{\exp(t/d_0) - 1}$  is derived for

the case of a semi-infinite step of depth  $\underline{d}$  and thickness  $\underline{t}$ . As  $\underline{t}$  decreases to very small values,  $\underline{d}$  increases and tends to the value  $\underline{d}_0$ . The maximum depth to the step is therefore  $\underline{d}_0$ ; the actual depth is likely to be considerably less than  $\underline{d}_0$  for reasonable values of density. The advantage of the formula for a step is that in cases where it is applicable it gives a smaller value of  $\underline{d}_0$  than either of the formulas of Bott and Smith, and the relation between actual depth and  $\underline{d}_0$  is very easily applied.

The expressions for  $\underline{d}$  and  $\underline{d}_0$  were originally

derived for use in the interpretation of the Bouguer anomalies near Holleford, Ontario, where a circular depression in the Precambrian basement is thought to represent a meteor crater. The maximum depth and thickness of the material filling the depression (shown by the negative anomaly to be lighter than the rocks exposed at the surface) were estimated to be 300 ft. and 650-1,600 ft., respectively; subsequent drilling to 1,100 ft. failed to reach the undisturbed Precambrian.--D. B. Vitaliano (courtesy Geophysical Abstracts).

2-1483. Woollard, George P., Ned A. Ostenso, and Edward C. Thiel. GRAVITY ANOMALIES, CRUSTAL STRUCTURE, AND GEOLOGY IN ALASKA: Jour. Geophys. Research, v. 65, no. 3, p. 1021-1037, 7 figs. incl. 2 maps, secs., 4 tables, March 1960, 18 refs.

A regional Bouguer gravity anomaly map of Alaska is interpreted in terms of variations in crustal structure and local geology. The crustal results are compared with those determined from seismic measurements in 2 areas in Alaska and one in Alberta, Canada. These comparisons indicate the importance of including crustal composition in the derivation of crustal thickness from gravity values and suggest that regional isostatic anomalies are probably more indicative of abnormal crustal composition than of abnormal crustal thickness. In the Prince William Sound area of Alaska the seismic measurements show that the positive anomaly associated with the area is related to a thick crust having a high density rather than to a thin crust as might be deduced on the assumption of a normal crustal density. The inverse effect noted in the common association of negative anomalies with granitic intrusions appears to prevail in southern Alaska, where the seismic value of crustal thickness over the Coast Range batholith is less than that which would be deduced from the gravity anomaly values. Although the data are too sparse to permit any extended study of local geologic features, several areas of local gravity anomaly are defined. In general, these can be correlated with known geologic features such as sedimentary basins, granite on basic rock intrusions, and the configuration of the buried crystalline rock surface.--Auth.

2-1484. Thiel, Edward C., A. P. Crary, Richard A. Haubrich, Jr., and John C. Behrendt. GRAVIMETRIC DETERMINATION OF OCEAN TIDE, WEDDELL AND ROSS SEAS, ANTARCTICA: Jour. Geophys. Research, v. 65, no. 2, p. 629-636, 7 figs. incl. map, graphs, 2 tables, Feb. 1960, 5 refs.

The use of the gravity meter for measurement of ocean tides is illustrated by studies on the floating ice shelves of Antarctica. The observations are complicated by high-frequency oscillations of the ice, attributed to oceanographic influences. Factors involved in the reduction of the gravimetric data are analyzed. Amplitude and phase are computed for the more significant tidal components, and the energy spectra from 0.03 to 4 cycles per day are presented. The Weddell Sea tide has both diurnal and semi-diurnal components. The Ross Sea tide is diurnal, with the solar component predominating. The tidal range is greater in the Weddell Sea than in the Ross Sea. Correlation of tidal currents with changes in surface elevation provides an estimate of the inward dimension of the Ross Ice Shelf.--Auth.

2-1485. Balsley, James R., and A. F. Buddington. MAGNETIC SUSCEPTIBILITY ANISOTROPY AND FABRIC OF SOME ADIRONDACK GRANITES AND ORTHOGNEISSES: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 6-20, map, diag., table, 1960, 4 refs.

The magnetic susceptibility anisotropy of about 150 specimens of granites and orthogneisses of the NW. Adirondack area has been determined. The maximum susceptibility of the orthogneisses is uniformly about parallel to the lineation, and the minimum susceptibility is characteristically about at right angles to the planar structure. A high ratio of maximum to intermediate magnetic susceptibility indicates a high intensity of linear structure. A high ratio of the mean of the maximum and intermediate to the minimum magnetic susceptibility indicates a high intensity of planar structure. Studies of magnetic susceptibility anisotropy have permitted the determination of lineation in many rocks where it could not be determined in the field and have afforded quantitative bases for correlations with broad geologic structures.--Auth.

2-1486. Kobets, N.V., and B.V. Komarov. SOME PROBLEMS OF METHODOLOGY IN PROSPECTING FOR PRIMARY DIAMOND DEPOSITS BY AERO METHODS: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 80-86, 4 illus., pub. 1959, 3 refs.

English translation of GeoScience Abstracts 1-434.

2-1487. Chaney, P.E., F.M. Mayes, Jack W. Jones, and John Bennett. DRILL STEM LOGGING TOOL (In: McGrain, Preston, and Thomas J. Crawford, eds. Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959: Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 61-66, illus., logs, 1959) ref.

A new tool has been developed to permit electric logging of freshly drilled formations. The device consists of a self-contained recording mechanism which is pumped down the drill pipe whenever a log is desired. An electrode assembly projects out through one of the mud passages of a jet bit. Logs are recorded during the normal process of coming out of the hole, or the tool may be recovered on a wire line. A sample log obtained with this tool is presented for comparison with a conventional log of the same section.--Auth.

2-1488. Campbell, Orton E. WHY SEDIMENTARY STRUCTURES SHOW HIGH SELF-POTENTIALS: World Oil, v. 150, no. 5, p. 97-98, port., 3 maps, Apr. 1960, 5 refs.

Mineralized halos are often associated with geologic structures and/or oil and gas accumulations for which no satisfactory explanation of the origin has been offered. It is fact that electrical earth currents, operating by electrolysis with differences of self-potentials, are present which are capable of producing the anomalies. Geological structure has a natural electrical field with the highest potential at the top of the structure. Stratigraphic traps are special and unique problems. Petroleum anomalies originate perhaps by upward rising migratory hydrocarbons which do not necessarily come from the petroleum accumulations. Hydrostatic and thermal



forces may be the ones which carry the minerals, but self-potential gradients are forces which determine the halo patterns.--K. M. Willson.

1489. Byerly, Perry. EARTHQUAKE MECHANISMS: Science, v. 131, no. 3412, p. 1493-1496, illus., 3 diags., May 20, 1960, 17 refs.

Theory going back to Stokes and Love indicates that a sudden force applied to a part of an infinite homogeneous elastic medium sends out waves of longitudinal and transverse types. The directional character of at least the beginnings of these waves is preserved at a distance. This character depends on the nature of the force system applied at the source. It is widely agreed that the energy released by seismic waves was gradually stored as potential energy of strain before the shock. The moot question is just how this strain energy is transformed into the kinetic energy of wave motion.--Auth. summ.

1490. Aki, Keiiti. STUDY OF EARTHQUAKE MECHANISM BY A METHOD OF PHASE EQUALIZATION APPLIED TO RAYLEIGH AND LOVE WAVES: Jour. Geophys. Research, v. 65, no. 2, p. 729-740, 4 figs. incl. map, table, Feb. 1960, 15 refs.

Rayleigh waves and Love waves are used for the study of the earthquake mechanism by the application of a method of phase equalization. In this method, an impulse response is computed from known phase-velocity data and instrument characteristics and is cross-correlated with an actual record. A comparative study of Love waves from Kern County [California] aftershocks of 1952 and those from Nevada shocks of 1954 strongly supports the hypothesis of a pair of couples rather than a single couple for the earthquake source. Source functions for 5 Kern County aftershocks are derived from the Rayleigh waves recorded at Weston and Palisades. It was found that the sense of principal motion in the source function is in agreement with the fault-plane solution obtained from the P-wave data. Mantle Rayleigh waves are found to be useful for this purpose also.--Auth.

1491. Gutenberg, Beno. THE SHADOW OF THE EARTH'S CORE: Jour. Geophys. Research, v. 65, no. 3, p. 1013-1020, 4 figs., 3 tables, March 1960, 10 refs.

There is very good agreement between the observed decrease in amplitudes of longitudinal waves diffracted at the boundary of the earth's core and corresponding theoretical results. Especially, no rapid decrease of amplitudes at the beginning of the shadow zone is to be expected. Theory and observations show that the amplitudes of diffracted short-period P waves decrease faster with distance in the shadow zone than those of longer waves. At epicentral distances of over  $110^\circ$ , short-period diffracted P waves emerge gradually, and their beginning can rarely be ascertained. Corresponding long-period waves arrive within the limits of error at the time calculated on the assumption of a straight-line travel-time curve.

Amplitudes of waves diffracted from a caustic increase rapidly with distance; this usually limits the range of their observation to roughly  $10^\circ$ . On the contrary, amplitudes of P waves diffracted at the core may be visible on records of a great earthquake at distances of  $103^\circ$  to  $180^\circ$ . Observations for diffracted S waves cover a much smaller range of dis-

tances, since S waves appear in a portion of seismograms which is disturbed by earlier motion. Otherwise, observations indicate similar behavior for diffracted P and S waves.--Auth.

2-1492. Shurbet, D.H. THE EFFECT OF THE GULF OF MEXICO ON RAYLEIGH WAVE DISPERSION: Jour. Geophys. Research, v. 65, no. 4, p. 1251-1255, map, 3 graphs, 2 tables, Apr. 1960, 10 refs.

The dispersion of Rayleigh waves crossing the Gulf of Mexico was measured. It is shown that the observed dispersion can be accounted for by considering the presence of unconsolidated sediments. The average sediment thickness required to account for the dispersion is 7 km.--Auth.

2-1493. Helburn, Nicholas. SOUTHWESTERN MONTANA EARTHQUAKE AND LANDSLIDE: Geog. Rev., v. 50, no. 1, p. 109-111, Jan. 1960.

On Aug. 17, 1959, a major earthquake shook southwestern Montana, eastern Idaho, and northwestern Wyoming; its focus is tentatively located at  $44^\circ 30'N.$ , depth 25 mi., intensity 7.8 on Richter scale. Three important geologic aspects of the earthquake are faults and associated scarps, the behavior of Hebgen Lake, and landslides. Hebgen fault and the Red Canyon fault were active. Although the Hebgen Lake dam was not raised, the capacity of the lake was increased, for reasons not yet determined. Local ground-water activity and geyser activity in Yellowstone Park were temporarily stimulated. The Madison Canyon landslide blocked the Madison River, thus requiring intensive engineering efforts to prevent a disastrous washout. The U. S. Forest Service has set aside the area as one of national scientific interest.--M. Russell.

2-1494. Hardy, Clyde T., and Grant I. Gaeth. FIELD INVESTIGATION OF UTAH EARTHQUAKE, MAY 23, 1953: Utah Acad. Sci., Arts, & Letters, Proc., v. 36, 1958-1959, p. 137-140, map, pub. 1959, ref.

An isoseismic map of the Utah earthquake of May 23, 1953, was constructed for the purpose of locating the center of the disturbance. The epicenter had previously been reported in a newspaper account as probably in the mountains E. of Lehi, Utah. Coordinates, based on this general location, were later listed in the annual earthquake summary of the U. S. Coast and Geodetic Survey. Interrogation of numerous persons, one week after the quake, showed the epicenter to be about 3 mi. W. of Lehi, Utah, near the northwestern corner of Utah Lake. The active fault is probably associated with hot springs at Saratoga Resort.--Auth.

2-1495. Hardy, Clyde T. FIELD INVESTIGATION OF UTAH EARTHQUAKE, FEBRUARY 4, 1955: Utah Acad. Sci., Arts, & Letters, Proc., v. 36, 1958-1959, p. 141-143, map, pub. 1959, ref.

An isoseismic map of the Utah earthquake of Feb. 4, 1955, made by interrogating about 40 persons on the day after the quake, shows the epicenter to have been in the northwestern part of Salt Lake City.

The trend of the fault responsible for the earthquake bears about  $N. 21^\circ E.$  and may be projected northward to a point about 2,000 ft. S. of Becks Hot

Springs. The fault and hot springs are probably associated.--From auth. introd. & concl.

2-1496. Ambraseys, Nicholas N. THE SEISMIC SEA WAVE OF JULY 9, 1956, IN THE GREEK ARCHIPELAGO: Jour. Geophys. Research, v. 65, no. 4, p. 1257-1265, 3 maps, 2 graphs, 2 tables, Apr. 1960, 17 refs.

The earthquake of July 9, 1956 (magnitude 7 1/2, epicentral intensity VIII-IX MM), with epicenter at 36°54'N. 26°12'E., was followed by a severe seismic sea wave which originated near 36°04'N., 26°09'E. The wave was probably produced by a series of landslides on the steep banks of the submarine trench of Amorgos. The amplitude of the wave near its source was 100 ft., and the agitation of the sea within the central area lasted several days. A list of seismic sea waves in the Greek archipelago and adjacent seas is given in the appendix.--Auth.

2-1497. Kisslinger, Carl. SEISMOGRAMS ASSOCIATED WITH THE NEAR PASSAGE OF TORNA-DOES: Jour. Geophys. Research, v. 65, no. 2, p. 721-728, 6 figs. incl. map, Feb. 1960, 5 refs.

The principal features found on seismograms obtained on 3 occasions when a tornado passed near the station are described. Long periods, 50 sec. and greater, predominate, with periods of 30 sec. and less and bursts of very short period activity superimposed. The classical solution for a vertical load on an elastic half-space is applied, and the results indicate that elastic tilts due to the static loading effect of the tornado contribute substantially to the basic form of the record. The shorter periods are ascribed to the dynamic loading of the high winds on obstacles in the path. Part of the activity is not ground motion but is due to the buoyancy effect caused by the rapid pressure fluctuations.--Auth.

2-1498. Gilbert, Freeman, and Gordon J.F. MacDonald. FREE OSCILLATIONS OF THE EARTH. I. TOROIDAL OSCILLATIONS: Jour. Geophys. Research, v. 65, no. 2, p. 675-693, 13 figs., incl. diags., graphs, 8 tables, Feb. 1960, 32 refs.

The free periods of toroidal oscillations of the earth have been computed for 2 earth models. The lowest period for the Gutenberg model earth is 2,651 sec. and for the Jeffreys-Bullen model 2,732 sec. The surface amplitudes of the oscillations have been computed for 3 kinds of delta function stress sources - a unit force, a unit couple, and a unit torque - at depths of 600, 250, 100, and 30 km. The amplitudes decrease with increasing depth of the source. For a unit couple at 600 km. the maximum amplitude of the lowest period for the Gutenberg model is  $1.59 \times 10^{-25}$  cm., and for the Jeffreys-Bullen model it is  $0.70 \times 10^{-25}$  cm. By using the free periods of oscillation we have extended Love wave phase velocity dispersion curves to include long-period Love waves.

The method used to compute the periods and amplitudes of the free oscillations is an extension of the Thomson-Haskell matrix method used in plane layered media.

An example is presented to show the correspondence between the free oscillations and ray theory.--Auth.

2-1499. Band, William. STUDIES IN THE THEORY OF SHOCK PROPAGATION IN SOLIDS: Jour. Geo-

phys. Research, v. 65, no. 2, p. 695-719, 9 figs. incl. graph, profiles, table, Feb. 1960, 18 refs.

In Pt. 1 a single-parameter visco-elastic model of a shear-yielding solid is defined for which a permanent-regime solution of the equations of motion exists for any finite compression. The 'profile' of this compression as a function of distance is obtained in the form of an integral which can be evaluated when the velocity of propagation is known as a function of final compression. It is assumed that the permanent-regime solution approximates actual shock waves and that the velocity of the permanent-regime profile equals the shock velocity. Observed shock speeds are used to compute shock profiles in a number of metals. The maximum slope of the profile for any one metal increases with increasing compression. The limiting value of the maximum slope as the volume is extrapolated to zero gives a numerical estimate of the viscosity parameter, and this has been done for Al, Pb, Sn, Zn, and Zr.

In Pt. 2 Zener's linear theory of anelasticity has been generalized to materials with cubic crystal structure. The theory of the propagation and attenuation of plane waves, both longitudinal and transverse, along a principal axis of the crystal is presented. The combined effects of relaxation mechanisms and thermal diffusion are included. The significance of the results for the theory of shock propagation are discussed, and several questions are raised for later discussion.

In Pt. 3 the general equations for propagation of steady-state compression profiles in shear-yielding, heat-conducting, anelastic solids are given. Methods of solution by successive approximation are developed for Hookean solids, with both adiabatic and isothermal (very steep) profiles, and for non-Hookean solids with shock profiles. The results of Pt. 1 are corrected to include the effects of thermal conductivity and anelasticity. Heating aftereffects of shocks are discussed, including the effects of irreversible heating due to viscous yielding, etc., and it is shown how the temperature of the solid after passage of a shock profile may be calculated.--Auth.

2-1500. Birch, Francis. THE VELOCITY OF COMPRESSIONAL WAVES IN ROCKS TO 10 KILOBARS, PART 1: Jour. Geophys. Research, v. 65, no. 4, p. 1083-1102, 5 figs., 6 tables, Apr. 1960, 52 refs.

The velocity of compressional waves has been determined by measurement of travel time of pulses in specimens of rock at pressures to 10 kilobars and room temperature. Most of the samples, mainly igneous and metamorphic rocks, furnished 3 specimens oriented at right angles to one another. The present paper gives experimental details, modal analyses, and numerical tables of velocity as function of direction of propagation, initial density, and pressure. Discussion of various aspects of the measurements is reserved for Pt. 2.--Auth.

2-1501. Rosenbaum, J.H. THE LONG-TIME RESPONSE OF A LAYERED ELASTIC MEDIUM TO EXPLOSIVE SOUND: Jour. Geophys. Research, v. 65, no. 5, p. 1577-1613, 31 figs., table, May 1960, 9 refs.

The long-time response of a layered elastic medium is considered for the particular case of a point-source explosion in a liquid layer lying above an infinitely deep liquid bottom. An asymptotic solution,



id for large values of the time variable, is ob-  
hed; it expresses the response in terms of har-  
nic vibrations of the liquid layer. Special empha-  
is placed on those vibrations which correspond  
waves with small angles of incidence and which,  
cause of radiation into the bottom, decay ex-  
ponentially with time. However, the well-known  
ded-wave phenomenon, first discussed by Pekeris  
1948, is also included in the present formulation.  
A detailed presentation is made of the method of  
alysis. This method is applicable to more com-  
licated problems of direct geophysical interest.  
merical results are presented for some typical  
amples, the behavior of the phase velocity for the  
ver modes being of particular interest.--Auth.

1502. Lovan, T. E. SONIC LOGGING - A METH-  
OF POROSITY DETERMINATION (In: McGrain,  
eston, and Thomas J. Crawford, eds. Proceed-  
s of the Technical Session, Kentucky Oil and Gas  
society, Twenty-Third Annual Meeting, June 5,  
59: Kentucky Geol. Survey, Ser. 10, Spec. Pub.  
p. 51-60, 2 diags., 4 logs, table, 1959) 2 refs.

The Sonic log as a method of determining reser-  
ir porosity was introduced in the Tri-State area  
Jan. 1959. It has advantages over other logging  
ethods in that it can be used in rocks of varying  
rosities and for carbonate reservoir rocks as well  
sandstones. Some examples of this well-logging  
ethod are given.--Auth.

1503. Savit, Carl H. USE SEISMIC DATA TO  
ND STRATIGRAPHIC TRAPS: Oil & Gas Jour. v.  
no. 15, p. 182-184, port., 2 logs, Apr. 11,  
60, 10 refs.

Stratigraphy has been only incidentally determi-  
ble by the seismograph. Stratigraphic information  
s to be found by distinguishing the different kinds  
reflections from the different kinds of rocks.  
is implies getting rid of noise without filters and  
ontrolling amplitude while retaining contrasts.  
chniques are available but a great deal of experi-  
ent and evaluation must be done. A short resume  
its content accompanies each reference.--N.  
reet.

1504. MacFarlane, R. M., and R. K. Ault.  
CLEAR LOGGING IN THE APPALACHIAN BASIN  
: McGrain, Preston, and Thomas J. Crawford,  
5. Proceedings of the Technical Session, Kentucky  
i and Gas Association, Twenty-Third Annual  
eting, June 5, 1959: Kentucky Geol. Survey, Ser.  
i, Spec. Pub. 2, p. 13-27, 4 diags., 7 logs,  
59).

Gamma ray-neutron logs are being used for  
rosity determination under various bore hole con-  
itions. The gamma ray log is used to correct the  
utron response so that true porosity (effective)  
n be determined.

The combination gamma ray-dual spaced neutron  
s have been successful in detecting gas-bearing  
ormations in wet or dry cased and uncased bore  
es. The dual spaced neutron log is also used in  
ferentiating between liquid filled porosity and gas  
ed porosity. This combination of logs does not  
icate formations as gas bearing in the case where  
asion of bore hole liquid is greater than the depth  
investigation of the neutron device.--Auth.

1505. Ratcliffe, E. H. THE THERMAL CON-  
CTIVITIES OF OCEAN SEDIMENTS: Jour. Geo-

phys. Research, v. 65, no. 5, p. 1535-1541, 6 figs.,  
incl. illus., diags., graph, table, May 1960, 12 refs.

As a part of the measurement of heat flow through  
the ocean floor, the thermal conductivities of sam-  
ples of sediment from the Pacific, Atlantic, and  
Mediterranean areas have been determined in the  
laboratory by a steady-state method. Average values  
over a wide range of water content are found to de-  
pend more on water content than on solid phase  
constituents, and conductivities can be read from a  
diagram if the amount of contained sea water or the  
wet density of the sediment is known. A graph is  
included, showing the conductivities of wet, granular  
materials other than ocean sediment.--Auth.

2-1506. Pakiser, Louis C., Frank Press, and  
Martin F. Kane. GEOPHYSICAL INVESTIGATION  
OF MONO BASIN, CALIFORNIA: Geol. Soc. Amer-  
ica, Bull., v. 71, no. 4, p. 415-447, 2 maps (1  
fold.), sec., 14 profiles, 3 seismograms, 2 tables,  
Apr. 1960, 27 refs.

Gravity and seismic studies in Mono basin, Mono  
County, completed during the summer of 1957,  
revealed a large, roughly triangular block that had  
subsided about 18,000 + 5,000 ft. and received an  
accumulation of about 300 + 100 cu. mi. of light  
clastic sediments and volcanic material of Cenozoic  
age. The seemingly near-vertical faults that bound  
this great block are displaced toward the center of  
the basin from the surrounding mountain masses,  
but in general they are parallel to well-defined  
Basin and Range trends.

The gravity minimum anomaly associated with the  
Mono basin structure has a residual gravity relief  
of about 50 mgals. and the lowest gravity readings  
(on Paoha Island) yield a complete Bouguer gravity  
value of about - 260 mgals. with respect to the Inter-  
national Ellipsoid. The computed depth of subsidence  
is based on a density of 2.3 gms/cm<sup>3</sup> for the basin  
fill and 2.7 gms/cm<sup>3</sup> for the basement rocks. Seis-  
mic-refraction profiles at several places in the basin  
demonstrate that the Cenozoic deposits are thick  
where the gravity is low and relatively thin where  
the gravity is higher. Along common seismic and  
gravity profiles steep seismic dips coincide with  
steep gravity gradients. Numerous seismic re-  
flections are present within the basin fill. Anomalies  
on 4 aeromagnetic profiles are related in part to  
volcanic material within the Cenozoic section.

It is concluded that Mono basin may be a volcano-  
tectonic depression caused by subsidence along  
faults, following extrusion of magma from a magma  
chamber at depth. Volcanic rocks of Pliocene(?) and  
Pleistocene ages are exceptionally abundant in this  
area.--Auth.

2-1507. Press, Frank. CRUSTAL STRUCTURE  
IN THE CALIFORNIA-NEVADA REGION: Jour. Geo-  
phys. Research, v. 65, no. 3, p. 1039-1051, 9 figs.  
incl. map, graphs, 5 tables, March 1960, 15 refs.

A case is made for the combined use of 3 methods  
for the study of the earth's crust in a given region:  
seismic refractions, surface wave phase velocity,  
and gravity. Only by this approach can the fine de-  
tails of crustal structure be revealed. The standard  
phase-velocity curves are revised to take into ac-  
count recent refraction results in South Africa and  
the Gutenberg low-velocity zone of the upper mantle.  
Restrictions on the use of the phase-velocity method  
alone are discussed.

In the California-Nevada region seismic refrac-  
tions reveal the following structure below the sedi-

ments: 23 km. of granitic rock with velocities of 6.11 km./sec. and 3.49 km./sec. for compressional and shear waves; 26 km. of gabbroic-ultramafic rock with a compressional velocity of 7.66 km./sec. underlain by a zone of ultramafic rock with a compressional velocity of 8.11 km./sec. When this structure is used in computing theoretical Rayleigh

wave phase velocities and gravity anomaly, discrepancies are found with the observed values which can be resolved by reducing the mean shear velocity and the density in the crust. It is probable that this reduction is limited to the intermediate crustal layer, and several modifications consistent with all 3 exploration methods are discussed.--Auth.

## 7. GEOCHEMISTRY

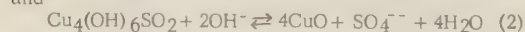
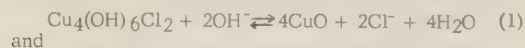
See also: Areal and Regional Geology 2-1353; Igneous and Metamorphic Petrology 2-1545, 2-1546, 2-1555; Mineral Deposits 2-1581; Fuels 2-1598.

2-1508. Roy, Rustum. HIGH PRESSURE - A NEW CHEMICAL TOOL: Mineral Industries, v. 29, no. 5, p. 1, 4-6, 8, 4 illus., 9 diags., Feb. 1960.

Until recently, pressure as a variable has been neglected in most physical-chemical studies relevant to geochemistry. Lately, however, it has become possible to study chemical reactions at pressures up to 1.5 million p.s.i. This paper describes briefly the different devices used for geochemical studies at high pressures, with particular emphasis on the apparatus currently in use at Pennsylvania State University, namely: externally and internally heated cold seal test tube bombs and externally heated uniaxial pressure devices. Included in this paper are some results obtained from the use of these pressure devices in such areas as: 1) synthesis of minerals and new materials, 2) melting of silicate material under volatile pressure, 3) stability of mineral phases and assemblages and their importance in petrology, and 4) application of results of high-pressure studies to geophysical problems.--D. Hawkins.

2-1509. Barton, Paul B., Jr., and Philip M. Bethke. THERMODYNAMIC PROPERTIES OF SOME SYNTHETIC ZINC AND COPPER MINERALS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 21-34, 7 diags., table, 1960, 10 refs.

Measurement of pH of a solution of known  $Zn^{++}$  or  $Cu^{++}$  concentration in equilibrium with the metal hydroxide or oxide allows the calculation of an apparent equilibrium constant ( $K' = [M^{++}] \cdot a^{2}_{OH^{-}}$ ) which varies with composition of the solution. Extrapolation of  $K'$  values to infinite dilution permits estimation of a true equilibrium constant ( $K = a^{2}_{M^{++}} \cdot a^{2}_{OH^{-}}$ ). This procedure is applied to the determination of the equilibrium constant for  $ZnO$  (zincite).  $CuCl_2$  and  $CuSO_4$  solutions precipitate the basic salts  $Cu_4(OH)_6Cl_2$  (atacamite), and  $Cu_4(OH)_6SO_4$  (brochantite) at concentrations greater than about  $10^{-4.5}$  molar. Extrapolation to infinite dilution gives the equilibrium constants for these basic salts. At higher pH values reactions such as



take place and permit the estimation of  $K$  for  $CuO$  (tenorite). For zincite,  $K = 10^{-17.0 \pm 0.1}$ ; for brochantite,  $K = 10^{-68.6 \pm 0.4}$ ; for atacamite,  $K = 10^{-69.4 \pm 0.4}$ ; and for tenorite,  $K = 10^{-19.7 \pm 0.4}$ . The variation of  $K$  with temperature ( $75^{\circ}$  to  $750^{\circ}C.$ ) permits the calculation of heat content for tenorite, brochantite, and atacamite, but for tenorite the more reliable heat of formation is obtained from the free energy data and the entropy measured by Hu and Johnston.

To the extent that the pH measurements represent the H ion activity the ratio  $\frac{[Cl^{-}]}{a^{2}_{OH^{-}}}$  gives the activity coefficient of the metal ion at any particular ionic strength. The activity coefficients of  $Cl^{-}$  and  $SO_4^{--}$  are derived from  $\frac{[Cl^{-}]}{a^{2}_{OH^{-}}} = K_1$  and  $\frac{[SO_4^{--}]}{a^{2}_{OH^{-}}} = K_2$

obtained by pH measurement from reactions (1) and (2).

The occurrence of tenorite in an ore body indicates that the concentration of  $Cu^{++}$  in the solutions which are in equilibrium with the oxidized ore is less than about  $10^{-4.5}$  molar.--Auth.

2-1510. Boyd, Francis R., and J. L. England. APPARATUS FOR PHASE-EQUILIBRIUM MEASUREMENTS AT PRESSURES UP TO 50 KILOBARS AND TEMPERATURES UP TO  $1750^{\circ}C.$ : Jour. Geophys. Research, v. 65, no. 2, p. 741-748, 6 figs. incl. diags., graphs, 2 tables, Feb. 1960, 7 refs.

Construction and calibration of apparatus utilizing a solid pressure medium for phase-equilibrium studies at elevated temperatures and pressures are described. Pressure calibration is carried out by measurement of the Bi I-Bi II and Tl II-Tl III transitions. A new determination of the Tl transition,  $37.1 \pm 1.3$  kilobars, is given. Tests indicate that talc is superior to pyrophyllite and boron nitride as a solid pressure medium for high-temperature work.--Auth.

2-1511. Korzhinsky, D.S. THE NUMBER OF FACTORS OF STATE IN SYSTEMS (A REPLY TO I.V. ALEXANDROV): Geokhimiya [in translation], 1958, no. 5, p. 641-645, pub. 1959, 7 refs.

The main area of disagreement between Alexandrov and Korzhinsky is the function and importance of intensive and extensive parameters of thermodynamic systems, where intensive parameters refer to such variables as temperature, pressure, and concentration, and extensive parameters are those dependent on mass, such as volume, entropy, and masses of components.

Korzhinsky states that the Gibbs phase rule, which considers only intensive parameters, finds frequent use in his previous publications, especially those dealing with analyses of mineral parageneses. Under certain conditions, however, the Gibbs formulation is inadequate and must be supplemented by Korzhinsky's equation,  $f_{in} + f_{ex} = k + 2$ , where  $f_{in}$  and  $f_{ex}$  are the intensive and extensive factors of state, and  $k$  is the number of components in the system. An example of such a condition would be a closed system in which the masses of the "k" components are fixed.--F. Manheim.

2-1512. Schairer, J. F., and Hatten S. Yoder, Jr. THE NATURE OF RESIDUAL LIQUIDS FROM CRYSTALLIZATION, WITH DATA ON THE SYSTEM NEPHELINE-DIOPSIDE-SILICA: Am. Jour. Sci.,



258-A (Bradley Volume), p. 273-283, 4 diagrs., 1960, 22 refs.

The phase-equilibrium relations in the system nepheline-diopside-silica have been determined. The albite-diopside, although it is not binary, divides the system into 2 parts, a silica-rich portion with albite as one of the solid phases in completely crystallized mixtures and a nepheline-rich portion with nepheline as one of the solid phases in the final crystallization product. In the system nepheline-diopside-silica the primary phases are diopside (probably an aluminous diopside solid solution), common nepheline solid solutions, carnegieite solid solutions, sodic plagioclases, forsterite, tridymite, and cristobalite. The join jadeite-diopside is not unique in this system at 1 atm. pressure, and jadeite is not encountered as a solid phase.

The results support Bowen's contention that through fractionation a magma is impoverished in early-crystallizing minerals (in this case diopside) and yields residual liquids rich in the alkali feldspar-silicates.

The nature of the albite-diopside join indicates that no single anhydrous composition can yield a liquid which through fractionation produces both a silica-rich residue and a nepheline-rich residue. There is a reaction relation between olivine and diopside. This reaction relation should be as important in the alkali basalts as the olivine-hypersthene reaction relation is in the tholeiitic basalts. -- Auth.

513. Skinner, Brian J., and Howard T. Evans. CRYSTAL CHEMISTRY OF  $\beta$ -SPODUMENE SOLID SOLUTIONS ON THE JOIN  $\text{Li}_2\text{O} \cdot \text{Al}_2\text{O}_3 \cdot \text{SiO}_2$ . Am. Jour. Sci., v. 258-A (Bradley Volume), p. 312-322, illus., 3 diagrs., 2 tables, 1960, 12 refs.

$\beta$ -spodumene ( $\text{LiAlSi}_2\text{O}_6$ ) is tetragonal, space group  $P4_32_1$ ,  $a = 7.5332 \pm 0.0008 \text{ \AA}$  and  $c = 9.1540 \pm 0.0008 \text{ \AA}$ . It is isostructural with keatite ( $\text{SiO}_2$ ) and forms an extensive solid solution towards  $\text{SiO}_2$  and  $\text{AlSiO}_4$ .  $\beta$ -spodumenes of various compositions have been synthesized and their unit cell edges determined.

On the assumption that  $\beta$ -spodumene is a "stuffed silica structure" based on the known keatite structure, probable locations of the introduced Li atoms are established from geometrical considerations. From the change of unit cell volume with composition, it is deduced that, up to the spodumene composition, no atoms are inserted in cavities on the twofold axes; beyond this composition, additional cations are inserted in cavities in general positions in the structure. Channels extend throughout the structure connecting these cavities, providing paths along which atoms can move under potential gradients. These channels account for the cation-exchange properties of  $\beta$ -spodumene.

Data are presented on the composition of co-existing  $\beta$ -spodumene solid solution and  $\beta$ -eucryptite solid solution at  $1350^\circ\text{C}$ . and 1 atm. and on the composition of  $\beta$ -spodumene solid solution coexisting with tridymite at  $1350^\circ\text{C}$ . and 1 atm. -- Auth.

514. Pistorius, Carl W.F.T., and George C. Kennedy. STABILITY RELATIONS OF GROSSULARITE AND HYDROGROSSULARITE AT HIGH TEMPERATURES AND PRESSURES: Am. Jour. Sci., v. 258, no. 4, p. 247-257, 2 diagrs., 2 tables, Apr. 1960, 12 refs.

The equilibrium relations between hydrogrossu-

larite + quartz and anorthite + wollastonite + water have been determined to 15,500 bars and  $810^\circ\text{C}$ . by means of the "simple squeezer" high-pressure apparatus. There is a continuous solid-solution series from grossularite,  $3\text{CaO} \cdot \text{Al}_2\text{O}_3 \cdot 3\text{SiO}_2$ , through hydrogrossularite to hibschite,  $3\text{CaO} \cdot \text{Al}_2\text{O}_3 \cdot 2\text{SiO}_2 \cdot 2\text{H}_2\text{O}$ . The composition of hydrogrossularite, above  $500^\circ\text{C}$ . and in equilibrium with quartz, is determined almost entirely by temperature. Grossularite is stable above  $780^\circ$  and 15 kilobars  $\text{H}_2\text{O}$  pressure in the presence of quartz. -- Auth.

2-1515. Yoder, Hatten S., Jr., and Charles E. Weir. HIGH-PRESSURE FORM OF ANALCITE AND FREE ENERGY CHANGE WITH PRESSURE OF ANALCITE REACTIONS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 420-433, 3 diagrs., 3 tables, 1960, 38 refs.

The compressibility of 4 natural analcites has been obtained up to 10,000 atm. at  $25^\circ\text{C}$ . All specimens show abnormal compressibility [ $(\partial\beta/\partial P)_T > 0$ ] and one exhibits a reversible transition at about 8,400 atm. Abnormal increase of compressibility with pressure usually precedes a transition, and it is suggested that a similar transition lies at higher pressures than those investigated for the 3 remaining specimens.

The compressibility and density data were used with previously obtained data to compute the volume change and free energy change with pressure for the reactions  $2 \text{ analcite} \rightleftharpoons \text{ nepheline} + \text{ albite} + 2 \text{ water}$ ,  $\text{ analcite} \rightleftharpoons \text{ jadeite} + \text{ water}$ , and  $\text{ analcite} + \text{ quartz} \rightleftharpoons \text{ albite} + \text{ water}$ . The volume change for these reactions proceeding to the right is negative for all pressures investigated at  $25^\circ\text{C}$ .

The analcite exhibiting the transition showed an increase in birefringence after compression, which is attributed either to strain induced by the volume discontinuity or to retention of a possible lower symmetry of the high-pressure form. Lower symmetry in a high-pressure form is not uncommon in previously investigated substances. -- Auth.

2-1516. Dacheville, Frank, and Rustum Roy. HIGH PRESSURE STUDIES OF THE SYSTEM  $\text{Mg}_2\text{GeO}_4$ - $\text{Mg}_2\text{SiO}_4$  WITH SPECIAL REFERENCE TO THE OLIVINE-SPINEL TRANSITION: Am. Jour. Sci., v. 258, no. 4, p. 225-246, 14 diagrs., 4 tables, Apr. 1960, 31 refs.

The system  $\text{Mg}_2\text{GeO}_4$ - $\text{Mg}_2\text{SiO}_4$  has been studied to the experimental limits of available hydrothermal and uniaxial high pressure apparatuses.

The inversion temperature for the  $\text{Mg}_2\text{GeO}_4(\text{spinel})$ - $\text{Mg}_2\text{GeO}_4(\text{olivine})$  equilibrium is  $810^\circ\text{C}$ . at atmospheric pressure. The  $\Delta V$  of the inversion is 3.5 c.c./mole;  $\Delta H$  is  $3,690 \pm 180 \text{ cal./mole}$ . The inversion temperature is raised by  $0.025^\circ\text{C./bar}$  for the first 5,500 bars. Infrared absorption spectra, X-ray intensities, and molar refractivities clearly show that  $\text{Mg}_2\text{GeO}_4$  is an inverse spinel.

Solid solution between  $\text{Mg}_2\text{GeO}_4$  and  $\text{Mg}_2\text{SiO}_4$  is complete in the olivine phase at temperatures above that of the inversion in  $\text{Mg}_2\text{GeO}_4$ . The maximum silicate content of the spinel solid solutions at lower temperatures increases steadily with pressure, from 10 mole % at 700 bars to 50 mole % at 60,000 bars at  $542^\circ\text{C}$ . Extrapolation places the spinel-olivine transition for  $\text{Mg}_2\text{SiO}_4$  at  $100,000 \pm 15,000$  bars. The change in the lattice spacings of the spinel solid solutions of  $\text{Mg}_2(\text{Ge,Si})_4\text{O}_4$  shows that  $\text{Mg}_2\text{SiO}_4(\text{spinel})$

has a cell edge of 8.22Å. Therefore the  $\Delta V$  for this transition is 2.0 c.c./mole. The pressure dependence of the transition in  $Mg_2SiO_4$  is estimated by extrapolation at 0.013°C./bar.

Experiments show that substitution of  $Fe^{2+}$  for  $Mg^{2+}$  markedly increases the maximum silicate content of the spinel solid solution under corresponding p-t conditions.

The geophysical implication of these results is that an olivine-spinel transition in the mantle of the earth does appear to be a reasonable explanation of the seismic and density discontinuities starting at 400 km.--Auth

2-1517. Verma, R.K. ELASTICITY OF SOME HIGH-DENSITY CRYSTALS: Jour. Geophys. Research, v. 65, no. 2, p. 757-766, illus., diag., 10 tables, Feb. 1960, 22 refs.

The adiabatic elastic constants of 2 garnets (spessartite-almandite and almandite), spinel (synthetic), rutile (synthetic), and olivine are reported. The stiffness constants  $C_{pq}$  were determined from the velocities of acoustic wave propagation in crystals. The velocities of wave propagation were measured by McSkimin's method. A frequency range of 6 to 12 Mc/s was used.--Auth.

2-1518. Boyd, Francis R., and J.L. England. THE QUARTZ-COESITE TRANSITION: Jour. Geophys. Research, v. 65, no. 2, p. 749-756, illus., 2 graphs, 4 tables, Feb. 1960, 13 refs.

The quartz-coesite transition curve has been determined over the temperature range 700°C to 1700°C. in the pressure range 20 to 40 kb. The equation for the curve is  $P = 19.5 + 0.0112T$ , where P is in kilobars and T is in degrees centigrade. The determination was made with an internally heated tungsten carbide pressure vessel utilizing talc for a pressure medium. The results of some runs on the quartz-coesite transition with an anvil apparatus are described. These data are in poor agreement with the determination cited above, and the difference is interpreted as an effect of pressure gradients in quartz-coesite runs in the anvil apparatus. Quartz could not invert to coesite in the earth at depths less than about 100 km., and it is therefore unlikely that coesite has formed in crustal rocks.--Auth.

2-1519. Gast, Paul W. LIMITATIONS ON THE COMPOSITION OF THE UPPER MANTLE: Jour. Geophys. Research, v. 65, no. 4, p. 1287-1297, 5 tables, Apr. 1960, 50 refs.

New determinations of the isotopic composition of Sr and of the concentration of K, Rb, Cs, Sr, and Ba in rocks and meteorites are given. The isotopic abundance of  $Sr^{87}$  in the upper mantle and the crust appears to be lower than that found for chondrites. Furthermore, for a chondritic earth model, the concentrations of K, Rb, and Cs in the earth's crust are anomalous when compared with those of U, Ba, and Sr. These 2 concurring arguments indicate that the upper mantle and crust of the earth do not contain K, Rb, Cs, U, Ba, and Sr in the proportions found in chondrites and that the alkali metals are depleted relative to U, Sr, and Ba. This depletion may be an indication of a nonchondritic earth composition; it may also result from an earth differentiation in which K, Rb, and Cs were concentrated or retained in the lower mantle.--Auth.

2-1520. Green, Jack. SOME ASPECTS OF EXTRA-TERRESTRIAL GEOCHEMISTRY: Geochim. Soc., Geochim. News, no. 21, p. 3-8, 7 illus., diag., 3 tables, Apr. 1960, 6 refs.

Defluidization of the mantle of the moon and associated volcanism may have produced the following results: 1) Extensive production of silicic extrusives including welded tuffs in non-mare areas and subsequent evolution of more mafic extrusives as mare. 2) Formation of calderas in and along zones of collapse and at fracture intersections. 3) Formation of relatively thick volcanic dust in non-mare areas, and only thin spallation dust in mare areas. However, local accumulation of relatively thick dust in mare areas is probable in view of post-mare development of craters in maria. 4) Appreciable concentration of compounds in eternally shadowed zones particularly of Cl, Br, S, B, I, and F, and to a lesser extent, Cs, As, Hg, Sb, Ar, and Pb in non-mare areas relative to the average terrestrial sedimentary rock. Most of these elements have high neutron capture cross-sections. 5) Concentration of these compounds as minerals formed in environments low or lacking in O and under partial vacuum. These minerals might correspond to the closed-tube sublimates obtained in determinative mineralogical techniques.--From auth. introd.

2-1521. Hoffman, J.H., and Alfred O. Nier. COSMIC-RAY-PRODUCED HELIUM IN THE KEEN MOUNTAIN AND CASAS GRANDES METEORITES: Jour. Geophys. Research, v. 65, no. 3, p. 1063-1068, 5 figs., 3 tables, March 1960, 10 refs.

The  $He^3$  and  $He^4$  distributions have been measured in the Fe meteorites Keen Mountain [Virginia] and Casas Grandes [Mexico]. In the former, a small meteorite (6.75 kg.), the  $He^3$  and  $He^4$  concentration did not depend upon position. In the latter, a large meteorite (1,550 kg.), a 'depth effect' was observed and contours of constant  $He^3$  and  $He^4$  content could be drawn. An attempt is made to explain the result in terms of the model earlier presented in connection with similar work on the Grant meteorite.--Auth.

2-1522. Larsen, Esper S., 3d, and David Gottfried. URANIUM AND THORIUM IN SELECTED SUITES OF IGNEOUS ROCKS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 151-169, 16 diags., 3 tables, 1960, 22 refs.

The U and Th contents of 199 igneous rocks from a variety of petrographic provinces is here summarized. Data are given for the Mesozoic calc-alkalic batholiths of the western United States; volcanic and hypabyssal rocks of the tholeiitic magma type from Hawaii and Virginia; and effusive calc-alkalic, alkalic, and subsilicic-alkalic rocks from the western United States and Hawaii.

The batholithic rocks show an increase of both U and Th from gabbro to quartz-monzonite and granite. The more extreme differentiates, chiefly muscovite quartz monzonites, contain considerably less U and Th than the quartz monzonites and granites, although the Th/U ratios are nearly the same. The volcanic and hypabyssal rocks in general show a similar increase in both Th and U toward the more felsic members. The alkalic basalts of the Honolulu volcanic series show an anomalous decrease in U and Th toward the right on a variation diagram.

The Th/U ratios remain fairly constant from the mafic to the felsic members of each series here studied; each series also has a more or less charac-



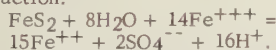
istic Th/U ratio, ranging from 2 1/2 to 5. The scatter of Th contents, U contents, and Th/U ratios larger with increasing complexity of magmatic differentiation. The batholithic rocks have the greatest scatter and have a significantly lower Th/U ratio in the gabbros than in the more felsic members. Our data show no increasing loss of U relative to Th from the magma during the later stages of crystallization.--Auth.

1523. Skougstad, Marvin W., and C. Albert. OCCURRENCE OF STRONTIUM IN NATURAL WATERS: U.S. Geol. Survey, Circ. 420, 6 p., map, tables, 1960, 8 refs.

The regions where the stable Sr content of surface waters is relatively low (less than 0.50 p.p.m.) include the Pacific Northwest, northeastern United States, and the central lowlands, particularly the lower Mississippi basin and the western Gulf Coast area. Moderate concentrations of Sr (0.50 to 1.5 p.p.m.) are found in streams of southeastern United States, most of the Great Plains region, the western mountain and plateau regions, and California. Relatively high concentrations of Sr occur in the surface waters of an area that includes northern and western Texas and southern New Mexico and Arizona. Exceptions to the above distribution are due to local geologic conditions.--Auth.

1524. Garrels, Robert M., and M.E. Thompson. OXIDATION OF PYRITE BY IRON SULFATE SOLUTIONS: Am. Jour. Sci., v. 258-A (Bradley Volume), 1960, 57-67, 9 diags., 2 tables, 1960, 2 refs.

The rate of oxidation of pyrite specimens from 3 localities was measured in acid iron sulfate solutions. The rate of reduction of ferric ion in these solutions is considered to be a measure of the rate of oxidation of the pyrite. The over-all oxidation reaction, in which 100% to 0.1%  $\text{Fe}^{+++}$  is in accord with the classical reaction:



The average rate of reduction from 100%  $\text{mFe}^{+++}$  to 50%  $\text{mFe}^{+++}$  is constant for pyrite from a given locality, and differences in the average rates between specimens from different localities could not be related to minor compositional variations of the solutions or of the pyrites. The pH did not affect rate in the tested range, pH 0 to pH 2.

The instantaneous rate of reduction of ferric ion diminishes with decrease in the ferric-ferrous ratio; it is postulated that instantaneous rate is controlled by differential absorption of ferric and ferrous ions on the pyrite surface, and is proportional to the fraction of pyrite surface occupied by ferric ion.--Auth.

1525. Broecker, Wallace S. EXTENSION OF TECHNIQUE FOR NATURAL TRITIUM MEASUREMENTS: Columbia Univ., Lamont Geol. Observatory, Final Rept., Contract no. AF19(604)-4076, 25 p., illus., 8 tables, Dec. 1959.

An attempt has been made to define the source

and time variations in the natural background of large-volume proportional counters in order that more sensitive counters might be constructed for the tritium assay of natural waters. An 8-liter counter was constructed and tested with a number of different shielding arrangements, with different cathode materials, and with various filling gases. A description is given of the counter, including the main counter shell, 5 types of counter inserts tried, and counter shielding. A background of 22 c.p.m. is the minimum value obtained so far for methane at 1 atm. filling pressure. Hg and full water shielding are required to secure this low value; with increasing pressure, the background rises about 5 c.p.m. per atm. The dominant components of residual background are probably neutrons, gamma rays, and beta particles; alpha particles and mesons are of minor significance. Beta rays are as yet a question mark although they may be the most important contribution of all. Results obtained indicate the need for specific experiments which should permit a full analysis of background together with techniques for keeping it to a minimum.--M. Russell.

2-1526. Oana, Shinya, and Edward S. Deevey. CARBON 13 IN LAKE WATERS, AND ITS POSSIBLE BEARING ON PALEOLIMNOLOGY: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 253-272, diag., 7 tables, 1960, 38 refs.

Surface waters of dimictic lakes in Connecticut have an average  $\delta\text{C}^{13}$  of -8.6‰, but the hypolimnion is more depleted ( $\delta\text{C}^{13} > -22.6\%$ ). Seston, mud, and mollusc flesh average ca. -30‰, but submerged pondweeds are richer. Four sources, partly evaluable from  $\text{C}^{13}$  ratios, can contribute to hypolimnetic  $\text{CO}_2$ : a) surface  $\text{CO}_2$  ( $\delta\text{C}^{13} \sim -9\%$ ), inherited from circulation; b) respiratory  $\text{CO}_2$  ( $\delta\text{C}^{13} \sim -30\%$ ) from oxidation of seston; c) fermentation  $\text{CO}_2$  ( $\delta\text{C}^{13} \sim -5\%$ ), here shown to be enriched at the expense of  $\text{CH}_4$  ( $\delta\text{C}^{13} \sim -70\%$ ); d) carbonate  $\text{CO}_2$  ( $\delta\text{C}^{13} \sim 0\%$ ), important in hard waters. Metabolism = b + c; and d, both heavy, can be separated by  $\text{C}^{14}$  assay, but this work is incomplete. In Quassapaug (softest), where d is negligible, c represents 4 to 20% of total metabolism. In Queechy (hardest), d could account for all heavy  $\text{CO}_2$ . In Linsley Pond (medium-hard), after allowance for d, c appears to provide 40% of total metabolism. Discovery that fermentation  $\text{CO}_2$  is heavy raises important questions about groundwater. Applied to paleolimnology,  $\text{C}^{13}$  assay may be ambiguous for the same reason: fermentation tends to raise ratios lowered by aerobic metabolism. In one meromictic lake in a humid region, permanent stagnation has not produced fractionation, presumably because of c; owing to dominance of d, a meromictic saline lake also fails to show fractionation.

The earliest organic deposits of the last Pleistocene Searles Lake are enriched in  $\text{C}^{13}$ , presumably because the lake redissolved much carbonate, and the  $\text{C}^{13}$  ratio fell only gradually. In the middle Huronian Fe-banded cherts of Michigan, the  $\text{FeCO}_3$ -rich layers may be poorer in  $\text{C}^{13}$  than the  $\text{SiO}_2$ -rich layers, and the depletion of these siderites supports the lacustrine hypothesis.--Auth.

## 8. MINERALOGY AND CRYSTALLOGRAPHY

See also: Igneous and Metamorphic Petrology 2-1552, 2-1553; Sedimentary Petrology 2-1563; Mineral Deposits 2-1586.

2-1527. Silverman, E.N., and Thomas F. Bates. X-RAY DIFFRACTION STUDY OF ORIENTATION IN THE CHATTANOOGA SHALE: *Am. Mineralogist*, v. 45, no. 1/2, p. 60-68, 3 figs. incl. illus., diag., 2 tables, Jan.-Feb. 1960, 4 refs.

An X-ray diffraction technique has been used to obtain a quantitative measure of the degree of orientation of the (001) planes of 10 Å illite and mica flakes in 58 samples from a drill core taken from the Chattanooga shale. Thin sections cut perpendicular to the bedding are placed on the X-ray spectrometer in such a position that, upon rotation of the section in its own plane, the geiger counter records all diffraction from mica flakes with (001) perpendicular to the plane of the section. Analysis of the resulting curve, using the Chi-square statistic divided by the area under the curve, gives an "orientation index" for each specimen which is reproducible and independent of the amount of 10 Å material present.

An evaluation of the data with respect to the U content of the samples shows that there is no significant correlation between the concentration of the element and this particular measure of the texture of the rock. This supports other evidence which indicates that the U was precipitated during deposition of the sediment and that the distribution of U has not been affected by movement of solutions during or following compaction.

The technique is applicable to the study of lattice orientation of any sufficiently fine-grained mineral in natural or synthetic substances.--Auth.

2-1528. Clark, Joan R. X-RAY STUDY OF ALTERATION IN THE URANIUM MINERAL WYARTITE: *Am. Mineralogist*, v. 45, no. 1/2, p. 200-208, 2 illus., 4 tables, Jan.-Feb. 1960, 5 refs.

X-ray examinations of the mineral wyartite, assigned the chemical formula  $\text{UO}_2 \cdot 6\text{UO}_3 \cdot 2\text{CO}_2 \cdot 3\text{CaO} \cdot 12\text{-}14\text{H}_2\text{O}$  by Guillemin and Protas, show that with passage of time the mineral is altering. Up to the present time 2 phases, both orthorhombic, have been found from repeated precession X-ray examination of a single crystal: wyartite I (comparable to "ianthinite" described by Bignand),  $a = 11.25 \pm 0.03$ ,  $b = 7.098 \pm 0.03$ ,  $c = 20.80 \pm 0.06$  Å, space group  $\text{P}2_12_12_1\text{-D}^{42}_2$  (no. 19), pseudo space-groups  $\text{P}2_1\text{cn-C}^{92}_{2v}$  (no. 33) or  $\text{Pmcn-D}^{16}_{2h}$  (no. 62); wyartite II (not previously described),  $a = 11.25 \pm 0.03$ ,  $b = 7.098 \pm 0.03$ ,  $c = 16.83 \pm 0.05$  Å, space groups  $\text{P}2_1\text{cn-C}^{92}_{2v}$  (no. 33) or  $\text{Pmcn-D}^{16}_{2h}$  (no. 62), pseudo space-group  $\text{Pbcn-D}^{14}_{2h}$  (no. 60). Indexed X-ray powder diffraction data for both phases are given. The alteration occurs without visible external change in the crystals, so that interpretation of chemical analyses on macroscopic samples may be open to question. The precise nature of the alteration is still unknown, although oxidation of  $\text{U}^{+4}$  to  $\text{U}^{+6}$ , with formation of uranyl ions, possibly also accompanied by some dehydration, seems to be a plausible explanation.--Auth.

2-1529. McConnell, Duncan. THE CRYSTAL CHEMISTRY OF DAHLITE: *Am. Mineralogist*, v. 45, no. 1/2, p. 209-216, illus., 3 tables, Jan.-Feb. 1960, 13 refs.

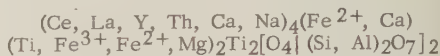
Chemical and powder diffraction data are given for fossilized dental enamel from a mastodon tooth.

The specific gravity is 2.968 ( $20^\circ/4^\circ\text{C.}$ ) and  $a_0 = 9.454$ ,  $c_0 = 6.892$  Å.

Calculations, analogous to those previously made for francolite, indicate that the carbonate groups enter the structure as substitutions for phosphate groups. However,  $(\text{CO}_3\text{OH})$  does not appear to be an accurate representation. The recent supposition that carbonate groups can substitute for  $(\text{F}, \text{OH})$  groups of normal apatite is untenable.--Auth.

2-1530. Gottardi, Glauco. THE CRYSTAL STRUCTURE OF PERRIERITE: *Am. Mineralogist*, v. 45, no. 1/2, p. 1-14, 3 figs., 7 tables, Jan.-Feb. 1960, 6 refs.

The crystal structure of perrierite (space group  $\text{C}2/\text{m}$ ) has been determined by X-rays. The following chemical formula has been attributed to the mineral:



The unit cell contains 2 formula units.

Perrierite is a sorosilicate, with a high number of O-atoms not bonded to Si. In its structure there are 2 types of Ti-O octahedral chains along the b-axis; the chains are connected in order to give a sheet parallel to (001). Each Ce atom is surrounded by 10 oxygens; each  $\text{Fe}^{2+}$  by 6 oxygens. Interatomic distances are of the expected magnitude.--Auth.

2-1531. Brown, G.M. THE EFFECT OF ION SUBSTITUTION ON THE UNIT CELL DIMENSIONS OF THE COMMON CLINOPYROXENES: *Am. Mineralogist*, v. 45, no. 1/2, p. 15-38, 5 diag., 2 tables, Jan.-Feb. 1960, 26 refs.

The unit-cell dimensions of 11 analyzed augites and ferroaugites from the Skaergaard intrusion [Greenland] are presented. The variation of the unit cell dimensions  $b$  and  $a \sin \beta$  is examined in relation to ion substitution, for the 23 analyzed common clinopyroxenes for which data are now available. For the 3 major substitution series,  $\text{Mg}^{2+}:\text{Fe}^{2+}$ ,  $\text{Ca}^{2+}:\text{Fe}^{2+}$  and  $\text{Ca}^{2+}:\text{Mg}^{2+}$ , the relationship between these 2 unit-cell dimensions and the ion substitution is regular, and linear equations are given for each series. The slope of the graph relating  $b$ -dimension to  $\text{Mg}^{2+}:\text{Fe}^{2+}$  substitution is identical to that for the orthorhombic pyroxenes. The 3 series are combined in a triangular diagram relating unit-cell dimensions to  $\text{Ca}^{2+}:\text{Mg}^{2+}:\text{Fe}^{2+}$  ratio, for the common clinopyroxene trapezium. The amount of substitution of  $\text{Al}^{3+}$  in octahedral coordination, and its effect on the  $b$ -cell dimension, is reconsidered.--Auth.

2-1532. Fleischer, Michael. STUDIES OF THE MANGANESE OXIDE MINERALS. III. PSILOMELANE: *Am. Mineralogist*, v. 45, no. 1/2, p. 176-187, 2 figs., 3 tables, Jan.-Feb. 1960, 23 refs.

Study of 14 analyses of psilomelane, including 7 new ones, verifies Wadsley's formula,  $(\text{Ba}, \text{H}_2\text{O})_4\text{Mn}_{10}\text{O}_{20}$ . X-ray study of the products of dehydration shows that the water is lost zeolitically up to  $500^\circ\text{C.}$  and that the mineral is converted to hollandite at about  $550^\circ\text{C.}$  Thermogravimetric and differential thermal analysis curves are given for 2 analyzed samples.--Auth.

2-1533. Smith, William Lee, Jerome Stone, Daphne R. Ross, and Harry Levine. DOVERITE,



POSSIBLE NEW YTTRIUM FLUOCARBONATE  
FROM DOVER, MORRIS COUNTY, NEW JERSEY:

Am. Mineralogist, v. 45, no. 1/2, p. 92-98, 4  
figs., Jan.-Feb. 1960, 5 refs.

Doverite, a possible new yttrium fluocarbonate,  
is found at the Scrub Oaks Fe mine at Dover. The  
mineral occurs in brownish-red aggregates mixed  
with xenotime, hematite, and quartz. The aggregates  
are irregular - some of them as large as 1 in. in  
diameter - and most of them have rims of bastnae-  
site.

In parts of the mine, the doverite aggregate con-  
stitutes several percent of the gangue. The aggre-  
gate has indices of refraction in the range from  
1.00 to 1.685. It is physically inseparable from the  
other components of the aggregate. Hematite and  
doverite were leached from the aggregate leaving  
a residue of quartz and xenotime. From the X-ray  
powder data, doverite is suggested to be an yttrian  
analogue of synchisite and to have a general formula  
 $\text{YCO}_3 \cdot \text{CaCO}_3$ , the Y in the formula including several  
elements of the rare-earth group.

The X-ray powder pattern of doverite bears a  
marked similarity to that of synchisite. The powder  
pattern of synchisite was indexed in terms of a  
pseudo-orthorhombic C-centered cell with  $a = 4.10$ ,  
 $b = 7.10$ , and  $c = 9.12$  Å. From these data the cell  
constants of doverite were calculated to be  $a = 4.07$ ,  
 $b = 7.06$ , and  $c = 9.12$  Å.

Inasmuch as doverite has not been satisfactorily  
identified from the aggregate, the possibility remains  
that doverite may be an Y-bearing synchisite. Until  
it can be proven to be an end-member of a series,  
its status as a new species is tentative. -- Auth.

1534. Milton, Charles, and Joseph J. Fahey.  
CLASSIFICATION AND ASSOCIATION OF THE CAR-  
BONATE MINERALS OF THE GREEN RIVER FORMA-  
TION. Am. Jour. Sci., v. 258-A (Bradley Volume),  
p. 242-246, map, 3 diagrs., table, 1960, 15 refs.

About 1/4 of all known carbonate minerals occur  
in the Green River formation of Wyoming, Utah, and  
Colorado, a sequence of Eocene lacustrine deposits.  
Most are alkali (Na) or alkali-earth (Ca, Mg, Ba) or  
calcium-alkali earth carbonates; a few contain rare  
elements, or Al, Cl, or phosphate. Twenty well-defined  
species are listed, and arranged in accordance with  
Fosberg's system, 7th ed. The particular assemblage  
found in each of the 3 states implies corresponding  
variation in the over-all geochemical and geological  
conditions affecting their deposition. -- Auth.

1535. Leo, G.W. AUTUNITE FROM MT.  
SPOKANE, WASHINGTON: Am. Mineralogist, v.  
45, no. 1/2, p. 99-128, 3 illus., map, 5 diagrs.,  
10 graphs, 10 tables, Jan.-Feb. 1960, 36 refs.

Near Mt. Spokane, Washington, coarsely crystal-  
line autunite is developed in vugs, fractures, and  
near zones in granitic rock. With the exception of  
dispersed submicroscopic uraninite particles, autun-  
ite is the only ore mineral in the deposits. A study  
of associated granitic rocks reveals that apatite,  
the most abundant accessory constituent, has been  
preferentially leached and corroded in mineralized  
zones, suggesting that it may have provided a  
source of lime and phosphate for the formation of  
autunite. Leaching may have been effected partly  
by meteoric water, but more probably was due to the  
action of ascending connate solutions that may also  
have carried U from unoxidized, as yet undiscovered  
deposits at depth.

Autunite from the Daybreak mine has been studied  
optically, chemically, and by X-ray diffraction. The  
autunite is commonly zoned from light-yellow mar-  
gins to dark-green or black cores, and autunite from  
the inner zone has a higher specific gravity and  
higher refractive indices than peripheral light mate-  
rial. X-ray powder diffraction patterns of dark and  
light meta-autunite formed from this autunite show  
no significant differences in the  $d$  spacings; how-  
ever, diffraction patterns of 9 zoned samples each  
show uraninite to be present in the dark, and absent  
from the light, phase.  $\text{UO}_2$  and  $\text{UO}_3$  determinations  
range from 0.66-0.70% and 57.9-58.0%, respectively,  
for light autunite, whereas dark autunite shows a  
range (in 7 determinations) of  $\text{UO}_2$  from 1.2 to 4.0%,  
and  $\text{UO}_3$  from 55.1 to 58.8%. The wide range of  
 $\text{UO}_2$  values in dark autunite is tentatively attributed  
to nonuniform distribution of discrete uraninite  
particles, which may also account for the dark color  
and higher density.

Thermogravimetric and differential thermal anal-  
yses of autunite suggest discrete water losses at  
about 90°, 145°, and 220°C. The first water loss  
probably represents dehydration to meta-autunite II,  
also recognizable by marked changes in optical prop-  
erties and the X-ray diffraction pattern. The form  
of the DTA curve above 90°C. resembles that of  
montmorillonite, suggesting that the dehydrations  
at about 145°C. and 220°C. may involve interlayer  
water as in montmorillonites, and the analogy with  
montmorillonite is further indicated by X-ray pat-  
terns of meta-autunite II heated just above these  
temperatures. Autunite heated to red heat shows a  
diffraction pattern distinct from all others. The  
cation exchange capacity of autunite, about 2.5 meq.  
per 100 g., is substantially lower than that previ-  
ously reported for artificial material. -- Auth.

2-1536. Milton, Charles, Edward C.T. Chao,  
Joseph M. Axelrod, and Frank S. Grimaldi. REED-  
MERGNERITE,  $\text{NaBSi}_3\text{O}_8$ , THE BORON ANALOGUE  
OF ALBITE, FROM THE GREEN RIVER FORMA-  
TION, UTAH: Am. Mineralogist, v. 45, no. 1/2,  
p. 188-199, 4 illus., map, 3 diagrs., 3 tables, Jan.-  
Feb. 1960, 6 refs.

Reedmergnerite,  $\text{NaBSi}_3\text{O}_8$ , is a B analogue of  
albite, occurring with eitelite, shortite, nahcolite,  
searlesite, leucosphenite, acmite, analcite, and  
magnesianiebeckite in unmetamorphosed dolomitic  
oil shales of the Green River formation in several  
oil wells in Duchesne County, Utah. It is triclinic,  
colorless, in stubby prisms with characteristic  
wedge-shaped ends; biaxial, negative,  $2V = 80^\circ$ ,  
indices of refraction  $\alpha 1.554$ ,  $\beta 1.565$ ,  $\gamma 1.573$ , all  
 $+0.001$ . X-ray diffraction and crystallographic  
data are given. -- Auth.

2-1537. Drysdall, A.R., and A.R. Newton. BLUE  
ASBESTOS FROM LUSAKA, NORTHERN RHODESIA,  
AND ITS BEARING ON THE GENESIS AND CLASSI-  
FICATION OF THIS TYPE OF ASBESTOS: Am.  
Mineralogist, v. 45, no. 1/2, p. 53-59, 2 tables,  
Jan.-Feb. 1960, 12 refs.

An occurrence of blue asbestos near Lusaka was  
investigated and proved to consist of an anasto-  
mosing stockwork of predominantly slip-fiber veins  
in a semicalcareous host rock. The host rock is  
probably lenticular and occurs in a strongly de-  
formed dolomite formation. Chemical analysis of  
the asbestos shows that it is magnesianiebeckite  
according to the classification of Miyashiro; the  
geological evidence strongly suggests that it is of

metasomatic origin. A similar occurrence has been described from Bolivia, and both are in strong contrast to the normal South African and Australian occurrences, where the asbestos is of riebeckite composition and probably of purely metamorphic origin. The name crocidolite has been widely used to describe blue asbestos, but as it applies only to fibrous varieties of riebeckite and magnesioriebeckite, it should not be applied as a mineral name.--Auth.

2-1538. Outerbridge, William F., Mortimer H. Staatz, Robert Meyrowitz, and Alfred M. Pommer. WEEKSITE, A NEW URANIUM SILICATE FROM THE THOMAS RANGE, JUAB COUNTY, UTAH: *Am. Mineralogist*, v. 45, no. 1/2, p. 39-52, illus., 5 tables, Jan.-Feb. 1960, 11 refs.

Weeksite,  $K_2(UO_2)_2(Si_2O_5)_3 \cdot 4H_2O$ , is a new uranyl silicate mineral named for Dr. Alice D. Weeks. It is a soft yellow nonfluorescent mineral with a waxy to silky luster, and it crystallizes chiefly in radiating fibrous clusters. In the Thomas Range, where it was first found, it occurs in opal veinlets in rhyolite and as replacements of pebbles in a tuffaceous conglomerate. Weeksite also has been identified from 8 other localities in Pennsylvania, Wyoming, California, New Mexico, Mexico, Arizona, and Texas.

Weeksite is biaxial negative  $2V = \text{about } 60^\circ$ ; dispersion  $r > v$ , strong;  $\alpha = 1.596$ ,  $\beta = 1.603$ ,  $\gamma = 1.606$ ;  $X = b$ , colorless,  $Y = c$ , pale yellow green,  $Z = a$ , yellow green. The measured specific gravity is about 4.1.

A chemical analysis of material from the Thomas Range showed:  $K_2O$  5.5,  $Na_2O$  0.7,  $BaO$  1.4,  $CaO$  1.1,  $UO_3$  51.5,  $SiO_2$  33.6,  $H_2O$  6.6,  $Al_2O_3$  0.6,  $CO_2$  0.3; total 101.3%. A synthetic Na analogue of weeksite showed  $Na_2O$  7.2,  $UO_3$  52.0,  $SiO_2$  33.9,  $H_2O$  (in part from wood) 8.1, C (from wood) 0.3; total 101.5%.

Weeksite is orthorhombic, pseudotetragonal; the space group is  $Pnnb-D_{2h}^{12}$ ;  $a = 14.26 \pm 0.02$  Å,  $b = 35.88 \pm 0.10$  Å,  $c = 14.20 \pm 0.02$  Å;  $a:b:c = .3974:1:.3958$ . Cell contents 16  $[K_2(UO_2)_2(Si_2O_5)_3 \cdot 4H_2O]$ . The 9 strongest lines of the X-ray powder pattern of weeksite are: 7.11 (10), 5.57 (9), 8.98 (8), 3.55 (7), 3.30 (7), 2.91 (6), 3.20 (5), 2.37 (5), 2.28 (5); and those of the synthetic Na analogue are 7.11 (10), 5.57 (9), 9.03 (8), 3.56 (7), 3.30 (7), 2.94 (6), 3.19 (5), 2.37 (5), 2.28 (5).--Auth.

2-1539. Serdyuchenko, D.P. THE COMPOSITION AND STRUCTURE OF DUMORTIERITE: *Akad. Nauk SSSR, Bull., Geol. Ser.*, in translation, 1958, no. 2, p. 102-106, table, pub. 1959, 22 refs.

English translation of GeoScience Abstracts 1-468.

2-1540. Kamb, W. Barclay, and William C. Oke. PAULINGITE, A NEW ZEOLITE, IN ASSOCIATION WITH ERIONITE AND FILIFORM PYRITE: *Am. Mineralogist*, v. 45, no. 1/2, p. 79-91, 5 illus., 3 tables, Jan.-Feb. 1960, 9 refs.

A new zeolite, for which the name paulingite is proposed, has been found in vesicles in basalt from the Columbia River near Wenatchee, Washington. Paulingite is cubic, and occurs as perfect, transparent rhombic dodecahedra. The very large cubic cell has  $a_0 = 35.10 \pm .02$  Å, and the space group is probably  $Oh^h - Im\bar{3}m$ . Ca and K are chief cations, but Ba and Na are also present. The atomic ratio of Si to Al is roughly 3. In association with the paulingite

is the rare zeolite erionite. It occurs as relatively large acicular hexagonal prisms, and also as tufts of very fine fibres. The hexagonal unit cell has dimensions  $a = 13.27 \pm .02$  Å,  $c = 15.05 \pm .02$  Å, and the space group is either  $D_{6h}^{24}$ ,  $C_{6v}^{12}$ , or  $D_{3h}^{12}$ . Erionite is a member of the chabazite group of zeolites. Also in association with the paulingite and erionite is pyrite of remarkable filiform habit.--Auth.

2-1541. Bassett, William A. ROLE OF HYDROXYL ORIENTATION IN MICA ALTERATION: *Geol. Soc. America, Bull.*, v. 71, no. 4, p. 449-455, 2 col. illus., 3 charts, 2 diags., table, Apr. 1960, 3 refs.

Phlogopite and biotite are generally more susceptible to alteration by ion exchange than muscovite. The key mechanism responsible for this difference is considered to be hydroxyl orientation.

Single-crystal study with infrared spectrophotometer established the orientation and bonding of the hydroxyl ions. The dipole moments of the hydroxyl ions in phlogopite, a trioctahedral mica, are perpendicular to the cleavage, whereas those in muscovite, a dioctahedral mica, are obliquely oriented.

The hydroxyl ions in phlogopite absorb a higher frequency ( $3,700 \text{ Cm}^{-1}$ ) infrared than those in muscovite ( $3,600 \text{ Cm}^{-1}$ ). Biotite and Fe-rich phlogopite generally show absorption at 2 frequencies, which indicates both trioctahedral and dioctahedral configurations. The inclined hydroxyl ions in dioctahedral mica place the K ions in a more negative environment than the perpendicular hydroxyl ions in trioctahedral mica. The more negative environment binds the K ions more securely in the lattice.--Auth.

2-1542. Williamson, W.O. SOME EFFECTS OF DEFORMATION ON THE STRUCTURE AND PROPERTIES OF CLAY: *Mineral Industries*, v. 29, no. 7, p. 1, 3-5, 8, 4 illus., 3 diags., Apr. 1960, 20 refs.

The particles of many clay minerals are in the form of very small, flat platelets and thus may develop planar parallelism during deposition, or during mechanical working. Such parallelism induces vector properties, e.g., drying and firing shrinkages tend to be differential. These vector properties play an important role in ceramic fabrication processes.

The effects of artificial deposition or deformation on the microstructure of clays have significance not only for ceramists but for geologists interested in the natural fabrics of rocks.

Engineers, soil physicists, and others may postulate a structure for plastic clay which is unacceptable to many ceramists; a possible resolution of the conflict of opinion is suggested.--Auth.

2-1543. Zeitner, June Culp. RARE GEMS OF THE MIDWEST: *Earth Sci.*, v. 13, no. 2, p. 58-62, illus., Apr. 1960.

Although the midwestern United States lacks the amount of gemstones found in the far West, it possesses a large variety of high-quality gem material. A brief description and locality designation is given of the following: thomsonite, chlorastrolite, datolite, Fairburn agate, Tepee Canyon agate, Superior agate, blue chalcedony, binghamite, dendritic opalite, Ohio flint, Petoskey stone, fresh-water pearls, gemfossils including cycad, *Lepidodendron*, osmundites, Iowa coral, and dendritic ivory.--M. Russell.



1544. Lapham, Davis M., and Alan R. Geyer. GENERAL COLLECTING IN PENNSYLVANIA: Pennsylvania Geol. Survey, Bull. G-33, 74 p, 58 figs. 1. illus., maps, 1959, approx. 45 refs.

The introductory topics covered in this bulletin state the hobby of mineral collecting to the science of mineralogy and to various aspects of the science of geology. In addition to the usual definitions, special sections are devoted to the geologic associa-

tions of minerals and to the use of topographic maps. Twenty-eight of Pennsylvania's best collecting localities are described by county. Each description includes word and map location directions; the types, forms, and availability of minerals present; a brief, nontechnical geologic summary of the occurrence; and specific references to that locality. Selected general references, a list of museum collections open to the public, and a mineral index conclude the bulletin.--Auth.

## 9. IGNEOUS AND METAMORPHIC PETROLOGY

also: Stratigraphy 2-1419; Geochemistry 2-1522.

1545. Friedman, Gerald M. CHEMICAL ANALYSES OF ROCKS WITH THE PETROGRAPHIC MICROSCOPE: Am. Mineralogist, v. 45, no. 1/2, p. 1-78, 2 tables, Jan.-Feb. 1960, 20 refs.

The chemical composition of the granite (G-1) was calculated from quantitative mineralogical (modal) analyses of petrographic thin-sections. The percentages of the major constituents,  $\text{SiO}_2$ ,  $\text{Al}_2\text{O}_3$ ,  $\text{Fe}_2\text{O}_3$ ,  $\text{FeO}$ ,  $\text{MgO}$ ,  $\text{CaO}$ ,  $\text{Na}_2\text{O}$ , and  $\text{K}_2\text{O}$  were computed. These calculations were compared with the results of wet chemical analyses made by 30 chemists in 25 analytical laboratories. Each constituent calculated from the thin-section analysis of the rock standard lies within the range reported by the chemists. The arithmetic mean of the percentage of each constituent determined by the chemical analyses agrees closely with the analysis calculated from the mode.

A norite sample was also analyzed by petrographic and chemical procedures. Agreement between the 2 techniques for this sample was close.

This study indicates that for rocks in which the composition of the individual minerals has been obtained by optical measurements, the major constituents can be determined with the petrographic microscope by modal analyses, and the results may be considered reliable.--Auth.

1546. Murata, K.J. A NEW METHOD OF PLOTTING CHEMICAL ANALYSES OF BASALTIC ROCKS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 247-52, 4 diag., 1960, 15 refs.

In studying the differentiation of basaltic magmas by means of the chemical analyses of basaltic rocks, very useful diagrams are obtained by plotting percent  $\text{CaO}$ ,  $\text{MgO}$ , or  $\text{Na}_2\text{O}+\text{K}_2\text{O}$  against the  $\text{Al}_2\text{O}_3/\text{SiO}_2$  weight ratio. The diagrams show the tholeiitic and the alkalic series of basaltic rocks nicely separated and parallel in their variations. Fractional crystallization of clinopyroxene appears to be the principal mechanism by which tholeiitic magmas are converted to alkalic magmas.--Auth.

1547. Waters, A.C. DETERMINING DIRECTION OF FLOW IN BASALTS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 350-366, 5 illus. on 2 pls., 5 figs. incl. secs., diag., 1960, 23 refs.

A number of structural features can be used to determine the direction of advance of individual lava flows in the great accumulations of flood basalts such as form the Columbia River plateau and the lava platform of the Oregon Cascades. Three are particularly valuable, however, because they are abundant, easily interpreted, and give unambiguous results: 1) inclined columnar joints in flow interiors

pitch toward the source of the lava; 2) spiracles (also pipe vesicles and vesicle cylinders) rise from the bottoms of flows and bend or trail out downstream; and 3) primary foreset bedding develops in most palagonite complexes.--Auth.

2-1548. White, Walter S. THE KEWEENAWAN LAVAS OF LAKE SUPERIOR, AN EXAMPLE OF FLOOD BASALTS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 367-374, sec., table, 1960, 19 refs.

Individual lava flows of Keweenaw (late Precambrian) age in the Lake Superior basin are characterized by volumes that may exceed 100 cu. mi. and by original surface slopes of the order of 10-20 ft. per mi. or less. In these respects they differ greatly from the flows of shield volcanoes from which so much of our knowledge of lava flowage stems. The volumes are estimated from field measurements. The slopes are estimated by subtracting reasonable initial dips of interbedded conglomerate beds (obtained by comparison with modern fan deposits of similar coarseness) from the angle between top and bottom of flows overlying conglomerate beds, since the lava flows and the gravel-depositing streams are believed to have flowed in opposite directions. Difference in volume alone seems adequate to explain a number of the physical differences between the Keweenaw flood basalts and those of shield volcanoes without appeal to differences in composition, volatile content, or temperature.--Auth.

2-1549. Ham, William E. GLASSY PEBBLES IN SOUTHWESTERN OKLAHOMA - OBSIDIAN VS. TEKITE: Oklahoma Geology Notes, v. 20, no. 4, p. 92-95, 3 illus., table, Apr. 1960, 5 refs.

A pebble, recently discovered at Lugert, Oklahoma, has the superficial appearance of a tektite. Consideration of the closely similar physical and chemical properties of tektites and obsidians shows that the pebble is an obsidian (principally because it has planar flow structure), as is also an earlier discovery from Delhi, Oklahoma, which has been reported to be a tektite.--A. Nicholson.

2-1550. Kudryashova, V.I. ON THE ORIGIN OF ELLIPSOIDAL LAVAS OF THE LOWER TUNGUSKA RIVER: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 87-92, 3 illus., secs., pub. 1959, 18 refs.

English translation of GeoScience Abstracts 1-470.

2-1551. Merlich, B.V. EXPLOSIVE BRECCIA DIKES OF TRANS-CARPATIA: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 41-50, 6 figs. incl. map, table, pub. 1959, 9 refs.

In the paper there are described the breccia dikes of Trans-Carpathia, previously regarded as tectonic. Proof is given of their explosive origin. According to the latest review, explosive dikes are not described in our literature, and data on such formations can only be obtained from American sources. Specific properties of explosive dikes are described, and the mechanism of formation of the dikes is outlined.--Auth.

2-1552. Zen, E-An. METAMORPHISM OF LOWER PALEOZOIC ROCKS IN THE VICINITY OF THE TACONIC RANGE IN WEST-CENTRAL VERMONT: *Am. Mineralogist*, v. 45, no. 1/2, p. 129-175, 6 diags., 5 tables, Jan.-Feb. 1960, 52 refs.

Rocks from the Taconic Range and adjacent areas, near Castleton in W.-central Vermont, have been studied for their mineral assemblages. Most of the rocks are argillaceous or arenaceous, of Early Cambrian to Middle Ordovician age, and are from the Taconic allochthone. A number of samples, including carbonates, are from the Lower Ordovician to Middle Ordovician rocks of the surrounding autochthonous marble belt. Regional metamorphism was later than the emplacement of the Taconic allochthone; the metamorphic grades in the Taconic sequence cut across structures and conform to the trend in the surrounding autochthone. The grade decreases from E. to W., and the rocks range correspondingly from phyllites to slates and shales.

Study of the mineralogy of the samples made use both of the X-ray diffractometer and the petrographic microscope. Two chlorite samples, 1 from a slate and 1 from a phyllite and a chloritoid sample from a phyllite, in addition to 4 rock samples, were chemically analyzed.

The most significant mineral assemblages found are: 1) muscovite-chlorite-quartz, 2) muscovite-paragonite-chlorite-chloritoid-quartz, 3) muscovite-chlorite-chloritoid-epidote-hematite-magnetite-quartz, 4) muscovite-stilpnomelane-chlorite-albite-microcline-quartz, 5) muscovite-biotite-quartz-graphite, 6) calcite-dolomite-chlorite-muscovite, and 7) calcite-zoisite-chlorite.

The prevalence of assemblage (1), with its simple mineralogy, indicates that chlorite may be variable in the (Fe, Mg)/Al ratio in addition to the Fe/(Fe+Mg) ratio. This is borne out by the chemical analyses; the analyzed chlorites are among the highest reported in their Al/(Fe, Mg) ratios.

The basal spacings of coexistent muscovite and paragonite in assemblage (2), found in the phyllites, are 9.97(5) Å and 9.62(3) Å, respectively, indicating very limited solid solution between these phases at this metamorphic grade. No paragonite has been found in the slates; it is suggested that paragonite may possess a lower limit of stability, with the more hydrous assemblage, kaolinite plus albite, taking the place of paragonite plus quartz.

Rocks relatively rich in Al contain chloritoid (2, 3), and the chloritoid-chlorite association is common even in slates. The chlorite is always more Mg-rich although the analyzed chloritoid also shows significant solid solution of Mg for Fe. The only sodic phase found with chloritoid is paragonite. Rocks relatively poor in Al do not have chloritoid but contain stilpnomelane (4), in addition to microcline which is incompatible with high-Al phases. The location of stilpnomelane on the low-Al part of the A-F-M diagram is based on regarding all Fe<sup>+++</sup> in stilpnomelane as subsequently oxidized Fe<sup>++</sup>; this is justified by an analysis of the available chemical data on this mineral. Stilpnomelane- and

chloritoid-bearing assemblages are mutually incompatible in the area; that the control is chemical composition rather than physical conditions of metamorphism is borne out by field mapping of the lithologic units. The stilpnomelane-microcline-chlorite assemblage rules out the chlorite-Fe biotite pair, which in fact is not found in the area although biotite does occur (5).

Data pertaining to phases in the A-C-F diagram are scanty. The mineral assemblages show, however, that chlorite may be the first mineral to form upon metamorphism of a quartzose and aluminous dolostone (6). The presence of the zoisite-calcite pair (7) indicates that the calcite-kaolinite pair, found in sedimentary rocks, ceases to be stable in the rocks of this area.

The analysis of the mineral assemblages in the Castleton area indicates that these rocks by and large have attained chemical equilibrium during metamorphism. The assemblages obey the Phase Rule, and no incompatible mineral pair has been found. Textural data on the rocks also suggest extensive recrystallization has occurred. Both H<sub>2</sub>O and CO<sub>2</sub> may be treated as mobile components without violating the Phase Rule. Evidence, however, is inconclusive on the status of O<sub>2</sub>.--Auth.

2-1553. Litsarev, M.A. GROSSULARITE-WILLASTONITE [SIC] SKARNS OF THE EMEL'DZHAK PHLOGOPITE DEPOSIT (SOUTH YAKUTIA): *Akad. Nauk SSSR, Bull., Geol. Ser.*, in translation, 1958, no. 3, p. 26-40, 3 illus., sec., diag., 4 tables, pub. 1959, 19 refs.

This paper presents the results of studies on grossularite-wollastonite skarns, which are unusual for the Archean complex of the Aldan massif. The geologic distribution of the skarns and their structure and mineral composition are discussed. A description of the minerals and their behavior in individual diffusive metasomatic zones is given. The problem of extent and depth of the system is also discussed.--Auth.

2-1554. Burge, Donald L. INTRUSIVE AND METAMORPHIC ROCKS OF THE SILVER LAKE FLAT AREA, AMERICAN FORK CANYON, UTAH: *Brigham Young Univ., Dept. Geology, Brigham Young Univ. Research Studies, Geology Ser.*, v. 6, no. 7, 46 p., 9 pls. incl. illus., maps (1 fold. geol. map), sec., Aug. 1959, 21 refs.

This report concerns 10 sq. mi. in the Silver Lake Flat area. Most of the area, which includes the southeastern part of the Little Cottonwood stock of quartz monzonite and associated metamorphosed Precambrian and Paleozoic sediments, is situated on the southern slope of the divide between American Fork Canyon and Little Cottonwood Canyon.

Metamorphism of the thin mantle of Mississippian limestones overlying the southeastern margin of the stock produced a striking suite of metamorphic minerals. The limestones were divided into 4 zones in order of increasing metamorphism: Zone 1, unmetamorphosed dark gray, slightly dolomitic limestone; Zone 2, dark gray to black amphibolite limestone containing abundant graphite; Zone 3, white to light gray siliceous limestone characterized by the presence of wollastonite, diopside, and idocrase; and Zone 4, abundant brown garnet with lesser amount of olivine, epidote, zoisite, serpentine, and microcline. Zone 2 represents a simple rearrangement of elements existing in a premetamorphosed limestone similar in composition to that of Zone 1,



whereas the limestones of Zones 3 and 4 have been strongly altered by the metasomatic introduction of large amounts of silica and smaller amounts of Al, Mg, and Fe. C and carbonate were eliminated from the 2 zones nearest to the intrusive contact, probably in the form of volatile compounds. A change in the atomic structure of the diagnostic minerals with increasing metamorphic rank is represented by the predominance of the double chain inosilicate structure in Zone 2, the single chain inosilicate structure in Zone 3, and the independent tetrahedral orthosilicate structure in Zone 4.

Prismatic amphiboles characteristic of Zone 2 show preferred orientations of axes concentrated in bedding planes of the limestone, with directions parallel to the strike and at right angles to the strike of the bedding. The most frequent orientation of the axes occurs at right angles to the strike of the bedding in a northwestern direction, closely paralleling a line drawn toward the center of the stock. The W. axes were probably orientated by slippage along the bedding planes of the limestone resulting from stresses that radiated from the stock during its intrusion. Orientation of amphibole axes along the strike of the limestone bedding was most likely due

to secondary slippage produced by tensional stresses working at right angles to the compressive stresses.

Mines in the area have produced moderate amounts of Pb, Zn, and Ag in the past, but are presently inactive. Most of the ore production has been from northeasterly striking fissures in the Tintic quartzite.--Auth.

2-1555. Afanasev, G.D., and S.G. Tseytlin. ROCK RADIOACTIVITY STUDY IN THE NORTHERN CAUCASUS AND ITS IMPORTANCE FOR CERTAIN PETROLOGIC PROBLEMS (PRELIMINARY CONCLUSIONS): Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 12-25, graph, 10 tables, pub. 1959, 22 refs.

A schematic representation is given of the accepted separation of magmatic rocks in the N. Caucasus into complexes of various age and their distribution in the structural zones of the N. Caucasus folded province. The following 4 magmatic complexes are recognized 1) Cenozoic (Tertiary-Quaternary), 2) Mesozoic (Jurassic-Cretaceous), 3) upper Paleozoic (lower Carboniferous-Permian) 4) Urushten (Cambrian-Lower Devonian). Numerous age determination data are given.--M. Russell.

## 10. SEDIMENTARY PETROLOGY

See also: Structural Geology 2-1376; Stratigraphy 2-1398; Mineralogy 2-1527, 2-1542; Engineering Geology 2-1618.

2-1556. Flint, Richard Foster, John E. Sanders, and John Rodgers. SYMMICTITE: A NAME FOR CONSORTED TERRIGENOUS SEDIMENTARY ROCKS THAT CONTAIN A WIDE RANGE OF PARTICLE SIZES: Geol. Soc. America, Bull., v. 71, no. 4, p. 507-509, Apr. 1960, 12 refs.

Existing terminology of till-like sedimentary deposits tends to emphasize genesis without considering all possible modes of origin and to suggest relationships to tillite that may be erroneous. Therefore new terms (symmictite for sediment; symmictite, or lithified equivalents) are proposed as general names for essentially nonsorted, noncalcareous, terrigenous deposits composed of sand and/or larger particles in a muddy matrix.--Auth.

2-1557. Tourtelot, Harry A. ORIGIN AND USE OF THE WORD "SHALE": Am. Jour. Sci., v. 258-A (Bradley Volume), p. 335-343, 1960, 42 refs.

Shale is a word of Teutonic origin that developed its meaning of "laminated clayey rock" in the English mining district of Derbyshire. The first recorded use is in 1747, but clearly the term is much more ancient. By the first few years of the 19th century, it had superseded a number of local names as well as the classically derived schistus, which the early natural philosophers were accustomed to use. By the end of the 19th century, it was nearly the only term applied to clayey rocks of Jurassic age and older in England and to clayey rocks of all ages in the United States, except for the coastal plain strata of Cretaceous and Tertiary age. The word is now looked on with some disfavor because we use it both as the class name for fine-grained rocks and more precisely as the name for laminated clayey rocks. Both of these uses, however, seem permissible and justifiable.--Auth.

2-1558. Pryakhina, Yu. A. CARBONATE CONCRETIONS IN THE MAYKOP DEPOSITS OF THE

CENTRAL CIS-CAUCASUS: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 1, p. 18-31, 13 figs., incl. 5 illus., 3 maps, graphs, diag., 4 tables, pub. 1959, 6 refs.

Four groups of carbonate concretions occur in the Maykop deposits of the central Cis-Caucasus and, based on their chemical, mineral, and petrographic properties, are divided as follows: 1) clay siderites, 2) clay-phosphate siderites, 3) clay ankerites, and 4) clay calcites. The enclosing sediments are marine, deposited in a reducing environment.--M. Russell.

2-1559. Bagnold, R.A. SEDIMENT DISCHARGE AND STREAM POWER - A PRELIMINARY ANNOUNCEMENT: U.S. Geol. Survey, Circ. 421, 23 p., 1960.

This report describes sediment transport by water streams and, by reasoning, extends the simplest conditions of uniform grain size and steady flow to natural materials and unsteady flow. Sediment discharge is discussed as a function of stream power measured by a function of discharge and slope.--U.S. Geol. Survey.

2-1560. McKee, Edwin D. CYCLES IN CARBONATE ROCKS: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 230-233, fig., 1960, 10 refs.

Sedimentary cycles are readily recognized in rock sequences involving both marine and continental strata and to a lesser extent in strata that are entirely marine but include both carbonate and detrital types. In rock sequences composed entirely of carbonate rock, cycles commonly can be detected by trends in the size and abundance of terrigenous components. In carbonate formations that are largely devoid of detrital minerals, however, cycles must be established on the basis of differences in size of clastic carbonate grains which may indicate differences in turbulence but are not necessarily the results of transportation. Trends in distribution obtained from a detailed analysis of cycles, however,

may indicate whether the position of various size grades has been determined by transportation or by environment.--Auth.

- 2-1561. Cloud, Preston E., Jr. GAS AS A SEDIMENTARY AND DIAGENETIC AGENT: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 35-45, 13 illus., 1960, 38 refs.

Gas can be a significant factor in the transportation and alteration of sediments, and traces of its former presence may convey useful information about the depositional and diagenetic history of sedimentary rocks. Transportation of sediment by gas is probably most important in embayed littoral environments of broad extent and reduced offshore drainage, hence onshore transport and shoreward decreasing grain size. Stranded, blown, or drifting bubbles and evasion of upwelling gas are responsible for a variety of imprints and structures in sediments. The relations to other features of impressions of gas blisters at the base of a thin Devonian graywacke are analyzed and found to support deposition of the graywacke from a weak suspension current. Concentration of pyrite in irregularly subcylindrical structures within a Devonian limestone contributes to the interpretation of these structures as former sites of gas accretion. Relations between pyrite, matrix, and later mineralization reveal the main steps in the diagenetic history of the limestone. Some ancient sedimentary features that have been interpreted as raindrop imprints, tracks and burrows, and Precambrian jellyfish are probably also the work of gas.--Auth.

- 2-1562. Lovering, T.S., and Anna O. Shepard. HYDROTHERMAL ALTERATION ZONES CAUSED BY HALOGEN ACID SOLUTIONS, EAST TINTIC DISTRICT, UTAH: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 215-229, 6 illus., chart, 9 profiles, 3 tables, 1960, 6 refs.

On the Helen claim in Homansville Canyon in the East Tintic district, a small deposit of clay typical of argillic alteration, together with fluorite, diaspore, quartz, hematite, and manganese oxides, occurs at the intersection of a minor N-S fissure with a steep ENE-trending contact of Tertiary quartz latite lava and dolomitized Cambrian Herkimer limestone. The fissure ends at a small completely argillized monzonite plug about 100 yds. to the N.

Hydrothermally altered rocks are scattered through the roots of the deeply eroded Eocene volcanic field in which the East Tintic and Tintic districts lie, but in most areas of argillic alteration, abundant pyrite or alunite and barite suggest that the alteration was caused largely or in part by S acids, whereas in the Helen claim these minerals are virtually lacking.

Mineral zones parallel the lava-dolomite contact: the lava is strongly argillized in a zone about 4 ft. wide, which has much montmorillonite and a small amount of kaolinite. The dolomite outward from its contact with the lava has been replaced by hematite and quartz in a zone 5-8 ft. wide, which merges on the far side with a manganese zone 1-3 ft. thick containing abundant quartz, manganese oxides and hematite, minor manganapatite, and kaolin minerals; farther from the contact, a fluorite-kaolin-diaspore zone, a diaspore-kaolin-fluorite zone, a kaolin-diaspore-fluorite zone, and a kaolin-mica zone appear successively. The edge of the dolomite is changed to a weakly bonded granular dolomite

with some kaolin and mica; beyond this zone hard fresh hydrothermal dolomite extends outward for several hundred feet.

The field relations of the alteration zones, together with X-ray, mass spectrograph, and petrographic studies, suggest reaction with acid emanations carrying halides of Fe, Al, and Si at temperatures up to 300°C. The zoning is appropriate to precipitation responsive to an increasing pH caused by reaction with the carbonate wall rock.--Auth.

- 2-1563. Grim, Ralph E., Georges Kulbicki, and Albert V. Carozzi. CLAY MINERALOGY OF THE SEDIMENTS OF THE GREAT SALT LAKE, UTAH: *Geol. Soc. America, Bull.*, v. 71, no. 4, p. 515-519, map, Apr. 1960, 3 refs.

The modern lake sediments and those of post-Provo age immediately preceding the present Great Salt Lake have about the same clay-mineral composition. In these deposits, the montmorillonite shows a relatively poor organization which may be a consequence of the presence of considerable Na. The montmorillonite may be either detrital or precipitated or formed by alteration of volcanic ash.

The illite, probably detrital, is generally well organized showing only a very small amount of degrading. The presence of kaolinite in all the lake samples is noteworthy. It is almost certainly detrital. There is a general absence of substantial mixed-layer components as well as of any chlorite, and there is no suggestion of attapulgite-sepiolite clay minerals.

The Alpine and Provo sediments of the Lake Bonneville terraces have a clay-mineral composition somewhat different from that of the modern lake sediments; they show a highly degraded chlorite and a very well-ordered montmorillonite.

The green clays of early Pleistocene age also have a distinctive clay-mineral composition corresponding to a high content of illite and a low content of montmorillonite associated with the presence of chlorite.

The clays occurring in the Great Salt Lake desert W. of Knolls contain at a depth of about a foot a thin bed of nearly pure, unconsolidated dolomite.--Auth.

- 2-1564. Terriere, Robert T. GEOLOGY OF GROSVENOR QUADRANGLE, TEXAS, AND PETROLOGY OF SOME OF ITS PENNSYLVANIAN LIMESTONES: *U.S. Geol. Survey, Repts., Open-File Ser.*, no. 505, 171 p., 28 figs. incl. 2 maps, secs., diagrs., 10 pls. incl. illus., geol. map (in pocket), sec. (in pocket), 2 graphs (in pocket), Jan. 1960, 45 refs.

The geology of the Grosvenor quadrangle, Brown and Coleman counties, has been mapped as a part of a systematic mapping program by the U.S. Geological Survey. In support of stratigraphic studies of the rocks of the area, detailed petrographic examinations were made of key limestone beds.

The quadrangle is underlain by Pennsylvanian and lower Permian sedimentary rocks that dip gently to the WNW., overlapped by isolated patches of Lower Cretaceous strata which dip slightly to the SE., and covered by Quaternary alluvium along stream valleys. The Pennsylvanian and Permian rocks consist mainly of gray shale, but also contain limestone, siltstone, sandstone, conglomerate, and red shale. The sandstone and conglomerate occupy broad, shallow channels cut during Pennsylvanian and early Permian time. Shale and siltstone also fill parts of some channels.



These rocks belong to the Canyon, Cisco, and Wichita groups, in ascending order. The Canyon group contains thicker limestone units and a larger proportion of limestone than the Cisco and Wichita groups, but has few channel deposits. Channel deposits are numerous in the Cisco and Wichita, especially in the lower part of the Wichita. Where channel deposits are superimposed on one another, they can be differentiated only by detailed study of their relationships to limestone beds. Smaller channels that seem to be the result of local cut and occur within some of the main channels. Limestone units form a small part of this sequence, but are of particular interest because they constitute distinctive marker beds and because much thicker Pennsylvanian and Permian limestone units are important oil and gas reservoirs in western Texas. Seven limestone units in the Bluff Creek and Gunsight members of the Graham formation of the Cisco group, considered representative, were selected for detailed study. Smoothed and etched surfaces, insoluble residues, and thin sections were studied with the aid of binocular and petrographic microscopes. Point counts were made to determine the abundance of the various organic and inorganic components. Shale and sandstone associated with the limestone units were studied separately. The limestone beds are composed of 1) interstitial calcite ooze and sparry calcite cement, 2) grains or aggregates (allochems), including fossils, oöliths, pellets, and reworked limestone fragments (intraclasts), 3) terrigenous sand, silt, and clay, and

4) authigenic chert, barite, pyrite, ankerite, hematite, limonite, and psilomelane(?). The limestone beds are predominantly made up of fossils and fossil fragments in a mixture of calcite ooze and cement. The constituents most diagnostic of certain units are fusulinids, algae, corals, and the insoluble residues. The relative amounts of these constituents, along with texture and field appearance, were used to divide the limestone samples into 4 types.

Type I is unsorted, unwinnowed limestone with many fusulinids, dark gray where fresh but orange where weathered. It contains finely disseminated pyrite, suggesting little water circulation and hence deposition below wave base. Type II limestone is aphanitic and has many recrystallized algae and veinlets of sparry calcite. Locally it contains many horn corals. Probably type II limestone was deposited in extremely shallow water where waves were small and conditions were most favorable for chemical precipitation of carbonate ooze.

Type III limestone contains abundant evidence of strong wave action: sorting and rounding, reworked fossils, intraclasts, oöliths, and relatively large amounts of terrigenous sand. This type has erratic distribution and probably originated on offshore bars. Type IV limestone is light gray, aphanitic, and contains few fusulinids or algae, but much poorly sorted and poorly preserved fine fossil debris, especially smaller Foraminifera. Field relations between type IV and type I limestone suggest that they are regressive and transgressive equivalents, respectively, formed in otherwise the same environments. --Auth.

## II. GEOHYDROLOGY

See also: Geochemistry 2-1523; 2-1525; Sedimentary Petrology 2-1559.

1565. Leopold, Luna B. CONSERVATION AND WATER MANAGEMENT: U.S. Geol. Survey, Circ. 4, 4 pts., 1960.

A series of 4 brief, general addresses, presented before various groups in the United States during 1958 and 1959: Pt. A. Conservation and Protection; Pt. B. The Challenge of Water Management; Pt. C. The Conservation Attitude; Pt. D. Ecological Systems and the Water Resource.

1566. Magin, George B., Jr., and Lois E. Randall. REVIEW OF LITERATURE ON EVAPORATION SUPPRESSION: U.S. Geol. Survey, Prof. Paper 272-C, p. 53-69, 1960, 322 refs.

Evaporation suppression is the reduction of evaporation by controlling the rate at which water vapor escapes from water surfaces. The need for water saving is greatest in areas of little rainfall and low runoff. Water losses by evaporation from storage reservoirs must be minimized for greatest utility with limited supplies.

A review of the literature shows that evaporation can be retarded by either physical or chemical means. It can be controlled by enclosure in a building, application of plastic covers, oil films, and monomolecular films.

Monomolecular films hold the greatest promise for use on large areas of water such as reservoirs. From the literature available it is apparent that these films introduce many chemical and physical problems. The suppressants must be nontoxic and

odorless, especially if used on water-supply reservoirs. They should allow the free exchange of O<sub>2</sub> so that aquatic life is not endangered. The effects of water-borne bacteria on these films and the effects of the films on water of different qualities are not fully known. The physical problem of detecting the presence of the film has not yet been solved. Of the monomolecular films that have been investigated, the waxy alcohols such as hexadecanol and octadecanol show the greatest promise. --Auth.

2-1567. Stallman, Robert W. NOTES ON THE USE OF TEMPERATURE DATA FOR COMPUTING GROUND-WATER VELOCITY: U.S. Geol. Survey, Ground Water Notes, no. 39, 17 p., diag., Jan. 1960, 4 refs.

Analytical methods now in use for calculating the hydraulic constants of aquifers require the direct measurement of water velocity along some known boundary. This requirement limits the possibility of field determination of the desired constants to those areas where adequate testing facilities are available. Indirect methods of measuring ground-water velocity may serve to reduce this limitation.

The simultaneous transfer of heat and water underground results in head and temperature distributions amenable to description by differential equations. A sample equation is derived for a system isotropic and homogeneous to heat and fluid flow. Sample computations show that, for a "typical" set of conditions in water-bearing formations, the ground-water velocity has an appreciable influence on the underground temperature distribution. Thus it appears feasible to utilize temperature measure-

ments for calculating flow velocity. Also, a combination of head and temperature measurements might be used for calculating aquifer permeability.--Auth.

2-1568. Bear, Jacob. SCALES OF VISCOUS ANALOGY MODELS FOR GROUND WATER STUDIES: Am. Soc. Civil Engineers, Hydraulics Div., Jour., v. 86, no. HY2, pt. 1, p. 11-23, 6 diag., Feb. 1960, 13 refs.

Scales for viscous fluid models used for studying various problems of flow through porous media are described. Consideration is given to some model techniques and limitations. The discussion covers scales for length (including distorted scales), time, discharge, volume, storage coefficient, 2 liquid flows, and capillary effects.--Auth.

2-1569. Nixon, Paul R., and G. Paul Lawless. TRANSLOCATION OF MOISTURE WITH TIME IN UNSATURATED SOIL PROFILES: Jour. Geophys. Research, v. 65, no. 2, p. 655-661, 6 graphs, 6 tables, Feb. 1960, 3 refs.

Downward translocation of moisture in soil profiles under various types of natural vegetation and a denuded plot was observed during a prolonged rainless period. Moisture determinations were made to 20-ft. depths with a neutron-scattering moisture meter. The observations were made as part of a study of ground-water recharge by deep penetration of rain water. The significant magnitude that translocated moisture may reach is illustrated by data obtained in sand under brush cover. In this case, deeply translocated moisture was equal to 159% of evapotranspiration during the first rainless month. Approximately 31% of the moisture content of a 20-ft. profile under a denuded plot was lost by downward movement from the first to the 240th day of a rainless period. The moisture content  $W$  of various soil layers under the plot varied with the  $T$  according to the relation  $W = aT^{-b}$ .--Auth.

2-1570. Kirkham, Don. SEEPAGE INTO DITCHES FROM A PLANE WATER TABLE OVERLYING A GRAVEL SUBSTRATUM: Jour. Geophys. Research, v. 65, no. 4, p. 1267-1272, 3 figs., 2 tables, Apr. 1960, 4 refs.

The problem of flow of ponded water over a stratum of homogeneous soil overlying a stratum of gravel, the gravel in turn overlying an impermeable layer, is solved for the case in which equally spaced ditches penetrate into the gravel. Flow nets are obtained, as are formulas giving the quantity of water seeping through the ground into the ditches. The theory shows that unless the ditches are closer together than about 3 times the depth of the stratum of soil overlying the gravel, the proportion of water entering the ditches through the ditch walls is small compared with that reaching the ditches by way of the underlying gravel stratum.--Auth.

2-1571. Luthin, James N., and J. W. Holmes. AN ANALYSIS OF THE FLOW OF WATER IN A SHALLOW, LINEAR AQUIFER, AND OF THE APPROACH TO A NEW EQUILIBRIUM AFTER INTAKE: Jour. Geophys. Research, v. 65, no. 5, p. 1573-1576, 3 figs., table, May 1960, 5 refs.

The problem of predicting changes in the water table due to changes in the amount of water supplied

is analyzed for a specific case. The differential equation describing the situation is

$$(\epsilon/km) \partial\phi/\partial t = \sigma^2 \phi / \partial x^2 - A/k$$

where  $\epsilon$  is a storage factor,  $k$  the hydraulic conductivity,  $m$  the thickness of the aquifer,  $\phi$  the hydraulic head, and  $A$  the steady-state loss due to deep seepage and/or evapotranspiration.

The Laplace transformation is used for solving the above equation, subject to an appropriate set of boundary conditions.--Auth.

2-1572. Wood, Perry R., and George H. Davis. GROUND-WATER CONDITIONS IN THE AVENAL-MCKITTRICK AREA, KINGS AND KERN COUNTIES, CALIFORNIA: U.S. Geol. Survey, Water-Supply Paper 1457, 6 maps (3 in pocket), 3 secs. (in pocket), diag., 7 graphs (1 in pocket), 10 tables, 1959, 47 refs.

The Avenal-McKittrick area consists of about 850 sq. mi. on the SW. side of the San Joaquin Valley. It extends southeastward from the towns of Avenal and Kettleman City to the Elk Hills. Except for about 16,000 acres in several isolated localities, the area is uninhabited desert most of the year, although extensive grazing is carried on during winter and spring. The area includes Kettleman and Antelope plains, Antelope and McLure (also known as Sunflower) valleys, and the area S. of Tulare Lake.

The geologic units of the area are classed as consolidated rocks and unconsolidated deposits. The first group includes consolidated non-water-bearing rocks and semiconsolidated rocks of Jurassic to Pliocene age which in general contain connate water of poor quality; locally at shallow depth fresh water is present in them. Most of the formations contain fold- and fault-induced fractures, and these openings probably convey small quantities of water to the adjacent unconsolidated deposits. The second group consists of unconsolidated to loosely consolidated water-bearing deposits of continental origin which supply nearly all the water pumped from wells. These deposits include the structurally deformed Tulare formation of late Pliocene and Pleistocene(?) age and the alluvium of Pleistocene and Recent age, which in most places unconformably overlies the Tulare. Both units are composed of generally poorly sorted silty materials containing lenticular bodies of sand and gravel derived from the Coast Ranges and deposited by streams on extensive alluvial fans. The alluvium is only moderately permeable, and, although the water-bearing properties of the Tulare are little known, it probably is also only moderately permeable. Lacustrine and flood-basin sediments of Pleistocene and Recent age, present in the bed of Tulare Lake and along Buena Vista slough, consist of nearly flat-lying well-stratified silt, clay, and fine sand which generally are poorly permeable.

The alluvium and the Tulare formation range in thickness from a few feet along the western border of the valley to several thousand feet beneath the valleys and plains. However, because of comparatively late structural deformation in the area, the deposits vary considerably in thickness for short distances. Over the crests of buried anticlines the deposits thin considerably, and along the flanks and in adjoining synclinal troughs they thicken rapidly.

Recharge to the unconfined and semiconfined bodies of ground water may be by seepage loss from streams, by deep penetration of imported water applied for irrigation in excess of plant requirements,



losses of imported water through irrigation canals and ditches, and by deep penetration of rainfall. However, seepage loss from intermittent streams draining the Coast Ranges probably is the chief source of recharge. Poorly defined drainage channels that extend generally eastward across the area suggest that little water escapes as surface outflow. Inasmuch as annual precipitation averages less than 12 in., deep penetration of rainfall is a significant source of recharge only during infrequent years of exceptionally large precipitation.

Although the general movement of ground water is northeastward from recharge areas along the eastern border of the valley toward Buena Vista, although in the trough of the valley, the movement is interrupted or deflected in several places by north-eastward-trending anticlinal structures, principally those underlying the Pyramid, Kettleman, and Lost Hills. Movement is restricted in the narrow Dagany and Avenal gaps and the area between the Kettleman and Lost hills. Local pumping depressions have developed in McLure and Antelope valleys.

Of the 344 wells that were investigated in the field in 1951 and 1955, only 95 were used as irrigation wells in the autumn of 1955. Short-term drawdown and recovery tests made by the Pacific Gas and Electric Co. indicate that the yields ranged from about 100 to 1,700 g.p.m. (gallons per minute). Specific capacities ranged from 2 to 120 g.p.m. per ft. of drawdown. The depths of irrigation wells ranged from about 100 to 1,700 ft.

Irrigation pumpage by electrically powered pumps increased steadily from about 13,000 acre-ft. in 1947 to about 25,000 acre-ft. in 1953. From 1953 to 1956 the annual pumpage remained nearly constant, about 25,000 acre-ft. The total for the 10-year period was about 200,000 acre-ft.

In areas of large withdrawals water levels have undergone steady and marked declines. For the period 1936-1953, the levels at Devils Den ranch declined a maximum of about 100 ft.; for the period 1951-1956, levels in Antelope Valley declined a maximum of 40 ft.; and for the period 1947-1954, levels in McLure Valley declined 60 to 120 ft. In the areas of little ground-water development E. of Avenal gap and throughout most of Kettleman and Antelope plains, water levels remained essentially unchanged to 1955, the date of this report.

The stream waters tributary to the Avenal-McKittrick area differ greatly in both chemical character and mineral content. They were divided into 2 general types on the basis of their chemical character. The waters of Avenal and Polonio creeks are characterized by a comparatively lower mineral content and proportionally less sulfate but more Mg than the waters of Bitterwater, Media Agua, and Carneros creeks. These differences in mineral content are related to the lithology of the rocks in the respective drainage areas. The waters in which the concentration of sulfate is relatively low were derived from areas underlain predominantly by marine sediments of Cretaceous age and sedimentary, igneous, and metamorphic rocks of the Franciscan formation of Jurassic and Cretaceous age. The waters characterized by a high proportion of sulfate were derived from areas underlain chiefly by marine and continental sedimentary rocks of Tertiary and Quaternary age.

The ground waters of the area are fairly consistent in chemical character but differ greatly in mineral content. The typical waters are of sodium sulfate composition or are sulfate waters of intermediate cation composition. Locally, sodium chloride

waters are present. The mineral content of ground waters in the area ranges from 477 to 7,040 p.p.m. (parts per million), hardness ranges from 30 to 3,020 p.p.m. and B from 0.3 to 11 p.p.m. The high Na content of the ground waters in comparison to that of the stream waters evidently is the result of cation exchange with the sediments.

In most of the area the ground waters are of doubtful to unsuitable quality for irrigation. Further detailed studies of soil salinity and quality of ground water, and exploration and testing of undeveloped parts of the area will be required to determine whether ground-water supplies or ground-water storage can be utilized effectively in conjunction with imported surface-water supplies.--Auth.

2-1573. Sherwood, C.B. GROUND-WATER RESOURCES OF THE OAKLAND PARK AREA OF EASTERN BROWARD COUNTY, FLORIDA: Florida Geol. Survey, Rept. Inv. no. 20, 40 p., 9 maps, diag., 9 graphs, 4 logs, 2 tables, 1959, 13 refs.

The Biscayne aquifer is the source of all fresh ground water in the Oakland Park area. This aquifer extends from the land surface to more than 215 ft. below mean sea level and is composed chiefly of sandy marine limestone, calcareous sandstone, and beds of fine to medium quartz sand. The aquifer differs from place to place, but, in general, most of the layers of limestone and sandstone occur at depths below 60 ft. The permeability of the aquifer increases with depth.

Wells for small supplies generally obtain water at depths ranging from 60 to 80 ft., whereas wells for large supplies usually obtain water from the interval between 100 and 200 ft. Large-diameter wells obtain as much as 1,000 g.p.m. (gallons per minute) from the lower part of the aquifer.

Chemical analyses of ground-water samples indicate a hard limestone water that is suitable, naturally or with treatment, for most ordinary uses. Periodic determinations of chloride content of the ground water show that some salt-water encroachment has occurred in areas near the coast and in the Middle River basin.

Pumping-test data for deep wells in the Prospect well field area indicate approximate aquifer coefficients of transmissibility and storage of 2,000,000 g.p.m. per ft. and 0.015, respectively. However, the data indicate also that the hydraulic characteristics of the aquifer are complicated by the presence of beds of sand, silt, and clay in the upper 100 ft. of the aquifer and by recharge from surface-water sources. Quantitative data and areawide water-level and salinity data indicate that large quantities of ground water are available for future development if salt-water encroachment can be effectively controlled.--Auth.

2-1574. Peek, Harry M. THE ARTESIAN WATER OF THE RUSKIN AREA OF HILLSBOROUGH COUNTY, FLORIDA: Florida Geol. Survey, Rept. Inv. no. 21, 96 p., 15 maps (1 fold.), 2 secs., diag., 31 graphs, 7 tables, 1959, 17 refs.

The Ruskin area is defined to include about 200 sq. mi. in southwestern Hillsborough County. Sand, limestone, and shells of Pleistocene and Pliocene age crop out and vary in thickness from a few feet to about 60 ft. These deposits are underlain by Hawthorn formation (middle Miocene) which varies in thickness from less than 10 ft. in the northern part to more than 150 ft. in the southern part. The

Tampa formation (early Miocene) underlies the Hawthorn. Other formations encountered in water wells are the Suwannee limestone (Oligocene), Ocala group (upper Eocene) and Avon Park limestone (middle Eocene).

The Hawthorn and younger sediments are the source of some domestic and other small water supplies, but the large quantities of water required for irrigation and industrial use are obtained from older limestones. Suwannee limestone and Tampa formation are the principal sources of artesian water. The water in these formations occurs in permeable zones, separated by relatively impermeable beds. The water is replenished by rainfall in western Polk County and eastern Hillsborough County. The artesian aquifer has a transmissibility coefficient of about 115,000 g.p.d. per ft. and a storage coefficient of about 0.0006.

Significant fluctuations of artesian-pressure head result from daily and seasonal variations in withdrawal of water. The artesian-pressure head declined progressively in the coastal area during a period of extensive agricultural development from 1950 to 1952. Since 1952, however, seasonal fluctuations have decreased in magnitude, and a slight progressive rise in artesian head has occurred locally, as a result of a decrease in withdrawals. Records of water levels in wells not affected by local variations in discharge indicate that, regionally, the artesian head declined progressively in 1955-1956.--H. S. Puri.

2-1575. Peek, Harry M. RECORD OF WELLS IN THE RUSKIN AREA OF HILLSBOROUGH COUNTY, FLORIDA: Florida Geol. Survey, Inf. Circ. no. 22, 85 p., fig., pl., 2 tables, 1959.

This publication is a companion report to "The Artesian Water of the Ruskin Area of Hillsborough County, Florida" by the same author (see abstract above). This report contains tables of well records compiled during investigation of the artesian water in the Ruskin area.--H. S. Puri.

2-1576. Mogg, Joe L., Stuart L. Schoff, and E. W. Reed. GROUND WATER RESOURCES OF CANADIAN COUNTY, OKLAHOMA: Oklahoma Geol. Survey, Bull. 87, 112 p., 4 maps (2 in pocket), graph, 9 tables, 1960, 36 refs.

Canadian County, in central Oklahoma, has an area of 891 sq. mi. The average annual precipitation is about 29 in., and the normal annual temperature is about 60°F. Farming is the principal occupation and wheat the major crop. Production of oil and gas from the West Edmond pool, flour milling companies, the railroad shops at El Reno, and the Mustang power plant near Lake Overholser are the principal industries.

The rocks at the surface range in age from Permian to Quaternary. The primary purpose of this investigation was to determine the water-bearing properties of the Quaternary terrace deposits and alluvium along the North Canadian River, but the water-bearing character of the Permian rocks also was studied.

Alluvium and terrace deposits consist of inter-fingering lentils of clay, sandy clay, sand, and gravel. The average thickness of the saturated materials encountered in 156 test holes drilled into the alluvium of the North Canadian River valley was 27 ft. The alluvium of the North Canadian River is the source of water supply for the municipalities of El Reno, Yukon, Okarche, Geary, and Calumet.

Oklahoma City supplements its surface-water supply with water from the alluvium, and the Concho Indian Reservation gets its water from the same source. Additional users are the Mustang power plant, which uses ground water for cooling purposes and irrigation farmers, with 15 wells which are used to irrigate 609 acres in the river valley. The water in the alluvium and terrace deposits of both river valleys is generally hard and locally contains excessive quantities of sulfate. Yields from wells in these deposits may be as much as 600 g.p.m., depending upon the amount of saturated sand and gravel present and the construction of the well.

Wells in the Rush Springs sandstone, which crops out in the southwestern part of the county, may be expected to yield up to 80 g.p.m. depending on the thickness of the saturated sandstone present. Water from these sandstones is used for domestic purposes. The other Permian rocks in the county are considered to be poor aquifers which yield small amounts of highly mineralized water.

Aquifer tests were made on 3 wells in the alluvium of the North Canadian River valley. Results of these tests showed that the average coefficient of permeability is about 1,000 g.p.d. per sq. ft.; and the coefficient of storage, within range of fluctuation of the water table due to pumping wells, is about 0.10. For the entire thickness of water-bearing material, 0.15 is considered to be a better estimate of the average coefficient of storage.

The "safe yield" of the alluvium and terrace deposits is estimated tentatively to be about 270 acre-ft. per sq. mi. per year. This is based on an estimated recharge of 17.5% of the normal rainfall in the area and on the assumption that all of the natural ground-water discharge could be salvaged. The major part of the natural discharge is due to transpiration by phreatophytes. These phreatophytes, which consist principally of willows and cottonwoods along the North Canadian River, discharge about 40,000 acre-ft. per year - 6 times the present discharge by pumping. The amount of water stored in the alluvium and terrace deposits of the North Canadian River averages about 3,000 acre-ft. per sq. mi.--Auth.

2-1577. LeGrand, Harry E. GEOLOGY AND GROUND-WATER RESOURCES OF PITTSYLVANIA AND HALIFAX COUNTIES, VIRGINIA: Division of Mineral Resources, Bull. 75, 86 p., 6 figs., pl., 11 tables, 1960, 20 refs.

Ground water in Pittsylvania and Halifax counties is utilized by about two-thirds of the population. It is used in all rural areas, some industrial areas, and the town of Halifax. Surface water is used by the municipalities of Chatham, Danville, Gretna, and South Boston.

These counties are located in the S.-central part of Virginia and are in the Piedmont province. Topography within these counties consists of low, rounded hills with gentle slopes and a few isolated ridges. Three rivers in the counties, the Roanoke River to the N., the Bannister River in the middle, and the Dan River to the S., flow eastward in channels that were cut more than 100 ft. below the upland area. These rivers receive drainage from a close network of tributary streams. Parts of all streams flow directly on bedrock, and other parts flow over a few feet thickness of channel sand. Flood-plain deposits that contain an upper zone of clay and a lower zone of sand and gravel, occur as bordering parts of all streams. Bedrock is exposed on many steep slopes



## MINERAL DEPOSITS

acent to valley floors. A residuum that is combed of surface soil and as much as 60 ft. of soft weathered rock covers much of the upland area.

The geology of the counties is described, and the structural and topographic characteristics of the rock are shown to be important factors that govern the yield of individual wells. The water generally occurs in soft weathered rock, a few feet above the hard fresh rock, much of which is fractured. Small amounts of water are obtained from dug and bored wells in soft weathered rock. Adequate domestic supplies are obtained in most cases from drilled wells in fractured hard fresh rock. Drilled wells range in yield from less than 1 gallon per minute to more than 100 gallons per minute; they range in depth from about 60 to 500 ft. The amount of water available from a particular well is correlated with the surrounding topography. The average yield of wells located in draws is several times that of wells on hills and more than that of wells in other topographic locations. More than 90% of the wells are drilled in hilly, upland areas where conditions are unfavorable for large supplies of water.

The withdrawal of water from wells is only a fraction of that available for recharging the underground reservoir. Recharge, derived from about 1 in. of rainfall annually, occurs in the upland areas; discharge occurs mainly in adjacent lowlands. The annual recharge and discharge are in balance, and therefore there is no increase or decrease in the annual trend in the fluctuation of the water table. Good quality water is obtained from fracture zones within schists and gneisses. Water from only a few wells contains objectionable amounts of the Fe impurities. Water from wells in Triassic sediments is hard in many cases.--Auth.

1578. Berry, Delmar W. GEOLOGY AND GROUND-WATER RESOURCES IN THE RAWLINS AREA, CARBON COUNTY, WYOMING: U.S. Geol. Survey, Water-Supply Paper 1458, 74 p., 3 maps (col. geol. map in pocket, scale 1:63,360), diag., graphs, table, 1960, 17 refs.

The Rawlins area in W.-central Carbon County, W.-central Wyoming, includes approximately 634 sq. mi. of plains and valleys grading into relatively rugged uplifts. The climate is characterized by low precipitation, rapid evaporation, and a wide range in temperature. Railroad and ranching are the principal occupations in the area.

The exposed rocks in the area range in age from Precambrian through Recent. The older formations are exposed in the uplifted parts, the oldest being deposited along the apex of the Rawlins uplift. The formations dip sharply away from the anticlines and other uplifts and occur in the subsurface throughout the remainder of the area.

The Cambrian rocks (undifferentiated), Madison

limestone, Tensleep sandstone, Sundance formation, Cloverly formation, Frontier formation, and Miocene and Pliocene rocks (undifferentiated) yield water to domestic and stock wells in the area. In the vicinity of the Rawlins uplift, the rocks of Cambrian age, Madison limestone, and Tensleep sandstone, yield water to a few public-supply wells. The Cloverly formation yields water to public-supply wells in the Miller Hill and Sage Creek basin area. Wells that tap the Madison limestone, Tensleep sandstone, and Cloverly formation yield water under sufficient artesian pressure to flow at the land surface. The Browns Park formation yields water to springs that supply most of the Rawlins city water and supply water for domestic and stock use.

Included on the geologic map are location of wells and test wells, depths to water below land surface, and location of springs. Depths of water range from zero in the unconsolidated deposits along the valley of Sugar Creek at the southern end of the Rawlins uplift to as much as 129 ft. below the land surface in the Tertiary sedimentary rocks along the Continental Divide in the southern part of the area.

The aquifers are recharged principally by precipitation, by percolation from streams and ponds, and by movement of ground water from adjacent areas. Water is discharged from the ground-water reservoirs by evaporation and transpiration, by seeps and springs, through wells, and by underflow out of the area.

Although most water supplies in the area are obtained from springs, some domestic, stock, and public supplies are obtained from drilled wells, many yielding water under artesian pressure, and some flowing.

Dissolved solids in the water from several geologic sources, ranging from 181 to 6,660 parts per million (p.p.m.), indicate the varied chemical quality of ground water in the Rawlins area.

Water from the Cambrian rocks, Tensleep sandstone, Cloverly formation, Frontier formation, Browns Park formation, and Miocene and Pliocene rocks is generally suitable for domestic and stock use. However, water yielded to the only well sampled in the lower part of the Frontier formation contained a high concentration of fluoride. Water from the rocks mentioned above contains less than 1,000 p.p.m. of dissolved solids but in some places may contain Fe in troublesome amounts. Water from the Madison limestone and Tensleep sandstone combined, Permian rocks, and Sundance formation contains more than 1,000 p.p.m. of dissolved solids. Water in the Sundance, Cloverly, and Frontier formations is very soft.

More ground water can be obtained in the Rawlins area than is now being used. Many springs are undeveloped, and water can be obtained from additional wells without unduly lowering ground-water levels.--Auth.

## 12. MINERAL DEPOSITS

Also: Geophysics 2-1486; Fuels 2-1596.

1579. Williamson, D.R., and Lorraine Burgin. SOURCES OF INFORMATION FOR WESTERN STATES MINERAL INDUSTRIES: Colorado School Mines, Mineral Industries Bull., v. 3, no. 2, p. 1-10, March 1960.

A reconnaissance guide to the literature and other sources of information which aid working with raw

materials in the western United States. Emphasis is placed on including the most productive and complete sources of information, although these might merely be additional directories with no factual data. The following categories are included: indexes and guides to the literature; selected magazines, trade publications, and bulletins; directories; selected glossaries, handbooks, and texts; general state publications; maps; federal government agencies;

state government agencies; professional geologic societies; mining associations; sources of out of print books; books reprinted; public and government libraries; universities, colleges, and associated libraries; bibliographic centers; translations; drill core libraries.--M. Russell.

2-1580. Smellie, D.W. TECHNOLOGY - MINING EXPLORATION (ANNUAL REVIEW): Can. Mining Jour., v. 81, no. 2, p. 166-168, 2 illus., Feb. 1960, 33 refs.

Canadian exploration developments during 1959 are reviewed. The Mattagami Lake area and extensions accounted for a good proportion of exploration activities.

Mention is made and references cited in regard to developments in mining geophysics and geochemical prospecting.

The status of exploration geology is commented upon. References are cited to significant publications on the genesis of massive sulfides, on the environment of sedimentary Fe deposits, and on the origin of barren iron sulfide horizons in the Canadian Shield.--W. C. Peters.

2-1581. McKelvey, Vincent E. RELATION OF RESERVES OF THE ELEMENTS TO THEIR CRUSTAL ABUNDANCE: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 234-241, 2 diags., table, 1960, 12 refs.

The tonnage of minable reserves in short tons (R) for many elements in the United States is equal to crustal abundance in percent (A) times  $10^9$  to  $10^{10}$ , and all 26 elements for which data are available have ratios in the range  $R = A \times 10^9$  to  $10^{11}$ . The linear relation that appears to prevail between reserves and abundance for all except the most abundant elements is useful in forecasting reserves in large segments of the earth's crust or over the world at large. For purposes of estimating world reserves of unexplored elements a figure of  $A \times 10^{10}$  to  $10^{11}$  probably will give the right order of magnitude.--Auth.

2-1582. Ginzburg, I.I. PRINCIPLES OF GEO-CHEMICAL PROSPECTING. TECHNIQUES OF PROSPECTING FOR NON-FERROUS ORES AND RARE METALS. Translated from the original Russian by V.P. Sokoloff: 311 p., 72 figs. incl. maps, secs., diags., graphs, 28 tables, New York, Pergamon Press, 1960, approx. 300 refs.

The wide application of geochemical prospecting methods at all stages of the geologic service raises the need for developing some theoretical premises of such prospecting, as well as to the need for a summation of the bulky data accumulated empirically by geologic reconnaissance, geologic prospecting, and geophysical field parties.

The author has endeavored to throw as much light as possible on problems originating in geochemical prospecting, to evaluate the relative importance of different prospecting methods, and to give a generalized view of prospecting work in different geologic, pedologic, climatic, and orographic environments, as well as to approach a clarification and an explanation of certain regularities which could serve as the basis of a rational orientation of geochemical prospecting.--From foreword.

Contents:

1. Methods of Analytical Investigations, p. 1-11.
2. Geochemical Tracers (Indicators), p. 12-25.

3. Accumulation of Metals in Igneous and Metamorphic Rocks, p. 26-55.

4. Accumulation of Metals in Sedimentary Rock, p. 56-73.

5. Prospecting in Bedrocks, p. 74-90.

6. Prospecting for Deposits without Surface Outcrops, p. 91-109.

7. Accumulation of Metals in Unconsolidated Overburden (Dispersion Halos), p. 110-162.

8. Prospecting at the Surface of the Overlying Mantle, p. 163-185.

9. Migration of Metals in Waters, p. 186-215.

10. Hydro-geochemical Prospecting for Metals and Characteristics of Different Water Types Associated with Ore Deposits, p. 216-229.

11. The Bio-geochemical Method of Prospecting, p. 230-242.

12. The Geobotanical Method of Prospecting, p. 243-251.

13. General Conclusions in Reference to Geochemical Survey, p. 252-266.

2-1583. Bichan, W. James. THE ORIGINS OF THE MASSIVE SULPHIDES, PART I: Can. Mining Jour., v. 81, no. 3, p. 73-78, 2 illus., March 1960.

It is suggested that massive sulfide bodies are specialized examples of igneous rocks rather than deposits from dilute hydrothermal fluids. Massive sulfide ore bodies are presented as resultants of cumulative concentration from multiple diastrophic and diagenetic cycles, the final stage being rapid crystallization from a magma.

Reference is made to massive sulfide deposits in the Rio Tinto mining field of Spain. The genesis of the Spanish deposits is ascribed to magmatic segregation and flotation concentration during which bubbles of escaping volatiles carried fluid metallic sulfides toward the margins of the magma chamber, resulting in a fluid pyritic matte which crystallized rapidly in the usual paragenetic sequence. Some of the sulfides trapped within the increasingly viscous magma are considered to form disseminated deposits. Certain ore bodies are designated as results of penetration of roof rock by fluid sulfides.

Structural and textural characteristics are cited in support of the suggested method of magmatic segregation.

The proposal and supporting observations are expanded to include massive sulfide bodies at Tilt Cove, Newfoundland; Noranda, Quebec; Buchans, Newfoundland; and Bathurst, New Brunswick.--W. C. Peters.

2-1584. Swayne, William H., and Frank Trask. GEOLOGY OF EL SALVADOR: Mining Engineering, v. 12, no. 4, p. 344-348, 2 illus., map, sec., Apr. 1960.

El Salvador is a porphyry Cu deposit on the trend of the Great Copper Belt of the Chilean Andes in Atacama province. The area's oldest rocks are Cretaceous andesitic bedded volcanics. These have been folded, faulted, and intruded by Tertiary rhyolites. El Salvador is the southernmost and largest of several centers of better-grade Cu mineralization occurring in a mineralizing belt 3 mi. long and a mile wide. Primary sulfide mineralization, in which chalcopyrite is the principal ore, was followed by secondary enrichment resulting in the formation of a chalcocite ore blanket covering an area 1,700 by 4,000 ft. and up to 950 ft. thick. It is under a leach cap of from 180 to 1,500 ft. thickness.--M. Russell.



1585. Mills, Joseph W. GEOLOGIC SETTING OF THE NICKEL OCCURRENCES ON JUMBO MOUNTAIN, WASHINGTON: Mining Engineering, 12, no. 3, p. 272-274, illus., map, March 1960, 1 fig.

Discovery of Ni on Jumbo Mountain, Snohomish County, focused attention on the area in 1956. The paper covers the geology of the area, touching upon the following structures: sedimentary rocks, ultrabasic and basic igneous rocks, gabbro, acidic igneous intrusives, and faults. The geology of the Ni deposits is discussed.--Auth.

1586. Govorov, I. N. FEATURES OF THE MINERALOGY AND GENESIS OF THE TIN-BERYLLIUM-FLUORITE DEPOSITS OF THE FAR EAST: Izv. Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, p. 47-55, table, pub. 1959, 12 refs.

The features of a Sn-Be-fluorite mineralization are discussed. The deposits are significantly different from the deposits of Sn, Be, and fluorite described in the literature. Mica-fluorite ores and spaz greisens occur at the contacts of granites with ore bodies, and, therefore, the mica-fluorite rocks are unusual desilicated greisens formed from the replacement of limestones by fluorine solutions or in silica.--Auth.

1587. Lukin, L.I., and E.P. Sonyushkin. STRUCTURES OF HYDROTHERMAL URANIUM DEPOSITS AND SOME PROBLEMS OF THEIR STUDY: Izv. Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 1-11, 12 figs. incl. secs., diags., table, 1959.

In this article there are discussed the problems of the relation between hydrothermal U deposits and ore bodies as related to tectonic structures. The main structural types of the deposits are briefly characterized. The part played by the folded and faulted dislocations in the localization of U minerals is outlined.--Auth.

1588. Swanson, Vernon E. OIL YIELD AND URANIUM CONTENT OF BLACK SHALES: U.S. Geological Survey, Prof. Paper 356-A, 44 p., 3 maps, 12 figs., 27 diags., 2 graphs, table, 1960, 83 refs.

Some black shales contain as much as 100 times more U than other common sedimentary rocks, and they also contain organic matter that will yield oil when subjected to destructive distillation. Such shales may be referred to as uraniferous oil shales and have been considered as a potential source of both oil and U; oil yield and U determinations on more than 500 samples of these shales are recorded in this report.

Slightly more than half of these samples are from the Late Devonian Chattanooga shale and its partial equivalents in the eastern and midcontinent areas of the United States. In central Tennessee, the upper member of the Chattanooga shale is about 15 ft. thick, contains 0.006% U, and will yield about 10 gallons of oil per ton of shale. Limited data indicate that the Chattanooga shale in Alabama and southern Kentucky, the Antrim shale of Michigan, the New Albany shale of Indiana and northwestern Kentucky, and its stratigraphic equivalent in southern Illinois, and the Chattanooga and Woodford shales of the midcontinent area have slightly lesser quantities of oil and U. A channel sample of 5 ft. of the Houborn shale member of the Houy formation in

central Texas indicates that this unit has a U content of 0.009% and an oil yield of 21.8 gallons of oil per ton of shale.

Some of the marine black shales in the cyclothems of Pennsylvanian age in Illinois, Kansas, and Oklahoma contain between 0.004 and 0.010% U and yield 8 to 15 gallons of oil per ton of shale, but generally these shales are less than 3 ft. thick. Some shale units in the Phosphoria formation of Permian age in southwestern Montana, which are about 10 ft. thick, will yield 10 to 15 gallons of oil per ton of shale, but their U content of 0.002 to 0.004% is relatively low.

The few data available indicate the Sharon Springs member of the Pierre shale of Late Cretaceous age in the Great Plains area has an oil yield of less than 8 gallons per ton of shale and a U content of about 0.003%. The Green River formation of Eocene age in Colorado and Utah has beds of oil shale tens of feet thick that will yield more than 25 gallons of oil per ton of shale, but the U content of these beds is low, generally between 0.0003 and 0.0010%.

Both oil and U have been recovered in large quantities from the Upper Cambrian black shales of Sweden, which yield about 14 gallons of oil per ton of shale and about 0.023% U. Some other oil shales from foreign sources that yield 50 or more gallons per ton generally contain about 0.0005% or less U.

A fair positive relation between oil yield and U content exists for some of these shales, particularly for parts of the Chattanooga shale locally and the Antrim shale, but in other shales little or no relation is apparent. In some of the Pennsylvanian shales and in those in the Phosphoria formation the U is more closely related to the phosphate content.

Whereas the oil from these shales is inherent to and derived directly from the organic matter, most of the U is attached to or precipitated in the presence of organic matter just before or during the time of deposition of the organic-rich sediment. It is suggested that 2 types of organic matter should be distinguished, the sapropelic type derived principally from algae, pollen and spores, resins, and the fatty tissues of animals, and the humic type which is derived principally from cellulose and lignin or the woody parts of plants. The sapropelic type of organic matter generally yields 4 or 5 times more oil than the humic type, but, because of its general resistance to decay, is probably insignificant in the process of concentrating U. The humic type of organic matter, either in its solid form or as soluble humic acid extracts, or, indirectly, as it creates a reducing and acidic environment during its decay, is believed responsible for the precipitation or sorption of the U in black shales. Only where the proportion of sapropelic to humic type of organic matter remains the same in an otherwise homogeneous black shale will the oil yield and U content have a high positive correlation.--Auth.

2-1589. Dzhezdzalov, A.T. THE PATTERN OF DISTRIBUTION OF IRON ORE DEPOSITS IN THE SAKSAGANIAN REGION OF KRIVROY ROG: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 53-72, 8 figs. incl. map, plans, table, pub. 1959, 7 refs.

The conditions of formation of the ancient weathered zone in the Saksaganian belt of Krivoy Rog are elucidated. A basis has been found for correlating the rich ore beds with fracture zones. The general pattern of weathering of the ferruginous quartzite and pattern of ore formation are examined.--Auth.

2-1590. Shabynin, L.I. ON THE GENESIS OF IRON ORE DEPOSITS IN SOUTH YAKUTIA: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 1, p. 32-46, 5 illus., sec., pub. 1959, 32 refs.

The main features of the geological structure and composition of the rocks and ores of the southern Yakutian Fe ore deposits, including complex B-Fe ores, are described. Existing opinions on conditions of formation of these deposits are critically examined. The arguments of the adherents of a sedimentary-metamorphic origin of Fe concentration are refuted, and the ores are shown to be skarns.--Auth.

2-1591. Laurence, Robert A. GEOLOGIC PROBLEMS IN THE SWEETWATER BARITE DISTRICT, TENNESSEE: Am. Jour. Sci., v. 258-A (Bradley Volume), p. 170-179, sec., 2 tables, 1960, 36 refs.

Residual barite deposits, derived from vein and breccia fillings in carbonate rocks of Ordovician age have been mined extensively in 3 parallel strike belts near Sweetwater, Tennessee. Geologic problems germane to the further exploitation of this mineral resource include the origin and age of the primary deposits and the breccias which contain them; relation of the barite deposits to Zn deposits of E. Tennessee; weathering of the carbonate rocks, including accumulation of the residual mantle; solution, transportation, and deposition of barite, fluorite, and silica; and the extent of ancient terrace deposits. The district presents a promising field for coordinated research in economic geology, geochemistry, geophysics, and geomorphology.--Auth.

2-1592. Castellani, Farrell. UNDERGROUND MINING FOR CEMENT ROCK: Explosives Engineer, v. 38, no. 1, p. 21-24, 6 illus., Jan.-Feb. 1960.

The Missouri Portland Cement Company's new cement plant is in the Kansas City area at the confluence of the Kansas and Missouri rivers. First operated in 1906 as an open pit production, it changed to underground mining in 1910. The present plant is a several million dollar - 2-kiln operation.

The rock mined is 36 ft. thick. The top 10 ft. is the bottom of the Winterset formation [Pennsylvanian]; the middle is a shale and "peanut rock," the Galesburg shale [Pennsylvanian]; and the bottom is Bethany Falls limestone [Pennsylvanian]. The best grades of portland cement require Fe, alumina, and silica in addition to pure limestone. These are found in the argillaceous beds. The operation at this property is unique in that this 36-ft. thickness of material can be blended into a chemically correct and balanced material for making cement. Mining methods are described.--J. W. Skehan.

2-1593. Gooch, Edwin O., Robert S. Wood, and

William T. Parrott. SOURCES OF AGGREGATE USED IN VIRGINIA HIGHWAY CONSTRUCTION: Virginia, Div. Mineral Resources, Mineral Resources Rept. 1, 65 p., 19 pls. incl. map, 1960, 15 refs.

Materials used for highway aggregate in Virginia range from highly metamorphosed rocks of Precambrian age to unconsolidated sand and gravel deposits of Recent age. Limestone, dolomite, and quartzite are produced in the Ridge and Valley province. Granite, various types of gneiss, diabase, basalt, marble, limestone, sandstone, and conglomerate are produced in the Blue Ridge and Piedmont provinces. Sand and gravel are produced in the Coastal Plain province.

Aggregate is classified into grades A, B, and C on the basis of 3 physical properties: abrasion loss, specific gravity, and absorption. Tests used in determining the physical properties are described in the text and the specifications for each grade are listed.

The name and location of each aggregate producer is given in the text and on the accompanying map. The geology of each pit and quarry is described, physical test data on the aggregate are listed with each description, and the size of the operation or rate of production during 1957 or 1958 is given.--Auth.

2-1594. Grandone, Peter, and William E. Ham. THE MINERAL INDUSTRIES OF OKLAHOMA IN 1959: Oklahoma Geology Notes, v. 20, no. 3, p. 46-49, table, March 1960.

Total value of 1959 mineral production in Oklahoma is estimated at \$742 million, 2.6% less than the 1958 value and 8.3% less than the 1957 record value of \$809 million. Mineral fuels accounted for nearly 95% of the 1959 value, nonmetals for 5%, and metals less than 1%. Ten of the 17 minerals produced showed a gain in 1959 value over that of 1958. Helium was produced in the State during 1959 for the first time.

The major commodity is petroleum, of which 197 million barrels was produced. For the fourteenth consecutive year Oklahoma was the fourth largest producer of petroleum in the United States.--W. E. Ham.

2-1595. Pistsov, Yu. P. ON THE AGE OF THE NERTCHINSK-ZAVOD GROUP OF POLYMETALLIC ORE-DEPOSITS, EAST TRANS-BAIKAL REGION: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 3, p. 100-101, pub. 1959.

The polymetallic ore deposits of the Nertchinsk-Zavod group are post-Jurassic to pre-Lower Cretaceous.--M. Russell.

### 13. FUELS

See also: Geologic Maps 2-1346; Stratigraphy 2-1397, 2-1403, 2-1406, 2-1409; Geophysics 2-1487, 2-1488, 2-1502, 2-1503, 2-1504; Mineral Deposits 2-1588, 2-1594; Engineering Geology 2-1612.

2-1596. Anderson, James A., 3d. PERCENTAGE DEPLETION ALLOWANCE (In: McGrain, Preston, and Thomas J. Crawford, eds. Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959: Ken-

tucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 40-47, 1959).

The concept of percentage depletion evolved over an extended period of time as the unique answer to a unique problem. That problem was to determine how mineral producers should compute a deduction from gross income which would fairly reflect the gradual exhaustion of their underground capital assets. The Congress recognized as early as 1918



at the conventional method of depreciating physical assets - on the basis of their cost - was inappropriate, since mineral reserves could not be made or purchased at standard costs but had to be discovered. Therefore, legislation was enacted which based the depletion allowance on the value of the mineral in place at the time of discovery. This idea had theoretical merit but "discoveries" had to be adjudged and "values" had to be estimated, with the result that depletion rates, as between producers of the same mineral, varied for unpredictable and imprecise reasons. It was decided, therefore, that the idea behind discovery depletion should be translated into more practical terms and that the depletion allowance should be computed as a fixed percentage of the price received for the mineral at the time or well.

The law which provides for the computation of depletion by this method is now available to the producers of virtually all minerals - from antimony to Zn. Thus, the oil and gas industry has not been singled out for favoritism. Nor is the rate for oil and gas inordinately high. If a barrel of oil sells for \$3.00, the depletion deduction computed is 27 1/2% of that price is \$.82 1/2%. Certainly this amount will not normally exceed the in-place value of that barrel - of the unit of capital being sold. -- Auth.

21597. Hoffmeister, William S. **PALYNOLOGY: AN IMPORTANT ROLE IN OIL EXPLORATION:** *World Oil*, v. 150, no. 5, p. 101-104, port., chart, diag., Apr. 1960.

Age determinations, correlations, and guidance in a favorable environment are now possible through the study of microfossils which include 7 groups. These groups as a whole cover geologic time from Precambrian to Recent and in all types of environment, thus providing very desirable information. The study of microforams, spores, and pollen, is supplemented with the hystrichospherids and planktonic, foraminous, and coccolith bodies. In correlation uses, the appearance of the same forms, diagnostic first occurrences, or relative percentages is very helpful. Spores and pollen (Cambrian to Recent) abundance generally decreases away from shore, while the hystrichospherids (Precambrian to Recent) generally indicate brackish water, and microforams (Jurassic to Recent) are mostly of marine environment. The yet little used microfossil dinoflagellates (planktonic) are found world-wide, and the chitinozoans are particularly helpful in Paleozoic rocks which are otherwise poor in microfossils. Extraction techniques use centrifuge of material after treatment with hydrofluoric acid, sodium hypochlorite, and ammonium hydroxide plus heavy liquid separation. The studies began with spores in England in 1833. Since 1938 there has been rapidly increasing use, particularly in Sweden and Russia, with much application to oil exploration in recent years in the United States. --K. M. Willson.

21598. Pirson, Sylvain J. **HOW TO MAKE GEOCHEMICAL EXPLORATION SUCCEED:** *World Oil*, v. 150, no. 5, p. 93-96, port., diag., Apr. 1960, 26 refs.

The main concept of geochemical exploration is that hydrocarbon gases diffuse to the surface of the earth through fissures, fractures, and faults. Leakage patterns may be related or unrelated to structural deformations. Amounts of gas dissolved in waters increase with depth, but the waters are gen-

erally undersaturated. A new concept indicates that anomalies represent leakage which took place at the time the petroleum accumulation was being formed. Formation water moving around petroleum accumulations leaves an erratic zonal distribution of minerals as shown by radioactivity, mineral alteration products, and variation in the dielectric constants. Combination of various geochemical data with diverse geoscience information is necessary for correct solutions of expectedly erratic peak and halo patterns. --K. M. Willson.

2-1599. Brake, J. A. **UNDERGROUND STORAGE OF LIQUEFIED PETROLEUM GAS** (In: McGrain, Preston, and Thomas J. Crawford, eds. *Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959: Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 7-12, 1959*).

The fluctuating demand for its products requires that the hydrocarbon industry provide bulk storage for large quantities of liquid materials. A study conducted by Columbia Hydrocarbon Corporation showed that underground storage in caverns is the most economical type of bulk storage for liquefied petroleum gas and that the inconveniences associated with it are relatively insignificant.

This paper presents a step-by-step account of problems involved in locating a site for, constructing, and final testing of an underground cavern for bulk storage of liquefied petroleum gas. -- Auth.

2-1600. Jordan, Louise. **WOODWARD COUNTY'S GAS FIELDS:** *Oklahoma Geology Notes*, v. 20, no. 4, p. 83-89, 3 maps, 2 tables, Apr. 1960, 2 refs.

Natural gas in commercial quantity was discovered in 1956 in Woodward County, northwestern Oklahoma. At the end of 1959, a total of 67 tests has been made since 1930 in the county area and of these 23, or 34% of the total, are productive of gas (one of oil). Of the 8 gas-condensate productive areas, 3 new fields were found in 1959. Production is from Missourian sandstones (Musselleum or Cottage Grove and Tonkawa) and Morrowan sandstones of Pennsylvanian age and from limestones of Mississippian age. Although shows of hydrocarbons have been noted in older rocks (Hunton, Viola, and Simpson) in some of the few (7) wells penetrating Middle Ordovician, no commercial production has been developed. -- Auth.

2-1601. Roberts, Carl H. **ORISKANY FOUND IN PENNSYLVANIA SYNCLINE:** *Oil & Gas Jour.* v. 58, no. 16, p. 174, 178, map, Apr. 18, 1960.

A thick accumulation of Oriskany [Devonian] sand has been found in a synclinal position between 2 anticlines which have at best a thin impermeable veneer of this formation on their crests. Commercial gas production has been developed from this synclinal accumulation. --N. Street.

2-1602. **TODAY'S ACTIVE OIL FRONTS:** *Oil & Gas Jour.* v. 58, no. 17, p. 148-199, 2 illus., 11 maps, 3 secs., 10 logs, tables, Apr. 25, 1960.

A review of the world's most active and successful areas. It includes the following: Panhandle-Hugoton region covering Oklahoma and Texas panhandles, SW. Kansas, SE. Colorado, and NE. New Mexico; Trenton-Ordovician in Michigan; Oriskany-Devonian and lower rocks in West Virginia; Ohio; Pennsylvania; Kansas-Nebraska; the Lower Cretaceous-Edwards

in S. and SW. Texas and gas in deeper Wilcox in S. Texas; Mississippi (Cretaceous and beyond); the Green River basin, Wyoming; Alaska. In Canada the activity in the Swan Hills region in Alberta (the most active area on the continent) is reviewed, together with that in British Columbia, and in the Northwest Territories and Yukon. In France, the activity is in the Paris and Aquitaine basins.

The tables show producing formations, drilling details, etc. The text maps show the locations of the various areas of interest; sample electric logs are included for many of the areas.--N. Street.

2-1603. Ray, Edward O. RESUME OF DRILLING ACTIVITIES IN EASTERN KENTUCKY IN 1958 (In: McGrain, Preston, and Thomas J. Crawford, eds. Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959; Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 48-50, 1959).

During 1958, a total of 439 wells were drilled in the eastern Kentucky field: 185 gas wells, 140 oil wells, and 114 dry holes. Oil production increased due to intensified drilling activity in Breathitt County; however, a slight decline in gas production was experienced. Eastern Kentucky produced 3,671,854 barrels of oil and 67,886,867 MCF of gas during 1958. The deepest test during the year was United Fuel Gas Company's S.W. Williams no. 8613T, in Breathitt County, which was abandoned in the basement complex at a total depth of 11,130 ft.--Auth.

2-1604. Kentucky Geological Survey, comp. OIL PRODUCTION IN KENTUCKY FOR YEAR 1958 (In: McGrain, Preston, and Thomas J. Crawford, eds. Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959; Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 82, 1959).

An alphabetical listing of counties with the number of barrels of oil produced in each in 1958.

2-1605. Lytle, William S. HISTORY, PRESENT STATUS AND FUTURE POSSIBILITIES OF SECONDARY RECOVERY OPERATIONS IN PENNSYLVANIA: Pennsylvania Geol. Survey, Bull. M-41, [15 p.], 3 maps, chart, 3 graphs, 1960, 33 refs.; reprinted from: Interstate Oil Compact Commission Committee, Bull., v. 1, no. 2, p. 28-42, Dec. 1959.

Pennsylvania has had many firsts in the history of crude oil production. This state led the nation in the volume of crude oil produced from the discovery of the Drake Well in 1859 to 1895, and, though the quantity of production in the state has diminished since then, Pennsylvania still produces a preferred lubricating oil used throughout the world.

Vacuum was used in Pennsylvania as early as 1869 to increase the recovery of crude oil from the oil reservoir. The use of vacuum in the state was at its height from 1918 to 1925 and this technique is still being practiced today.

In 1880, John F. Carll first called attention to the

possibilities of employing water flooding to increase the recovery of oil from the oil sands. By the early 1890's, there was some intentional, but secret, water flooding in the Bradford field. Flooding developed from the early "dump" flood used at the turn of the century to the intensive "five-spot" introduced in 1924 and in use today. The Bradford region has gained recognition today as the greatest secondary recovery laboratory in the world. Many new improvements in oil recovery techniques came out of research in this region.

In 1890-1891 the first recorded attempt to obtain additional oil by the injection of gas was made in Venango County. The use of air or gas as a driving medium to increase oil production continued to increase and by the middle 1920's was widespread in the state. Production resulting from air or gas injection reached its peak in 1929. This method is used extensively today.

Many important phases of petroleum technology had their beginnings in Pennsylvania. Early experiments in determining porosity and permeability were run in the state. The first core laboratory, the first flood water laboratory, the first diamond core of an oil-bearing sand, the first cable tool core, and the first chip core, plus many others were introduced in Pennsylvania.

As of Jan. 1, 1959, Pennsylvania's oil fields had produced 1,219,070,000 barrels of crude oil of which 327,662,000 barrels were produced by water flooding and 19,496,000 barrels by air or gas drive. The proven reserves as of the same date were 120,018,000 barrels. There are thousands of acres in the Pennsylvania oil fields including the 20-mi. long Foxburg-Clarion field which have not been subjected to secondary recovery methods.--Auth.

2-1606. Agnew, Allen F. OIL TESTS IN BLACK HILLS FRINGE, SOUTH DAKOTA: South Dakota Acad. Sci., Proc., v. 38 (South Dakota, State Univ. Bull., ser. 60, no. 2), p. 60-65, 3 maps, table, 1959, pub. 1960.

In the past 50 years, 191 tests have been drilled for oil in the Black Hills area of South Dakota, and 80% of this number have been drilled in the last 15 years. Of the 175 holes for which information is available, approximately 60% tested only as deep as the Cretaceous Newcastle or Fall River sands. An additional 25% reached the Pennsylvanian Converse or Leo sands, but only 3% penetrated all the sedimentary rocks to the Precambrian basement.

The reasons for not testing deeper include 1) lack of money, 2) narrowness of objective, 3) prejudice, and 4) promotional aspects.

The oil possibilities of the Black Hills area deserve a reappraisal.--Auth.

2-1607. Yaskovich, B.V. NEW DATA ON THE OCCURRENCE OF BITUMENS IN THE CAMBRIAN ROCKS OF SOUTHERN FERAGANA: Akad. Nauk SSSR, Bull., Geol. Ser., in translation, 1958, no. 2, p. 99-101, 2 tables, pub. 1959, 2 refs.

English translation of GeoScience Abstracts 1-534.



## 14. ENGINEERING GEOLOGY

1608. Russell, Harold E. **DRIVING THE JAY-ARD: Explosives Engineering**, v. 38, no. 1, p. 7-12, illus., chart, Jan.-Feb. 1960.

Jaybird tunnel, a 14-ft.-diameter horseshoe-shaped bore, 21,300 ft. long, is a unit of the \$85 million hydroelectric power development of the Upper American River Project located 75 mi. E. of Sacramento, California, just N. of Riverton on Route 50. The project will utilize run-off from an area of approximately 250 sq. mi. on the western slopes of the Sierra Nevada. The completed project will include 10 dams, 4 powerhouses, and 7 tunnels with combined length of 82,400 ft.

Jaybird tunnel will have a grade of 0.0025 with the ore being driven through the Calaveras formation. The Calaveras consists of folded schists, phyllites, gneisses, quartzites, and hornfels. Andesitic agglomerate mudflow-type material in tuffaceous matrix overlies the metamorphics along part of the alignment.

Drilling procedure, blasting technique, and the mucking operation are described in detail. The rock breaks well under the procedure used, and a minimum of secondary blasting is required. In sound rock requiring no steel support, progress has averaged 42 to 48 ft. per day per heading, requiring 5 to 10 shots each shift.--J. W. Skehan.

1609. California, Dept. of Water Resources. **INVESTIGATION OF ALTERNATIVE AQUEDUCT SYSTEMS TO SERVE SOUTHERN CALIFORNIA. APPENDIX C. PROCEDURE FOR ESTIMATING COSTS OF TUNNEL CONSTRUCTION**, prepared by John W. Marlette, Ernest M. Weber, and others: 1. Bull. 78 (Appendix C), 74 p., 10 illus., 19 fold. 2. Incl. map, secs., graphs, 11 tables, Sept. 1959, 5 refs.

During the period 1956-1959, the California Dept. of Water Resources, in connection with the Feather River and Delta Diversion Projects, carried out an investigation of alternative aqueduct systems. Results were published in Bulletin no. 78. A large number of aqueduct routes leading from the San Joaquin Valley into southern California were considered. Because of the nature of the terrain, all economical possibilities for such routes would require substantial tunnel construction, for which cost estimates were needed.--From introd.

This report, published as Appendix C. of Bulletin 78, sets forth a rapid, standardized method of estimating tunnel construction costs, utilizing readily available geologic data, to assist geologists and engineers in determination of feasible routes. Data for this method was obtained through review of construction case histories of 100 modern tunnel projects, inspection of tunnels under construction, and interviews with contractors. Curves are presented relating estimated rates of heading advance, for varying diameter tunnels, tunnel excavation costs, and timber lagging for 8 possible physical rock conditions. Rock conditions depicted include: intact rock; stratified or schistose rock; massive, moderately jointed rock; moderately blocky and seamy rock; very blocky and seamy rock; completely crushed or unconsolidated rock; wet competent rock; and wet crushed or unconsolidated rock.

Method involves preparation of a geologic section based on geologic mapping, geophysical surveys and drilling, or driving exploration tunnels or shafts. Rock units along tunnel alignment are classified as to expected physical condition, and estimates are prepared for rock density, estimated

rock load, rib spacing and maximum depth of cover. Rates of heading advance, tunnel excavation costs, and timber lagging requirements are obtained from the appropriate cost curve for each physical rock type appearing on the geologic section.--R. C. Richter.

2-1610. Poland, George F. **ANCHORING A TUNNEL IN SAND: Civil Engineering**, v. 30, no. 3, p. 59-61, 2 illus., 2 diag., March 1960.

Thorington Construction Company encountered the problem of anchoring the bottom of a vehicular traffic tunnel being constructed under the New River in Fort Lauderdale, Florida. Water level in the area is about 2 ft. below ground surface. The geologic materials are alternating layers of sand and porous limestone. The tunnel bottom (which is a 4 ft. thick tremie seal) had to be anchored because its weight was less than the hydrostatic force below it. The original plan was to anchor the tremie seal with reinforcing bars grouted in the underlying rock. This plan was abandoned because 42% of test bars pulled out of the rock.

It is known that a friction pile distributes its load in the form of a bulb. It was reasoned that the form of the "negative" load of a tension pile would be similar, but with the forces in opposite directions. Using the principle of overlapping pressure bulbs, a volume of material was trapped by anchor piles sufficient to overcome the hydrostatic force. The spacing of the piles was such that the weight of material did not exceed the design load of the piles. The piles were set by 1) driving a casing with an expandable cap to the plane of anchorage, 2) filling the casing with grout, 3) setting a number 11 reinforcing bar in the grout, 4) withdrawing the casing, and 5) fastening the bar to the tremie seal.

The novel system was proved successful when the construction cofferdam was dewatered and the tremie seal remained in place.--R. E. Pendergast, Jr.

2-1611. Abel, John F., Jr. **PERMAFROST TUNNEL: Mines Mag.**, v. 50, no. 3, p. 12-17, 5 illus., map, 3 diag., 4 tables, March 1960, 5 refs.

A report on permafrost mining techniques, made under the direction and control of U.S. Army, Snow, Ice, and Permafrost Research Establishment, at Camp Tuto, Greenland. An experimental tunnel 300 ft. long, 10 ft. wide, and 9 ft. high was driven into a hillside of glacial till. Basic hard rock mining methods were used in portal, tunnel, and room excavation. The operation included drilling blast holes, loading, blasting, ventilating, mucking out, extending the rail, and placing protective arches. The portal required support to prevent caving of thawed and unconsolidated material, but none was required in the tunnel. The volume of material removed was 1,080 cu. yd.; 334 cu. yd. from the portal, 666 cu. yd. from the tunnel, and 80 cu. yd. from an experimental room (20 x 25 ft.) in the tunnel.

The portal required protection from rock falling from a 2-ft. active layer of unconsolidated material and material thawing out of the roof. Protective steel arches were used with lagging in the first 78 ft. of tunnel. The slot blasted into the hillside before tunneling was accomplished was 15 ft. high at the face.

The permafrost material is described as a poorly graded gravel under the unified soil classification system of the Corps of Engineers. "The till would be considered free draining and non-frost susceptible if thawed, in that no volume change would occur."

Normally no lensing would be likely to occur in the material. Ninety-five percent by weight of the material was a granite gneiss. Random fragments of hematite, epidote, olivine, and a competent silicified pink sandstone were present. The temperature in the tunnel wall ranged from  $-6.0^{\circ}\text{C}$ . to  $-10.8^{\circ}\text{C}$ . - in the same range as in an ice tunnel on the ice cap.

Two types of tunnel blasting rounds were used, a burn cut and a V-cut round. The holes were all drilled with 1-5/8 in. bits. Four types of explosives were used during blasting experimentation. Costs and the experimental blasting record are given. The drill round giving the best results was a 5-hole burn cut, with a total of 35 holes, using 50% nitroglycerine dynamite. Drilling rate records for the various bits and equipment are given. Bit life was 212 ft./bit. The average drill round used 36 blast holes, and drill bit and steel cost \$2.12/ft. of tunnel advance.

The report concludes that excavation of frozen glacial till is feasible by conventional methods, modified to allow for the temperatures. It recommends commercial percussion drilling of blast holes with anti-freeze drill fluid. Diamond drilling is recommended for clean wall holes in the frozen till. High detonation velocity explosives are recommended for proper fragmentation. --J. R. Hyland.

2-1612. Taylor, Lloyd C. NEW DEVELOPMENTS IN WELL STIMULATION TECHNIQUES (In: McGrain, Preston, and Thomas J. Crawford, eds. *Proceedings of the Technical Session, Kentucky Oil and Gas Association, Twenty-Third Annual Meeting, June 5, 1959; Kentucky Geol. Survey, Ser. 10, Spec. Pub. 2, p. 28-39, 10 diag.*, 1959).

Hydraulic fracturing is the most important well stimulation technique developed during the past 10 years. As the result of careful study of thousands of fracturing treatments already accomplished, much engineering information, useful in planning future fracturing jobs, is now available. From study of these data, it is believed that fracture planes and area of reservoir rock exposed to fracture pattern can be determined, proper choice of propping agents can be made, and reservoir penetration can be calculated, all of which will lead to more efficient and beneficial well stimulation. --Auth.

2-1613. Symposium on Geology as Applied to Highway Engineering, 10th, Atlanta, Georgia, 1959. PROCEEDINGS: 83 p., illus., 2 maps, secs., Atlanta, Georgia Institute of Technology, 1959, refs.

This meeting, held at the Georgia Institute of Technology on Feb. 20, 1959, was the tenth assembly of these symposia which had their start in Richmond, Virginia, in 1950. The 1959 symposium was under the joint sponsorship of the Georgia State Highway Dept. and the Georgia Institute of Technology. Reports are as follows:

Felix, George D. Geology as an Aid to Right of Way, p. 1-4.

When land is acquired for highway building it is commonly necessary for geologists to render advice on the worth of mineral deposits so that just compensation is given. Other questions which have been settled with geologists' advice include claims that highway vibration and blasting operations have caused or will result in damages to property. Specific examples are cited to illustrate the problems involved. --M. Russell.

Furcron, A. S. The Distribution and Character of Stone for Aggregate in Georgia, p. 5-25.

A brief but comprehensive review of the distribution of rocks suitable for aggregate, gravel, or crushed stone. The stone industry is next to the top mining industry in Georgia. Discussion is keyed to the state geologic map of 1939. The distribution and character of aggregates is shown on a map, scale 30 in. to 1 mi., in which 11 distinct lithologic types are delineated. In 8 of these, useful stone for aggregate is available.

In the Coastal Plain of Georgia, 60% of the state, the principal source of stone will be Tertiary limestone, particularly the outcropping belt of Ocala. These limestones are not as hard as present requirements for aggregate generally demand, but with judicious planning, a very large supply of limestone can be obtained. In the crystalline and highland area, 30% of the state, foliated granites and migmatites are producing most of the commercial crushed stone. Along the southern marginal belt of the Piedmont, meta-volcanic rocks are also supplying a considerable amount of aggregate. The granite rock types will remain in the future a dominant source of supply for this district. In the Paleozoic areas, all or parts of the 10 NW. counties (10% of the state), the principal rocks are the Conasauga limestone, Mississippian limestones and, to a lesser extent, Ordovician limestones. --From author introd. and concl.

Marshall, Harry E. Design Considerations in the Treatment of Soft Foundations, p. 26-47.

The paper deals with treatment of peat and associated soft clays common to glaciated areas of the U. S. Attention is given to problems encountered in the construction of the Ohio Turnpike across the northern part of the state and the Interstate Route from Columbus to Cleveland. The materials are classified on the basis of origin as organic clays, sedimentary peat, fibrous peat, and woody peat. Methods of treatment include avoidance, total excavation, partial excavation and displacement, vertical sand drains, and bridging. --M. Russell.

Upham, Charles M. The Use of Geological Investigation in Foreign Consulting Work, p. 48-56.

Geologists are extensively called upon to provide information on the best location for roads and the materials with which to build them. Geologic studies in the Nile delta of Egypt helped locate sand, gravel, and rock for crushing, saving time and haulage. Airphoto analysis, soil resistivity surveys, and seismic surveys are particularly useful tools for the geologist on road-building projects. --M. Russell.

Seeger, Ralph W. Highway Material Survey in West Virginia, p. 57-66.

West Virginia's share of the federal Interstate Highway System is about 300 mi. Average rock quantities for base course construction will be about 31,000 cu. yds. per mi. The West Virginia Highway Aggregate Research Project is being undertaken to locate and show by means of a special map, supplies of potential aggregate in the state. Photogrammetry is used extensively. The suitability for aggregate of the most common rocks of West Virginia is summarized on the basis of their geologic age and lithology. An observed relationship between a rock's ability to absorb liquid and percent wear by the Los Angeles Abrasion Test is the basis of a quick field method



of predicting L. A. A. T. results.--M. Russell.

Fletcher, G. A. *Geology in Foundation Engineering*, p. 67-73.

Civil engineers should avail themselves of geologic advice in planning foundations for heavy structures. The presence of highly surcharged and glacially overconsolidated organic deposits found during the driving of pilings in Milwaukee, Wisconsin, would have been discovered at less cost had a geologic reconnaissance been made. The presence of an extremely irregular bedrock surface under an industrial site in Childersburgh, Alabama, could similarly have been discovered by geologic survey before costly foundation revisions were necessary. Construction projects which have benefited from prior knowledge of geologic conditions include eastern Manhattan Island, New York, where it was found that overconsolidated glacial lake deposits would support friction pilings, thus eliminating the need to drive to bedrock, and Georges Bank off New England, where foundation construction for the Texas Tower was simplified by prior knowledge that the bottom consisted of glacial deposits.--M. Russell.

Belcher, Donald J. *Applied Geomorphology*, p. 83.

In order to refine geomorphology to the point where highway engineering requirements are served, a fourth order of 36 geomorphic units are defined, and a fifth order of micro-forms or micro-features is recognized. The role of geomorphic analysis in locating materials and indicating superior road alignments is illustrated by examples of work for the West Side Freeway, San Joaquin Valley, California; the Sumpter, South Carolina, by-pass; Virginia highway paralleling U.S. 1, S. of Washington, D.C.; the Massachusetts Turnpike, northern extension; and a Pennsylvania project. Photogrammetry is used extensively in this work.--M. Russell.

1614. Widmer, Kemble. *GEOLOGICAL PROBLEMS IN THE CONSTRUCTION OF DAMS*: New York Acad. Sci., Trans., v. 22, no. 4, p. 223-232, maps, secs., Feb. 1960, 3 refs.

Engineering geology is the application of all phases of geology to engineering problems. Geologic investigation of 4 dam sites in New Jersey required the use of regional and detailed structural geology, stratigraphy, petrography, petrology, geophysics, airphoto interpretation, paleontology, ground-water geology, geomorphology, historical geology, and Pleistocene geology.

At an unidentified dam site, the spacing, continuity, and direction of faults and jointing in quartzite are being determined. The faults and joints are potential zones of leakage into adjacent valleys that are several hundred feet lower than the reservoir.

At the Charlotteburg reservoir, joints and buried channels are potential zones of leakage. The directions of the joints were plotted on a Schmidt net to determine the best angle to drill grout holes in intensely folded and faulted Precambrian gneiss. A Pleistocene geologist determined the direction of ice movement in order to locate buried channels.

Geophysical surveys in Round Valley showed that diabase was a ring dike instead of a cone sheet. At this site a core sample identified as limestone sagged downward in a fault was reidentified petrographically as a badly fractured gneiss rehealed by calcite.

A dam and reservoir site on Spruce Run required

intensive investigation because of complicated faulting and folding. The stratigraphy was determined outside the project area, and key beds were traced into the site. Faults were determined by airphoto interpretation. Depth to sound rock was found by geophysical methods. A ground-water survey included location of the water table and examination of wells and springs. Two different ages of solution cavities were found by examination of core samples.--R. Van Horn.

2-1615. Rich, George R. *THE NIAGARA PROJECT-DESIGN*: Boston Soc. Civil Engineers, Jour., v. 47, no. 1, p. 1-17, Jan. 1960.

This power project, whose various phases lie within an area of 6 x 3 mi., consists of 1) a water intake 3 mi. above the Falls on the U.S. side; 2) waterways; 3) a pump-turbine generating plant with storage reservoir; 4) a conventional generating plant; 5) power transformation and transmission facilities, as well as parkway facilities.

The waterways consist of 2 horseshoe-shaped conduits of reinforced concrete 46 ft. wide and 66 ft. high, constructed by cut and cover methods in the calcareous beds of the Lockport formation. These will carry water from the intake to the forebay/afterbay of the Tuscarora pump-turbine generating plant.

The rocks at the site are well-bedded sediments dipping to the S. at 30 ft. per mi. They are jointed but not faulted. The upper beds are of the Lockport formation, containing numerous joints whose bedding planes are opened, and have been further affected by solution activity. These occur beneath thin clay, gravelly sand, and sandy silt soils which are 4 to 32 ft. thick along the waterway.

To insure against slides in the escarpment face at the Lewiston generating plant 1) excavation is made at an angle of 45°; 2) a grout curtain is provided well upstream of the intake mat and extending into the shale strata (the Rochester, Clinton, Power Glen, and Queenston shale formations), interbedded in calcareous formations; 3) a system of drains is installed to collect any possible percolation not intercepted by the grout curtain.

Maps of the various elements of the project are provided; their design and seismic loading for structures is discussed.--J. W. Skehan.

2-1616. Hall, W.M. *THE NIAGARA POWER PROJECT - ADMINISTRATION AND CONSTRUCTION MANAGEMENT*: Boston Soc. Civil Engineers, Jour., v. 47, no. 1, p. 18-30, Jan. 1960.

Includes a detailed account of construction problems encountered and the ways in which they were solved, particularly those arising from urban and industrial congestion. Problems of relocation of railroads and electrical transmission facilities, removal of 10 million cu. yds. of material from the side of the gorge, and construction problems presented by the steep slope on the side of the gorge are described.--J. W. Skehan.

2-1617. Praszker, Michael. *DEVELOPMENT OF MARGINAL LANDS IN SAN FRANCISCO*: Civil Engineering, v. 30, no. 2, p. 68-71, 5 illus., 2 diags., Feb. 1960.

Disastrous earth slides triggered by heavy rains in the hilly areas of San Francisco have forced the home building industry to use sound exploration, soil mechanics, and grading principles and procedures. In the West Twins Peak subdivision, exploration

tion revealed several bedrock ravines overlain by eolian sand and alluvial fine-grained soils containing lenses of coarse, water-bearing material. Bedrock is weathered arkosic sandstones, thin-bedded cherts, and basic intrusives. Active earth movement was discovered in a large ravine which was to receive fill to a depth of 100 ft. The bottom of the ravine was excavated, and cuts were made 5 ft. deep in bedrock to provide shear keys and shoulder drains. At the head of the ravine a slip zone of puttylike clay underlain by springs was located and subsequently excavated to form a series of parallel drainage ditches normal to the slope. In cuts that showed imminent signs of failure, water-bearing sandy layers were exposed which could have supplied water to initiate and lubricate slides. To assure rapid drainage of the slopes, holes were drilled in back of the head of the cut to the toe of the cut and filled with gravel. The final product is a terraced system which has withstood several "wet winters" and a sharp earthquake.

Exploration in topographically rugged Mira Loma Park subdivision revealed soft clay layers and rock terraces filled with alluvium. The material was removed, drains were emplaced, and the site was then filled, except where final grade was close to bedrock. In those places, concrete piers were emplaced from just above the soft layers to rock.-- R. E. Pendergast, Jr.

2-1618. Kuiper, E. SEDIMENT TRANSPORT AND DELTA FORMATION: *Am. Soc. Civil Engineers, Hydraulics Div., Jour.*, v. 86, no. HY2, pt. 1, p. 55-68, 2 maps, sec., diag., Feb. 1960.

This paper deals with sedimentation computations that were carried out in connection with dike design for the Saskatchewan Delta Reclamation Project. The first part of the paper gives a description of the regime of the Saskatchewan River in general and of the Saskatchewan delta in particular. The second part presents some observations on delta formation. The third part discusses the sedimentation problems that were encountered in the design of dikes. The fourth and final part presents the method that was adopted to determine the future rate of sedimentation on the floodplains adjacent to the projected dikes.-- Auth.

2-1619. Hansen, J. A., Donald Thomas Davidson, and Chalmer J. Roy. GEOLOGIC AND ENGINEERING PROPERTIES OF TILL AND LOESS, SOUTHEAST IOWA. PROGRESS REPORT, NOVEMBER 1, 1959: 48 p., 15 figs. incl. illus., maps, graphs, 5 tables, Ames, Iowa, Iowa Engineering Experiment Station, 1959, 37 refs.

Readily available sources of crushed aggregate suitable for road construction occur within only a small portion of Iowa's total acreage. Federal and state highway construction programs will impose great demands for road and paving materials, and some suitable material must be found to substitute for the aggregate. The Iowa Engineering Experiment Station, Iowa State College, is engaged in a study of

the loess and glacial till materials which have been estimated to cover 80% of the state. The field work upon which this report is based was conducted in the lower 3 tiers of counties in southeastern Iowa. The investigation is Project HR-1 of the Iowa Highway Research Board, and Project 283-S of the Iowa Engineering Experiment Station.

The loess in southeastern Iowa is present as a widespread deposit which blankets the terrain. It has been removed by the process of erosion from steeply rolling landscapes and from the alluvial flood plains. Actually it is doubtful if deposits of any significance ever formed on these flood plains, as removal should have been contemporaneous with deposition.

Marked variations in clay content occur in the solum of loess-derived soils as a direct result of the differences in soil-forming factors such as slope and vegetation. The C horizons are not as highly influenced by these factors and are thus more uniform in texture, color, and clay content. C horizons studied have all been leached and are noncalcareous. Slight variations in clay content are apparent in E-W. traverses and possibly reflect a process of eolian sorting.

Glacial till is encountered everywhere beneath the loess and is extensively exposed in rolling topography. The soil-forming factors which play an important part in the development of the loess-derived solum also act similarly in till horizons. Due to the greater thickness of till deposits and to the thickness of the overlying loess, leaching has progressed downward to an average of 6 or 7 ft. Below this depth tills are highly calcareous.

The sandy silt layer which underlies the loess and which sometimes forms a part of the till solum exhibits strongly bimodal sorting characteristics. It has undoubtedly originated under conditions of slope and rill wash contemporaneous with the initial increments of loess deposition.

Studies of particle-size distribution seem to be prerequisite to any attempt at soil stabilization. Well-graded soils such as C horizon glacial till are stabilized more easily by either mechanical compaction or by the addition of stabilizing agents than, for example, the poorly graded gumbotils. High-clay soils require significantly higher amounts of stabilizing agents than well-graded sands. It can thus be readily seen that prior knowledge of particle size distribution aids in the choice of the most durable and economic method of stabilization.

Clay minerals noted in this study were montmorillonite, illite, and kaolinite. Montmorillonite has the highest cation exchange capacity of these clays, and it may be that in certain kinds of chemical stabilization this high exchange capacity would permit a greater bonding action.

X-ray analysis of the detailed study samples is an extremely useful technique for the qualitative determination of the varied mineral types. Quantitative determinations are still in the trial and error stage, but rough estimates may be made from inspection of peak intensities on the strip-chart record. No distinctive criteria for the distinction of different aged tills were observed.--From auth. introd. & summ.



## 15. MISCELLANEOUS

1620. Southern Research Institute. **THE UN-  
DISCOVERED EARTH. PROCEEDINGS OF THE  
CONFERENCE, JUNE 11-12, 1959:** 56 p., illus.,  
pts., map, diags., graphs, tables, Birmingham,  
Alabama, 1959, refs.

The Southern Research Institute was the sponsor  
of this conference held June 11-12, 1959, at which  
specialists in several of the earth sciences gave  
lectures designed to inform nonspecialist research-  
ers, engineering executives, and businessmen, of  
the present status of research in their various fields.  
The speakers and lectures are as follows:

Martin, Thomas W. Opening Remarks, p. 8.

Kinzel, A.B. Keynote Address, p. 9-15.

English, Earl F. Harnessing Geothermal Heat,  
p. 16-18.

The results of study and experience in known  
fields of production and utilization of geothermal  
heat for steam and power.

Iselin, Columbus O'D. Problems of Marine Re-  
sources, p. 19-21.

Report on progress being made to exploit  
biological and chemical resources of the oceans.

Davidson, D.M. Tapping the Earth, p. 22-24.

Report on progress being made in mining  
methods and exploration.

Hibbard, W.R., Jr. Tomorrow's Materials, p.  
25-27.

The developments in metallurgy are cited to  
illustrate that new materials are constantly  
sought.

Hubbert, M. King. Energy Sources for the Next  
Fifty Years, p. 28-31.

Predictions are made on the probable require-  
ments for fossil fuels and nuclear energy.

Teller, Edward. Project Plowshare, p. 34-38.

Prospects for the peaceful use of atomic  
explosions.

Woollard, George P. The Crust and Core of the  
Earth, p. 39-44.

Review of current theory on the crust and  
core and methods for investigating them.

Leggette, R.M. Water for an Expanding World,  
p. 45-48.

Discussion of the water supply problems and  
methods involved in solving them.

Lill, Gordon G. The Deep Hole Project, p. 49-  
52.

A plan to drill a hole to the Mohorovičić Dis-  
continuity.

Nolan, Thomas B. Technology and the National  
Minerals Problem, p. 53-55.

Discussion of the basic and applied research,  
development and economic climate required to  
insure an adequate supply of mineral raw mate-  
rials.

Martin, Thomas W. Conference Summary, p. 56.

--M. Russell.

1621. Dart, Raymond A., with Dennis Craig.  
**ADVENTURES WITH THE MISSING LINK:** 255 p.,  
illus., New York, Harper & Brothers; London, Ham-  
ish Hamilton, 1959, 71 refs.

A personal and historical record from the original  
discoverer of *Australopithecus africanus* at Taungs  
in Bechuanaland of the circumstances surrounding

that discovery in 1924 and the mixed reception it  
received as a result of the prevalent anthropological  
opinions of that era.

The author then traces the sequence of dramatic  
discoveries of adult, adolescent, and infantile man-  
apies by the late Dr. Robert Broom and his assistant  
in the 3 Sterkfontein valley sites W. of Johannesburg  
and by himself and his associates at the famous  
Limeworks site in Makapansgat valley, central  
Transvaal.

Explaining the mysteries to be solved and the  
obstacles to be overcome, he shows how the steadily  
accumulating evidence during the past quarter of a  
century has overwhelmed opposition to the recognition  
of the Australopithecines, not only as the physical  
but also as the cultural representatives of proto-  
humanity, and has thereby completely transformed  
previous concepts of the course of human emergence  
and the factors provocative thereof.--R. A. Dart.

2-1622. Green, Jack, and Dael Wolfle. **GEA,  
DAUGHTER OF CHAOS:** Science, v. 131, no. 3407,  
p. 1071, Apr. 15, 1960.

It is obvious that the science we call geology, and  
many of the related "geo" sciences will be applied  
with increasing frequency to planets, moons, and  
other astronomical bodies. Purists hold that the  
root "geo" means earth, and hence new terms must  
be invented to apply to each new body being investi-  
gated, for example "selenology" to apply to study of  
the moon. To insist on this is unnecessary and un-  
wise; the principles of the science will be the same  
wherever they are practiced, or essentially so, and  
the name applied to the science should remain the  
same also. It is only necessary that the "geo" term  
so applied be followed by the word "of" when it is  
used in reference to an area other than earth. Thus  
for example, one should refer to "the geology of  
Venus."--M. Russell.

2-1623. Kartashov, I.P. **CERTAIN GEOLOGIC-  
GEOGRAPHIC TERMS:** Akad. Nauk SSSR, Bull.,  
Geol. Ser., in translation, 1958, no. 1, p. 65-68,  
pub. 1959, 10 refs.

The meaning and current usage of the following  
terms are discussed: denudation, destruction, weath-  
ering, deluvium, erosion, washing out, deposition.  
It is concluded that a review and unification of the  
geologic-geographic nomenclature is needed.--M.  
Russell.

2-1624. Curtis, Neville M., Jr. **PUBLISHED  
PAPERS ON OKLAHOMA GEOLOGY IN THE YEAR  
1959:** Oklahoma Geology Notes, v. 20, no. 3, p. 55-  
73, March 1960.

Annotated bibliography and index of 198 papers  
published in 1959, and concerned with the geology of  
Oklahoma.--Auth.

2-1625. National Advisory Committee on Research  
in the Geological Sciences, Ottawa. **NINTH ANNUAL  
REPORT 1958-59 (INCLUDING SURVEY OF CUR-  
RENT RESEARCH IN THE GEOLOGICAL SCIENCES  
IN CANADA, 1958-59):** 216 p., 6 tables, pub. 1960,  
refs.

The National Advisory Committee on Research in  
the Geological Sciences has a threefold purpose: to  
stimulate and coordinate geological research carried  
on in Canada; to suggest research projects that  
should receive attention; and to aid in having these

projects undertaken. Its function is to stimulate research by the universities, the federal and provincial departments of mines, and by other organizations equipped for the work.

The first part of the report gives a summary of the work of the Committee in the period Sept. 1, 1958, to Aug. 31, 1959. The second contains the reports of the subcommittees covering the different fields of the geological sciences. These record developments in 1958-1959 and suggest some further problems for study.

The report includes the annual survey of current geological research in Canada. This records information on research by the universities, federal and provincial departments of mines, research councils and foundations.--From introd.

2-1626. Muilenburg, Grace. **RESEARCH AND ACTIVITIES OF STATE GEOLOGICAL SURVEY OF KANSAS, FISCAL YEAR ENDING JUNE 30, 1959:** Kansas, State Geol. Survey, Misc. Rept., 40 p., 22 illus., 5 maps, diag., 8 graphs, tables, Jan. 1960, 21 refs.

The organizational plan of the Survey is outlined, and present and proposed projects of each division are discussed briefly. Topographic maps, Survey reports, and papers published elsewhere by Survey personnel are listed.--R. H. King.

2-1627. Alexandrov, Eugene A. **RED CHINA STEPS UP ITS GEOLOGICAL SERVICE:** Mining Engineering, v. 12, no. 3, p. 248-250, March 1960, 5 refs.

The article, based on 5 recent articles from the Soviet Union, surveys the mineral resources of China and recent stepped-up exploration activities. Administrative and educational activities in connection with mineral resources and exploration are covered.--Auth.

2-1628. **MAPPING IN ANTARCTICA:** Military Engineer, v. 52, no. 346, p. 144, March-Apr. 1960.

The U. S. Geological Survey's Topographic Division, with funds provided by National Science Foundation, is conducting a limited program of mapping in Antarctica. The activities are: the preparation of aerial photographic and flight specifications, and selection of areas to be flown, by a Navy squadron; establishment of positions and identification of control stations by field personnel; and the preparation of maps. Seven areas chosen for their scientific interest are: Thurston Peninsula, Sentinel Mountains, Executive Committee Mountains, Horlick Mountains, Queen Maud Range, mountains of South Victoria Land, and the mountains W. of the Ross ice shelf. Three maps of the Knox Coast at 1:500,000 are being completed on a deferred project. A map of McMurdo Sound, Ross Island, and neighboring mountains at 1:250,000 is to be prepared, as well as local areas at larger scale for geologic studies.--A. C. Mason.

2-1629. **LUNAR MAPPING:** Military Engineer, v. 52, no. 345, p. 62-63, Jan.-Feb. 1960.

The U. S. Army Map Service is making what is believed to be the first map of the moon by modern stereoscopic principles. Libration of the moon permits selection of pairs of photographs having angular differences up to  $16^\circ$ , sufficient for stereoscopic vision, although in earth mapping angular parallax of  $30^\circ$  to  $40^\circ$  is used. No method is feasible to make

presently available lunar photographs taken with astronomic cameras geometrically compatible with photogrammetric stereoplotters. Currently accepted earth-mapping standards cannot be achieved; the lunar map will be at a scale of 1:5,000,000 with relief depicted by large-interval contour lines supplemented by form lines, hachures, shaded relief, scarp symbols, and spot elevations. Mapping at a scale of 1:1,000,000 is anticipated as soon as mapping methods are perfected and photographic coverage specifications is acquired. Mapping at a scale of 1:250,000 may be possible by photographing the moon from aerial balloons. Atmospheric refraction is eliminated in photographs taken from artificial earth satellites, and maps at 1:100,000 are believed possible. The hidden part of the moon cannot be mapped until a moon-orbiting satellite is equipped to transmit suitable imagery to earth.--A. C. Mason.

2-1630. **PROJECTION FOR LUNAR MAP:** Military Engineer, v. 52, no. 345, p. 63, diag., Jan.-Feb. 1960.

A lunar map projection must not omit or distort polar areas and must show about 60% of the lunar surface without misleading distortion. Of presently used projections, the Lambert Azimuthal projection most nearly meets requirements, but it also produces undesirable distortion of shapes of features toward the edges. For its lunar map, the U. S. Army Map Service has modified the stereographic projection to suit the above requirements. The projection plane is tangent to the mean libration center, and rays are projected to it from a point approximately  $7/26$  of a diameter beyond the rear surface of the moon on a line normal to the plane of tangency. This projection reduces the scale and shape distortions to an acceptable compromise. Maximum radial displacement is 1.1, and maximum tangential scale distortion is 1.85 at the outer limits.--A. C. Mason.

2-1631. **LUNAR TERRAIN STUDY:** Military Engineer, v. 52, no. 346, p. 144, March-Apr. 1960.

The Military Geology Branch of the U. S. Geological Survey is making a terrain study of the surface of the moon in a joint project in which the U. S. Army Map Service is preparing a map of the moon to serve as a base for the study. The first terrain study will be at a scale of 1:5,000,000, and this is expected to be followed by a more detailed study at 1:1,000,000. An analysis of the lunar terrain is being made by photogeologists stereoscopically viewing selected libration pairs of lunar photographs. The surface of the moon will be divided into physiographic regions and subregions, and the distinguishing characteristics and principal physical features of each will be described. The study will include interpretation of the constituents, textures, and bearing power of the surface.--A. C. Mason.

2-1632. **Barnard, Richard H. TERRAIN FACTORS IN AIRBORNE OPERATIONS:** Military Engineer, v. 52, no. 345, p. 48-49, 3 illus., Jan.-Feb. 1960.

Sites are chosen by applying terrain intelligence to tactical needs. Maximum acceptable slopes are 2 1/2% longitudinal grade and slightly less for transverse grade for fixed-wing assault aircraft, 30% for paratroops, but 10% preferable, and 50% for helicopters, but 10% preferable. Rough ground, such as gullies, boulders, and bare rock, severely limit fixed-wing aircraft and paratroop landings, and even large-scale nonlanding helicopter operations. Soil



earing capacity is a factor for fixed-wing aircraft; it varies with weather, irrigation, and drainage conditions even for the same material. Some soils, s lateritic clay, cannot sustain repeated impacts from landing. The density of vegetation affects airborne operations; in some cases, as paddy fields, the associated ground conditions can be a greater obstacle than the vegetation itself. The various terrain factors require mapping in advance to enable rapid planning of airborne operations in emergencies. --A. C. Mason.

-1633. Whitmore, Frank C., Jr. TERRAIN INTELLIGENCE AND CURRENT MILITARY CONCEPTS: *Am. Jour. Sci.*, v. 258-A (Bradley Volume), p. 375-387, 1960, 33 refs.

The development of nuclear weapons for tactical use has been followed by a concept of warfare based on small, mobile combat units equipped with helicopters, aircraft which can land and take off on short, sketchily prepared runways, and low ground-pressure vehicles which can move across country to supply a rapidly shifting force. This requires that terrain intelligence, dealing with the effects of the ground on military construction and maneuver, be made available to forward units that may have to act quickly on their own initiative. Help in this effort will come from new means of gathering information, such as radar, infrared photography, and television. Use of statistical methods will accelerate terrain analysis, especially at the strategic level. Analysis

of landforms as they offer shelter from nuclear blast, and of soils which may emit neutron-induced gamma radiation, will assume an important place in terrain intelligence, as also will selection of safe routes for helicopters, tanks and supply vehicles; sites for rapid airfield construction and paratroop landings; and the suitability of soil for hasty excavation of shelters. The rapid pace of such warfare will require frequent preparation of situation maps showing areas where tree blow-down might slow troop movement or where nuclear contamination might be dangerous.--Auth.

2-1634. Baranov, V.I., and K.G. Knorre. THE SEVENTH SESSION OF THE COMMITTEE ON DETERMINATION OF THE ABSOLUTE AGE OF GEOLOGIC FORMATIONS: *Geokhimiya* [in translation], 1958, no. 5, p. 646-647, pub. 1959.

A brief summary of titles and subjects covered at the session, held in May 1958, at the Division of Geologic-Geographic Sciences of the Academy of Sciences, U.S.S.R. Reports included: a summary of the Precambrian geochronology of the Baltic shield by A. A. Polkanov, a report on the Precambrian history of the planet, by N. P. Semenenko, a first report on the application of the K/Ca dating method in the U.S.S.R., by N. I. Plevaya, and a communication of a new K-capture constant for  $K^{40}$  by E. K. Gerling. The use of the new constant in the Ar method leads to higher absolute ages than use of the formerly accepted constant.--F. Manheim.

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